

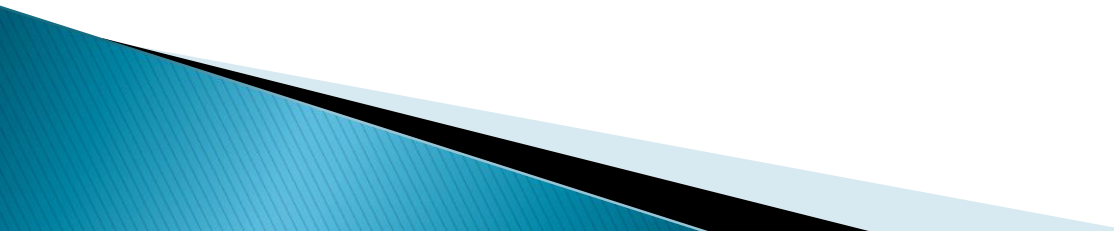
Training Institute on Adaptive Water– Energy Management in the Arid Americas



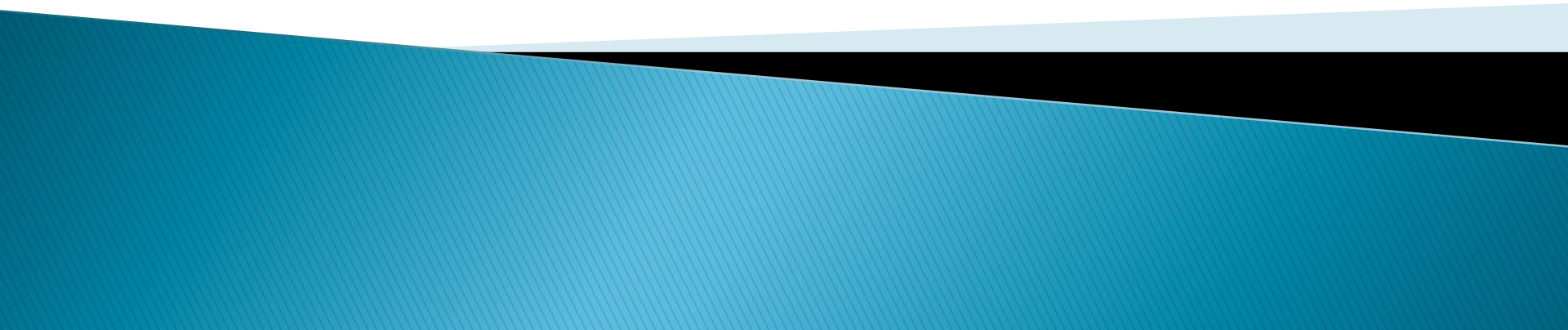
Climate change/variability, water and
energy, and adaptation

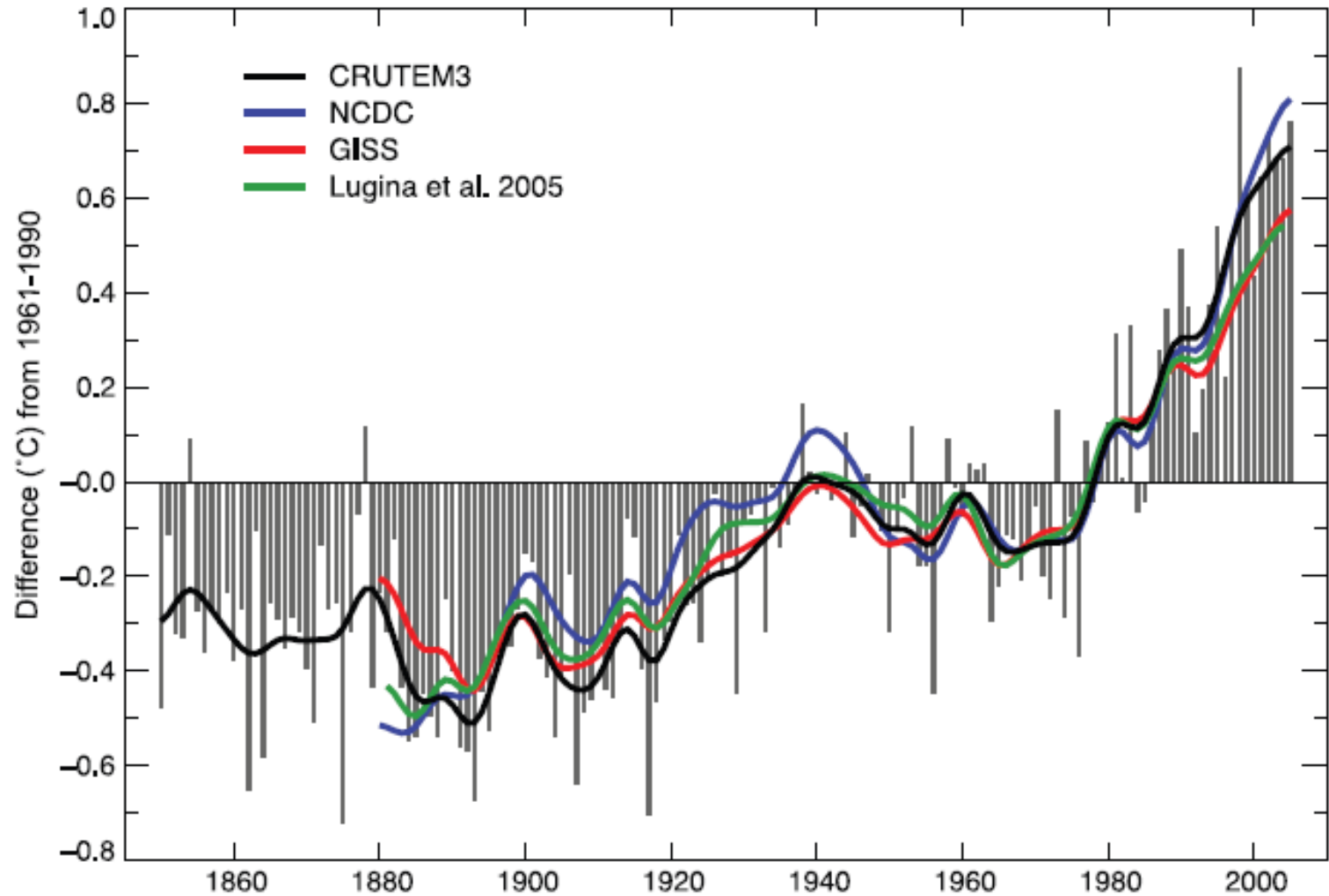
Alfredo Ribeiro
Federal University of Pernambuco

Topics

- ▶ IPCC Scenarios and Global Circulation Models
 - ▶ Bias Correction
 - ▶ Water Balance Simulation
 - ▶ Hydrological Model
 - ▶ Simulation and Optimization
 - ▶ Exercise
- 

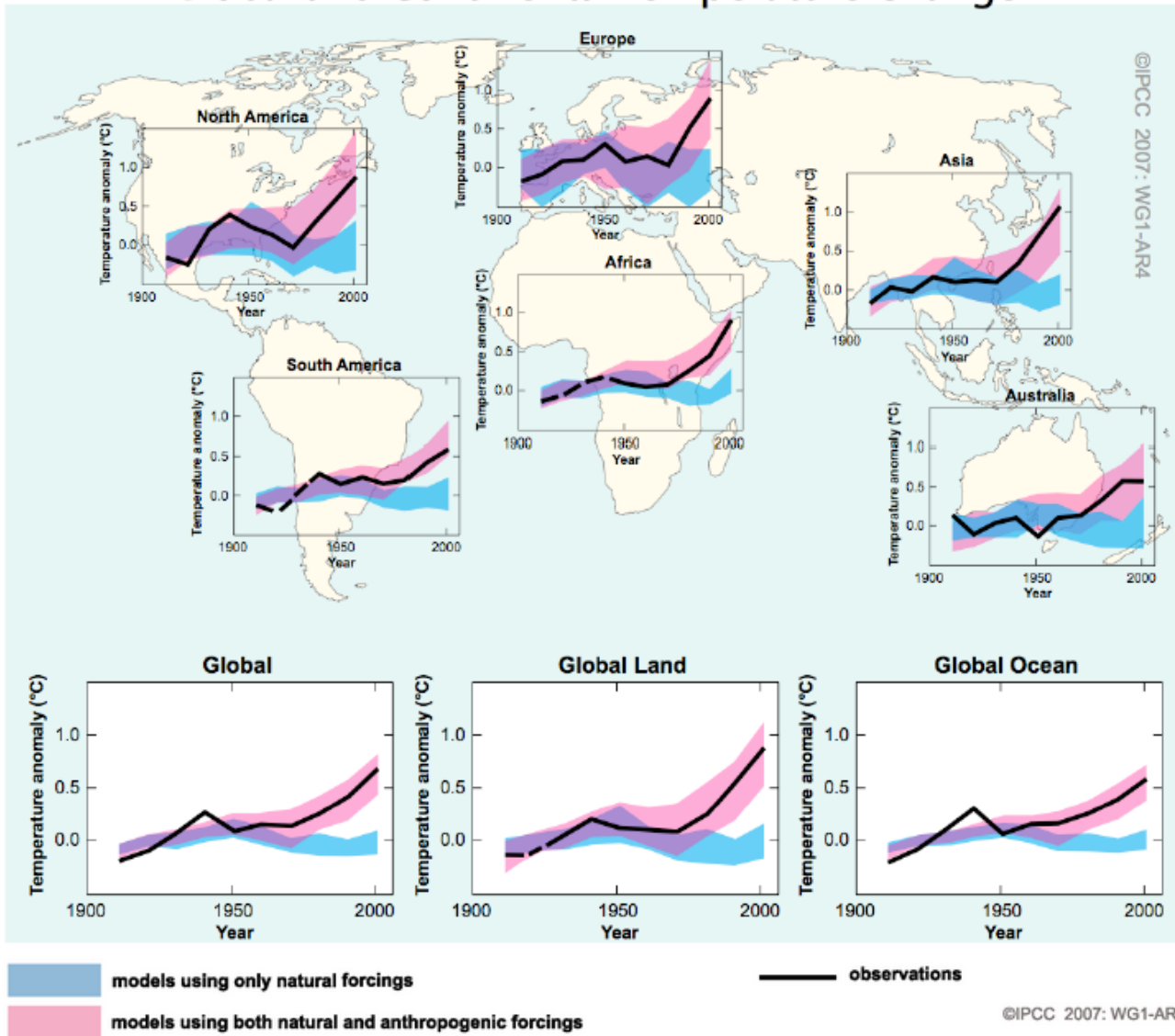
IPCC Scenarios and Global Circulation Models





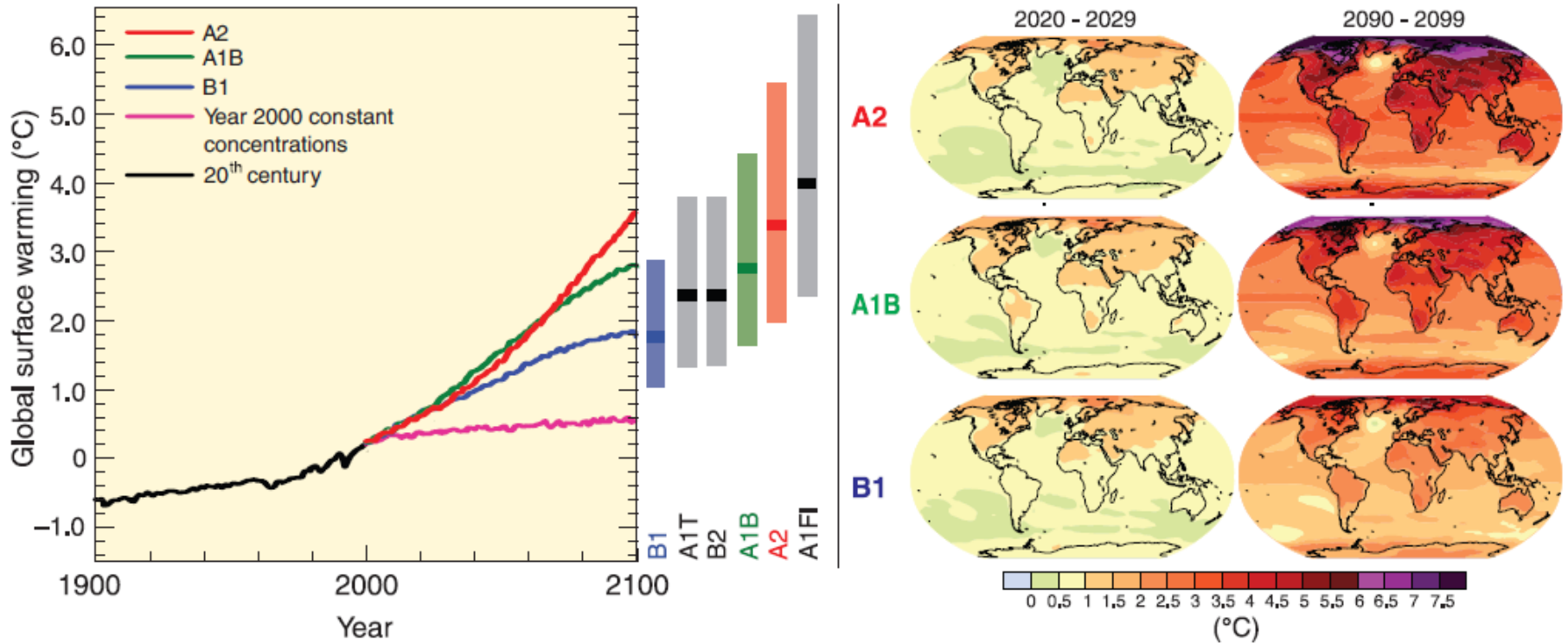
Annual anomalies of global land-surface air temperature (°C), 1850 to 2005, relative to the 1961 to 1990 mean

Global and Continental Temperature Change



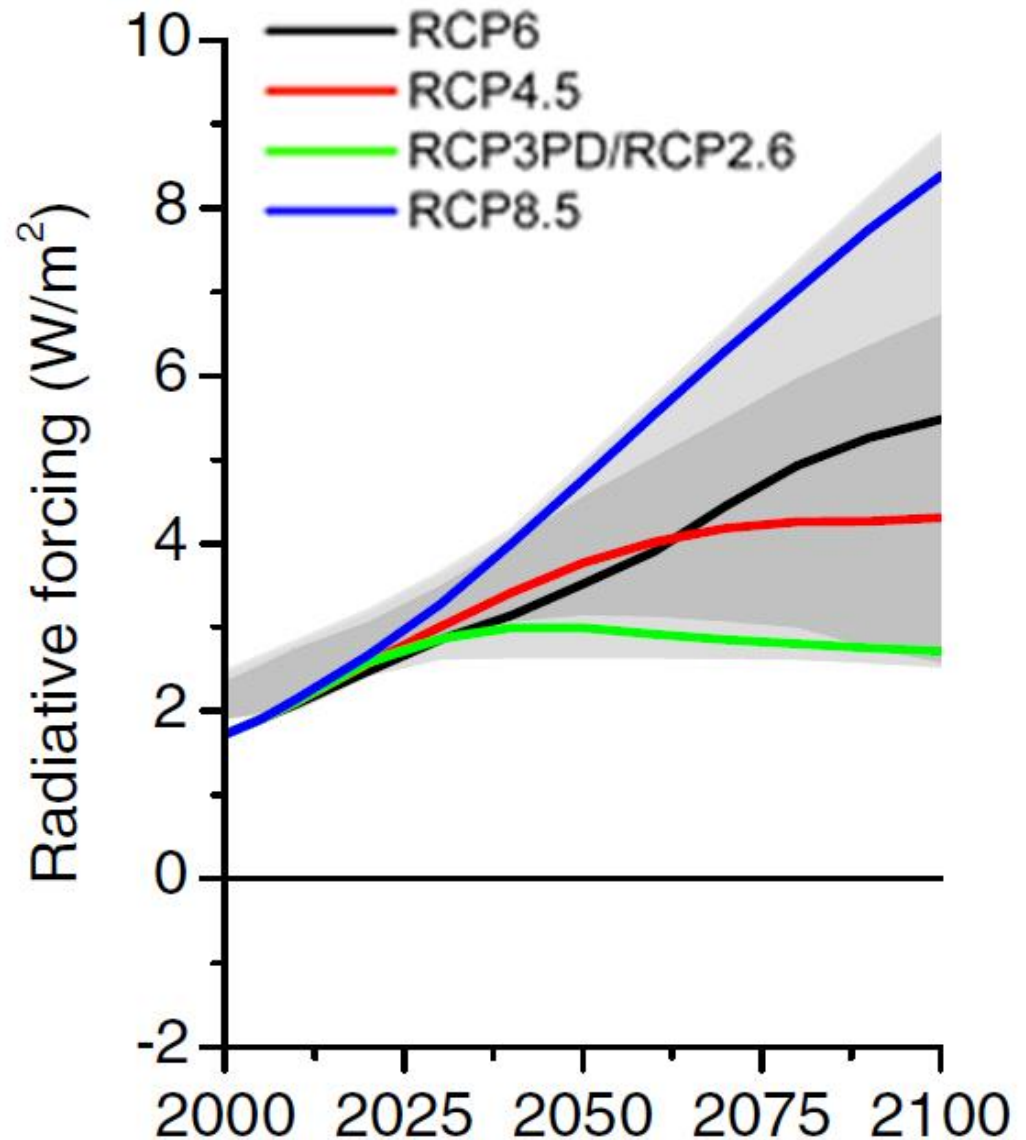
AR4-SRES Scenarios

Atmosphere-Ocean General Circulation Model projections of surface warming



AR5-RCP

Four RCPs were selected and defined by their total radiative forcing (cumulative measure of human emissions of GHGs from all sources expressed in **Watts per square meter**) pathway and level by 2100. The RCPs were chosen to represent a broad range of climate outcomes, based on a literature review, and are neither forecasts nor policy recommendations.



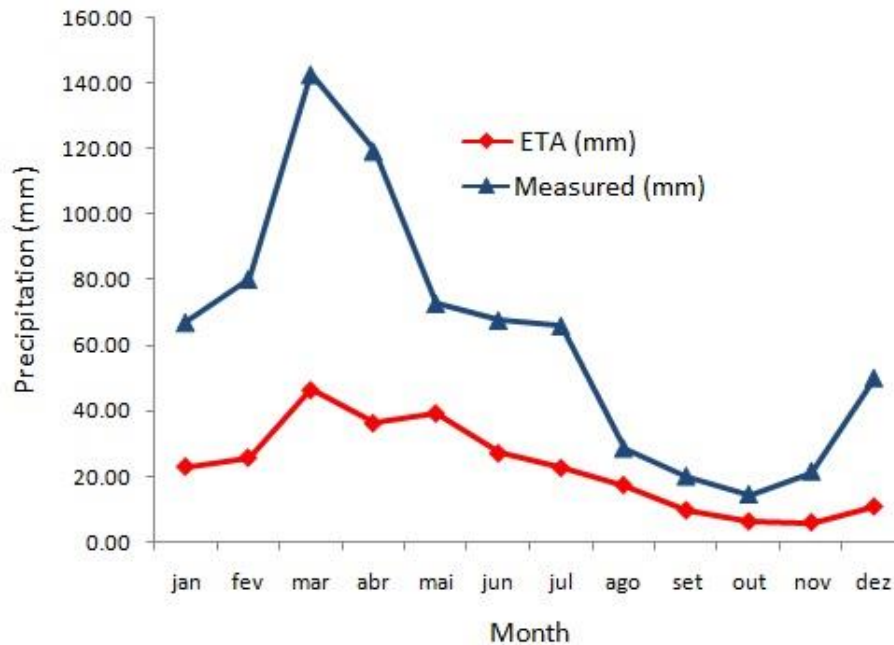
Bias Correction



Bias correction

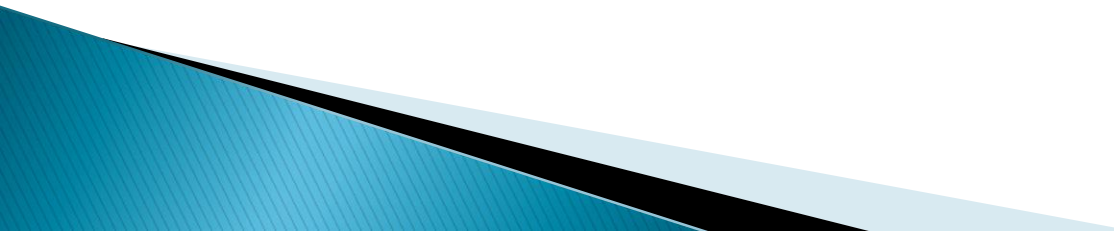
Eta CCS precipitation underestimate

Correction with cumulative density function (CDF)



Bias correction

Correction with cumulative density function (CDF)

1. For each grid point and for each month (January-December), it is necessary to compute the cumulative frequency of the model and observed rainfall;
 2. The second step is to determine the frequency of the rainfall model, and then replace the raw value with the amount of rainfall observed associated with the matching cumulative frequency;
- 

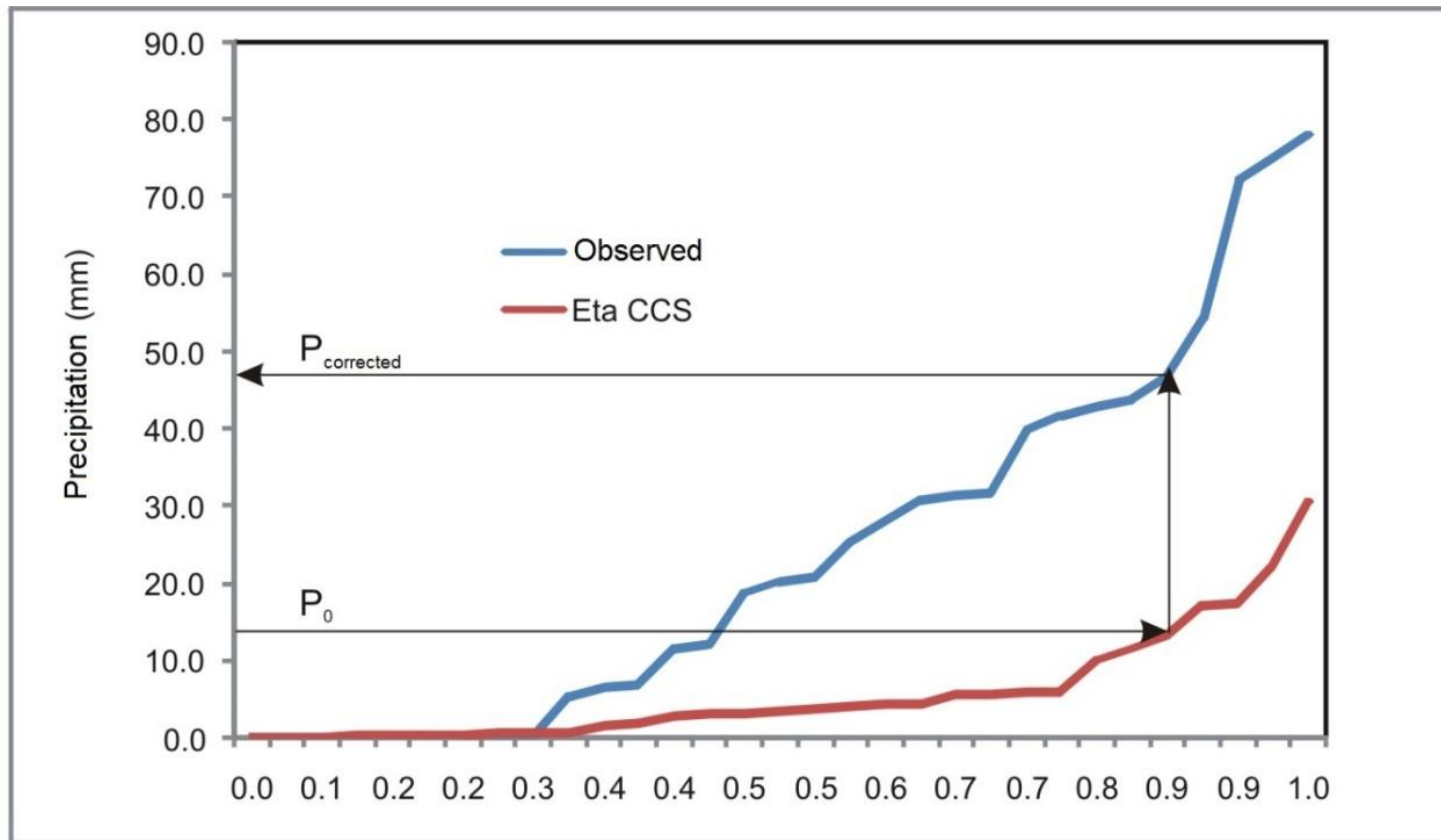
| Year | P (mm) | Order | P (mm) | F (%) | Year | P (mm) | Order | P (mm) | F (%) |
|------|--------|-------|--------|-------|------|--------|-------|--------|--------|
| 1960 | 27.5 | 1 | 12.97 | 3.23 | 1976 | 27.51 | 17 | 33.27 | 54.84 |
| 1961 | 77.68 | 2 | 13.59 | 6.45 | 1977 | 19.83 | 18 | 34.32 | 58.06 |
| 1962 | 21.47 | 3 | 13.82 | 9.68 | 1978 | 15.66 | 19 | 35.71 | 61.29 |
| 1963 | 13.99 | 4 | 13.99 | 12.90 | 1979 | 113.98 | 20 | 36.46 | 64.52 |
| 1964 | 70.78 | 5 | 15.66 | 16.13 | 1980 | 18.04 | 21 | 38.76 | 67.74 |
| 1965 | 16.67 | 6 | 16.63 | 19.35 | 1981 | 33.27 | 22 | 70.78 | 70.97 |
| 1966 | 16.63 | 7 | 16.65 | 22.58 | 1982 | 35.71 | 23 | 77.68 | 74.19 |
| 1967 | 12.97 | 8 | 16.67 | 25.81 | 1983 | 16.65 | 24 | 100.61 | 77.42 |
| 1968 | 133.39 | 9 | 16.87 | 29.03 | 1984 | 134 | 25 | 112.54 | 80.65 |
| 1969 | 194.41 | 10 | 18.04 | 32.26 | 1985 | 18.79 | 26 | 113.98 | 83.87 |
| 1970 | 100.61 | 11 | 18.79 | 35.48 | 1986 | 36.46 | 27 | 117.98 | 87.10 |
| 1971 | 29.27 | 12 | 19.83 | 38.71 | 1987 | 34.32 | 28 | 133.39 | 90.32 |
| 1972 | 117.98 | 13 | 21.47 | 41.94 | 1988 | 13.82 | 29 | 134 | 93.55 |
| 1973 | 270.72 | 14 | 27.5 | 45.16 | 1989 | 13.59 | 30 | 194.41 | 96.77 |
| 1974 | 16.87 | 15 | 27.51 | 48.39 | 1990 | 112.54 | 31 | 270.72 | 100.00 |
| 1975 | 38.76 | 16 | 29.27 | 51.61 | | | | | |

$$F = \frac{m}{n}$$

m = order

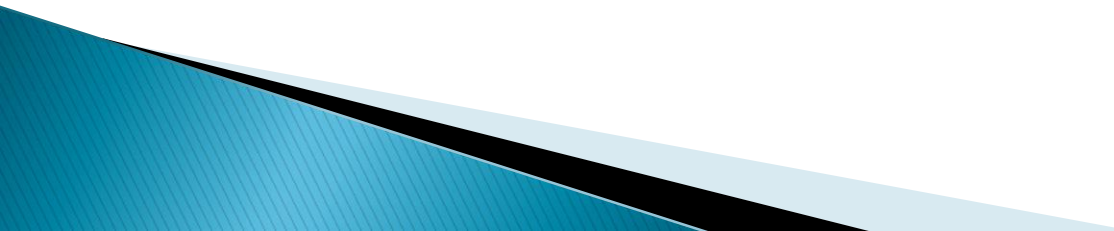
n = number of years

Bias correction



Bias correction

Correction with cumulative density function (CDF)

3. In the case of simulations of the future, instead of directly using CDF corresponding to the observed rainfall, the method first identifies the future rainfall value in the CDF of the model in the present time;
 4. The correction is made matching the quantile found in the CDF of the model to the same value in the CDF of the observed rainfall.
- 

Bias correction

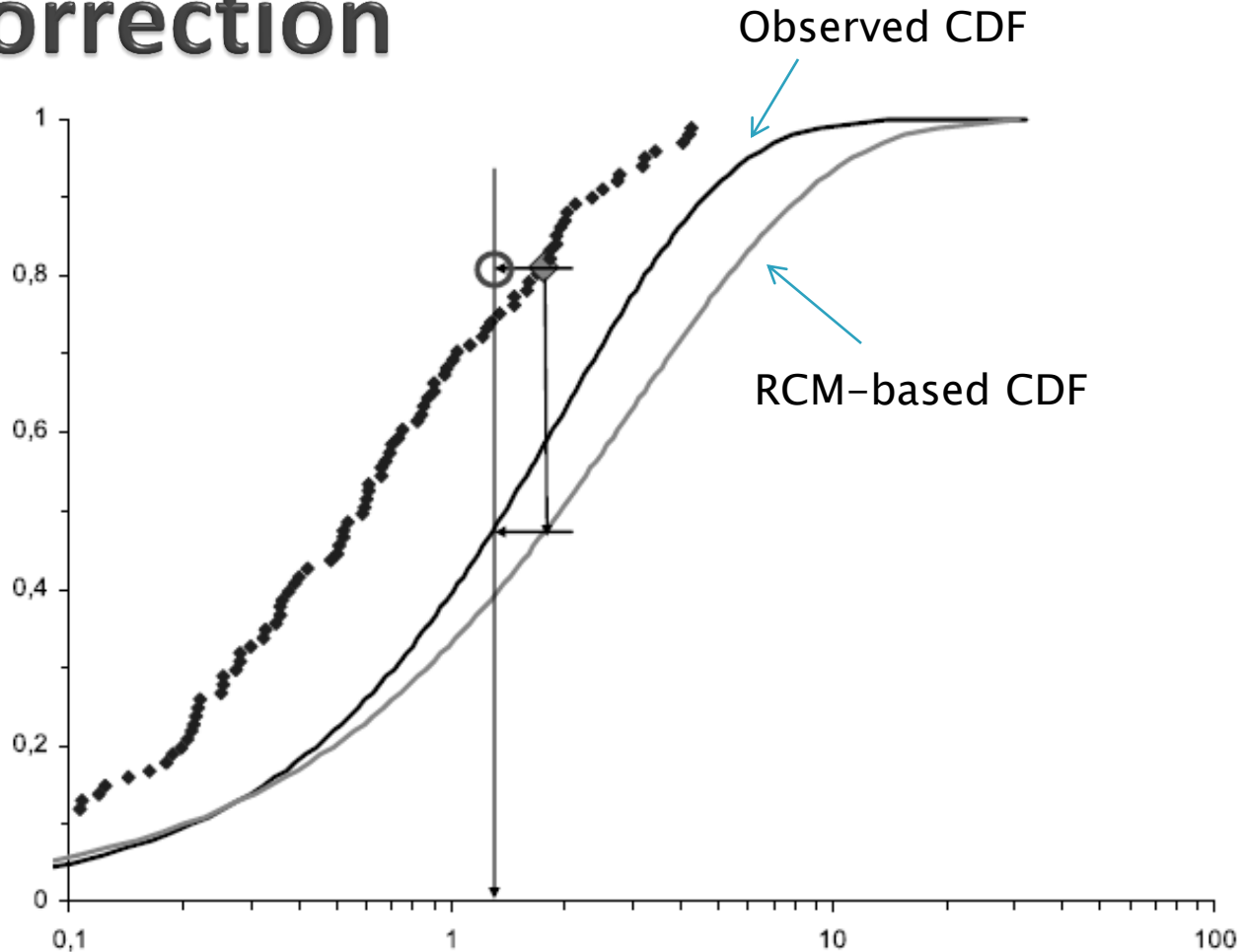


Figure 11. A sketch of the double quantile-quantile transform relating a given RCM future fdf to the observed cdf through the corresponding RCM cdf; the latter would be obtained from the observation period (see equation (2)). The vertical axes show cumulative probabilities, and horizontal axes are in millimeters. The sources are as follows: black line, observed cdf; gray line, RCM-based cdf (observation period); black diamonds, RCM future fdf. The RCM future point (large gray diamond) is matched with the same value on the observed RCM cdf, shifted at the same quantile level to the observed cdf and back to the same quantile level at the future value to give the point in the circle, in the process preserving the rank of the shifted value.

Water Balance Simulation

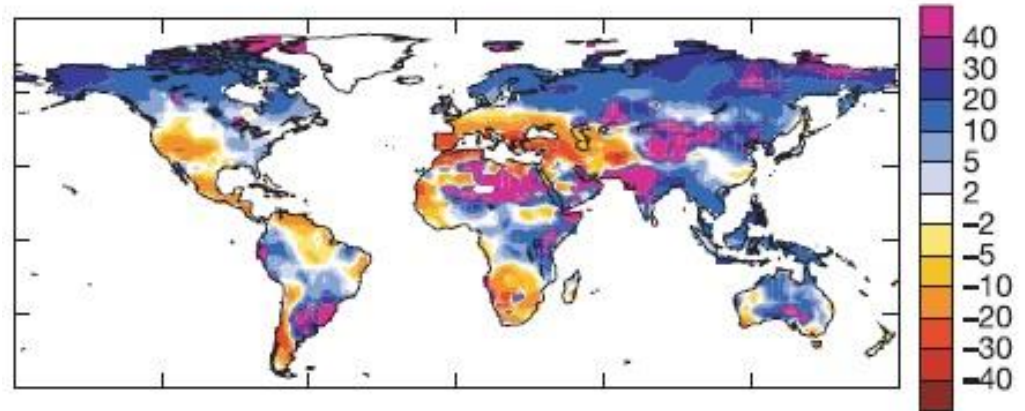


How to estimate the impact?

Use of Global Circulation Models (GCM)

~100-300 km resolution

Budget $P - E$



Use of Hydrological Models

Input precipitation and air temperature from GCM

Better spatial resolution

Water balance simulation

Objective

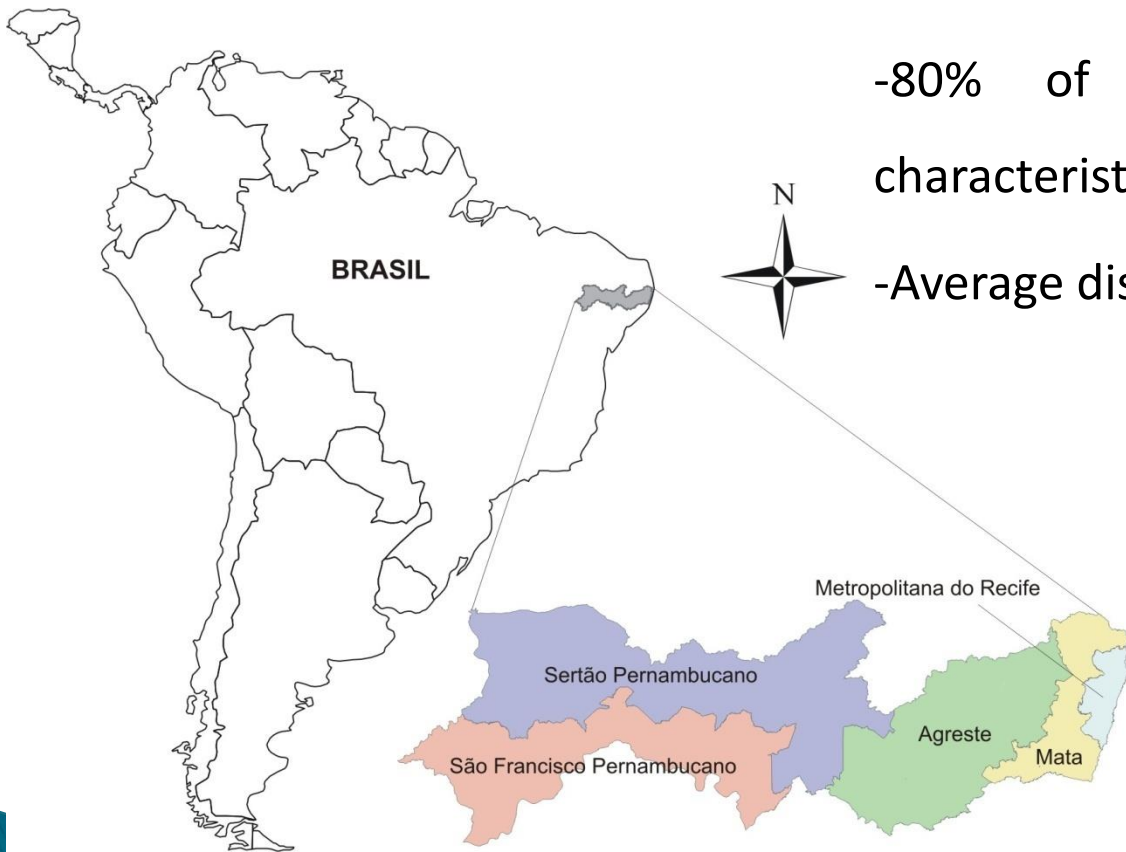
To estimate the impact of climate on surface runoff, evapotranspiration and soil moisture in the territory of Pernambuco State-Brazil.

Area Under Study

-99,123.00 km² in area;

-80% of the area with semiarid characteristics (precipitation 400-800 mm)

-Average discharge: 263 m³/s



Thornthwaite-Mather Method

Determine the water regime of a site using

- Water capacity of the soil
- Precipitation
- Potential Evapotranspiration

Water Capacity of the Soil - W_c

Dunne and Willmott (1996)

Global distribution of plant-extractable water capacity
of soil

0.5° x 0.5° spatial resolution



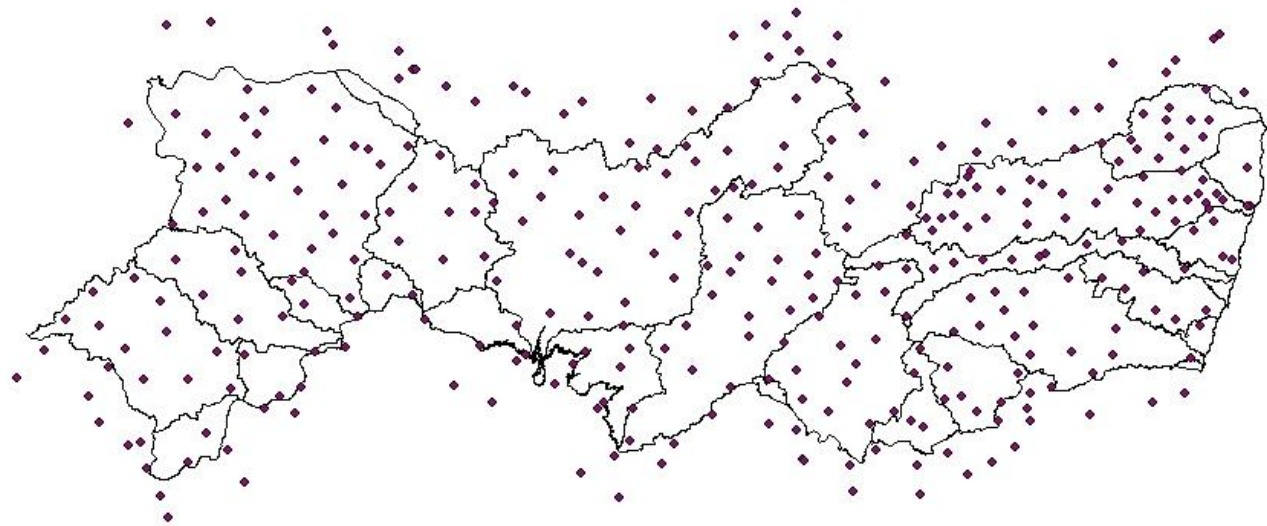
Precipitation - P

National Water Agency

348 rain-gauges

Monthly time step

Inverse distance weighted



Potential evapotranspiration - PET

Hargreaves method

Air temperature (°C) and relative humidity (%)

$$PET = F \cdot 0,158 \cdot (100 - UR)^{0,5} \cdot (32 + 1,8 \cdot T)$$

PET is potential evapotranspiration (mm/month)

F is the potential evapotranspiration factor (mm/month)

Thornthwaite-Mather Method

Surface runoff, water deficit, actual evapotranspiration and soil moisture at each time step.

$$W_{t+1} = W_t \cdot e^{\frac{P_{t+1} - ETP_{t+1}}{W_c}}, \text{ se } P - ETP < 0$$

The actual evapotranspiration (ET) and water deficit (DEF) are given by:

$$ET_{t+1} = P_{t+1} - (W_{t+1} - W_t)$$

$$DEF_{t+1} = ETP_{t+1} - ETR_{t+1}, \text{ if } ETR_{t+1} < ETP_{t+1}$$

$$DEF_{t+1} = 0, \text{ if } ETR_{t+1} = ETP_{t+1}$$

Thornthwaite-Mather Method

When $P - ETP \geq 0$, W , ET and DEF are:

$$W_{t+1} = W_t + P_{t+1} - ETP_{t+1}$$

$$ETR_{t+1} = ETP_{t+1}$$

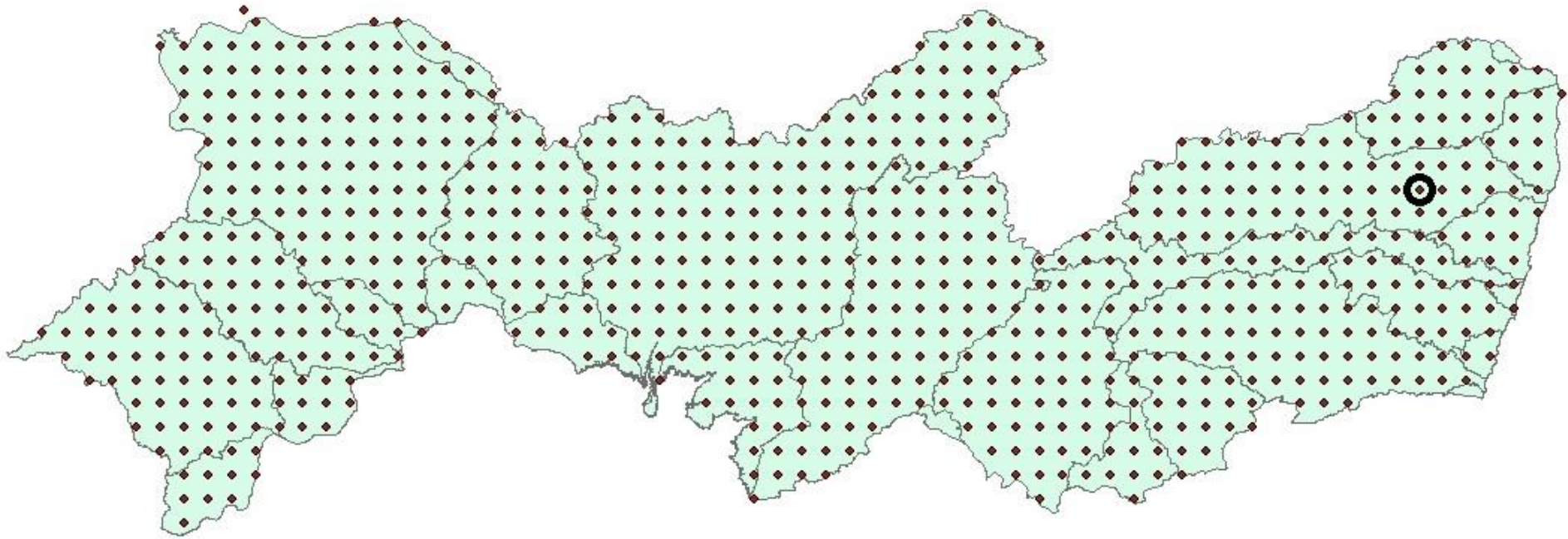
$$DEF_{t+1} = 0$$

The surface runoff (RO) is calculated by

$$RO_{t+1} = W_{t+1} - W_C, \text{ when } W_{t+1} > W_C$$

$$RO_{t+1} = 0, \text{ when } W_{t+1} \leq W_C$$

Thornthwaite-Mather Method



$$W_c = 98 \text{ mm}$$

Regional climate model

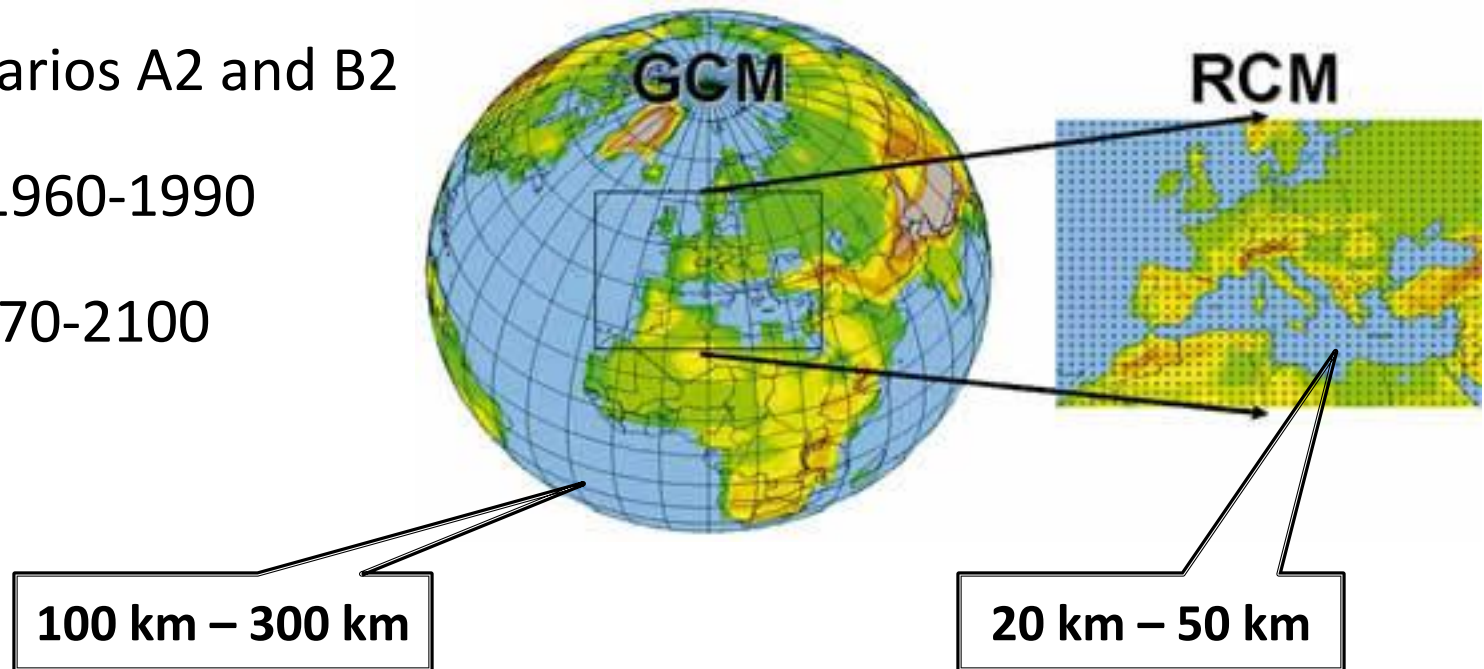
GCM → HadAM3P (UK Met Office Hadley Centre)

RCM → Eta CCS (50 km resolution)

IPCC scenarios A2 and B2

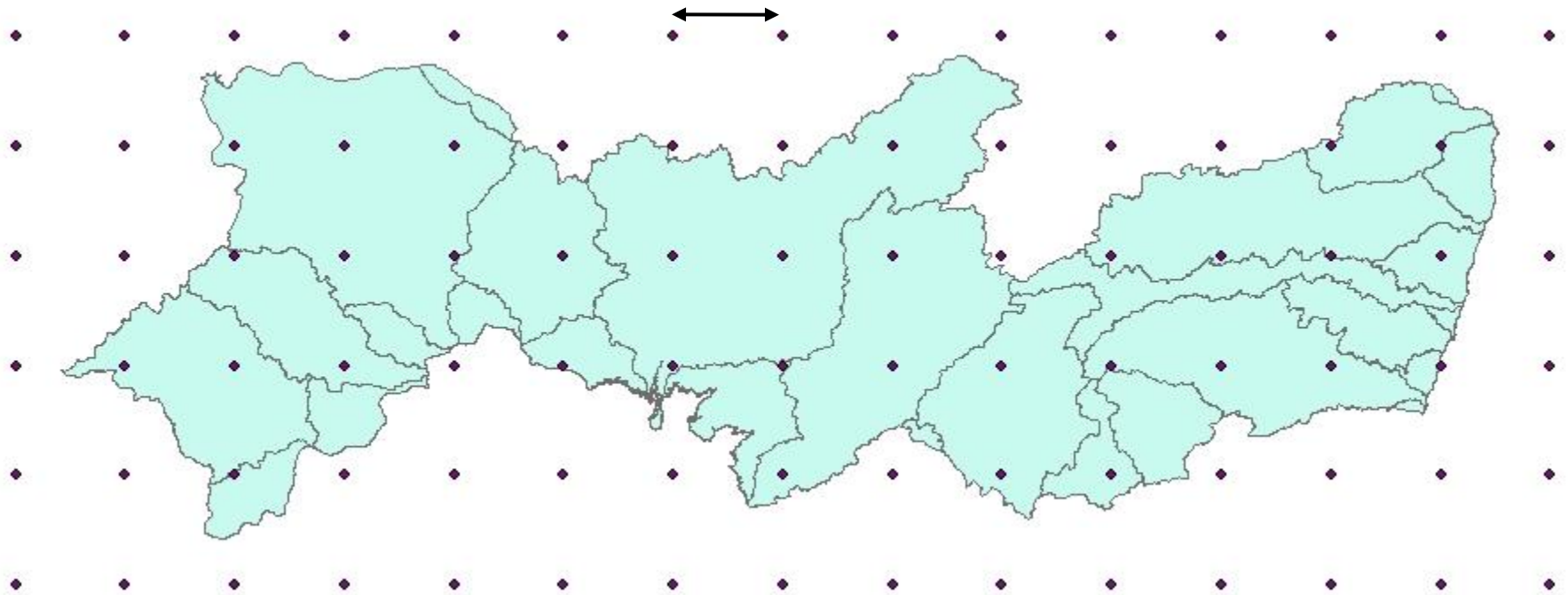
Baseline 1960-1990

Future 2070-2100



Regional climate model

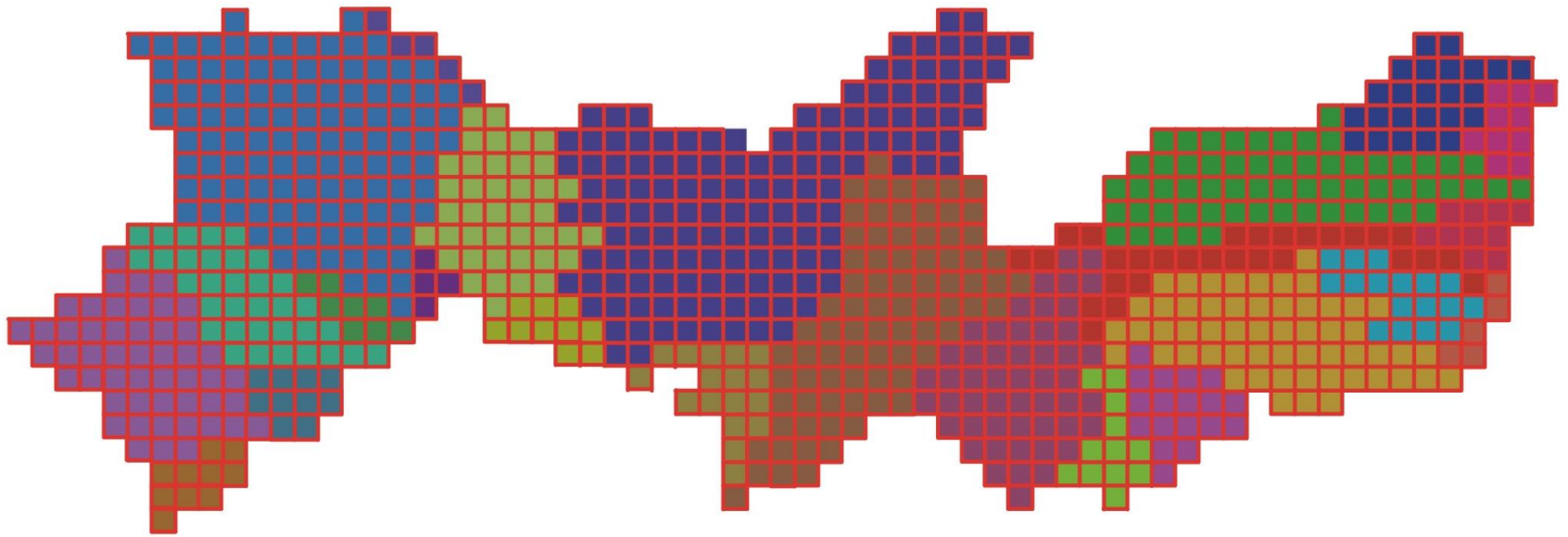
Spatial resolution: 50 km



Water balance

Spatial resolution: $0.1^\circ \times 0.1^\circ$ (about 10 km)

Application of Thornthwaite-Mather method in each cell

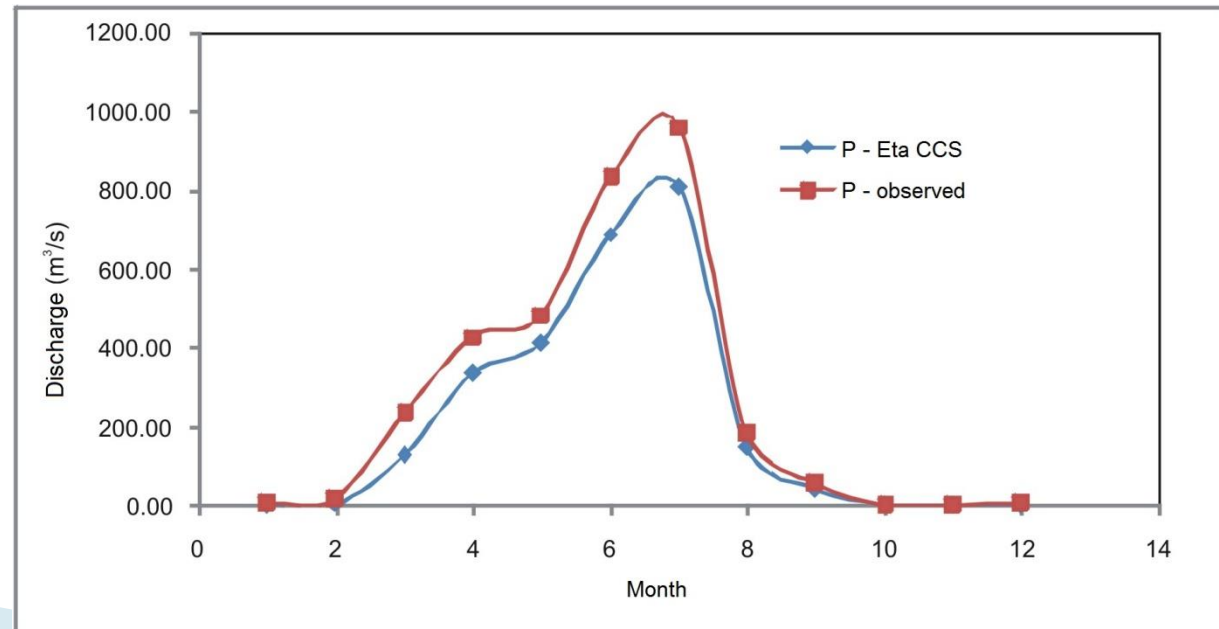


Water balance

Discharge simulated by water resources master plan of Pernambuco: **263.54 m³/s**

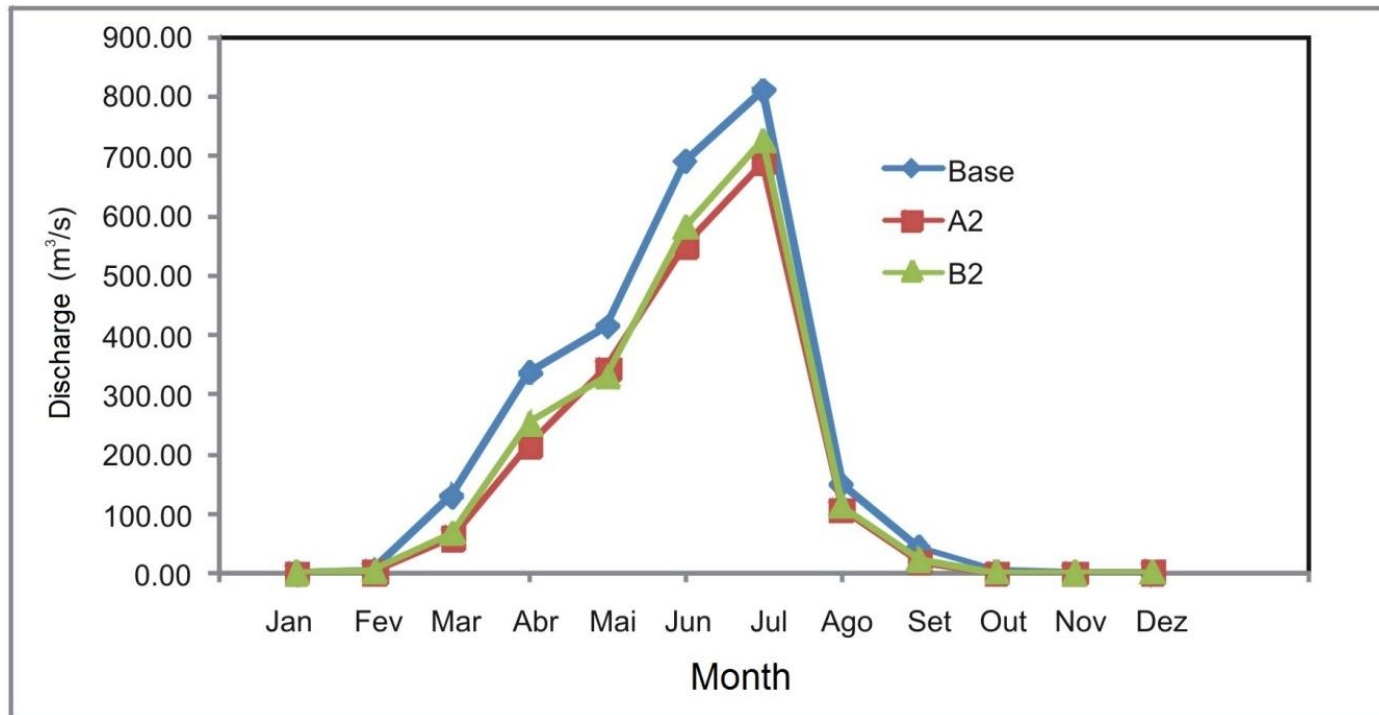
Discharge simulated with observed precipitation: **267.78 m³/s**

Discharge simulated with Eta CCS precipitation: **213.86 m³/s**



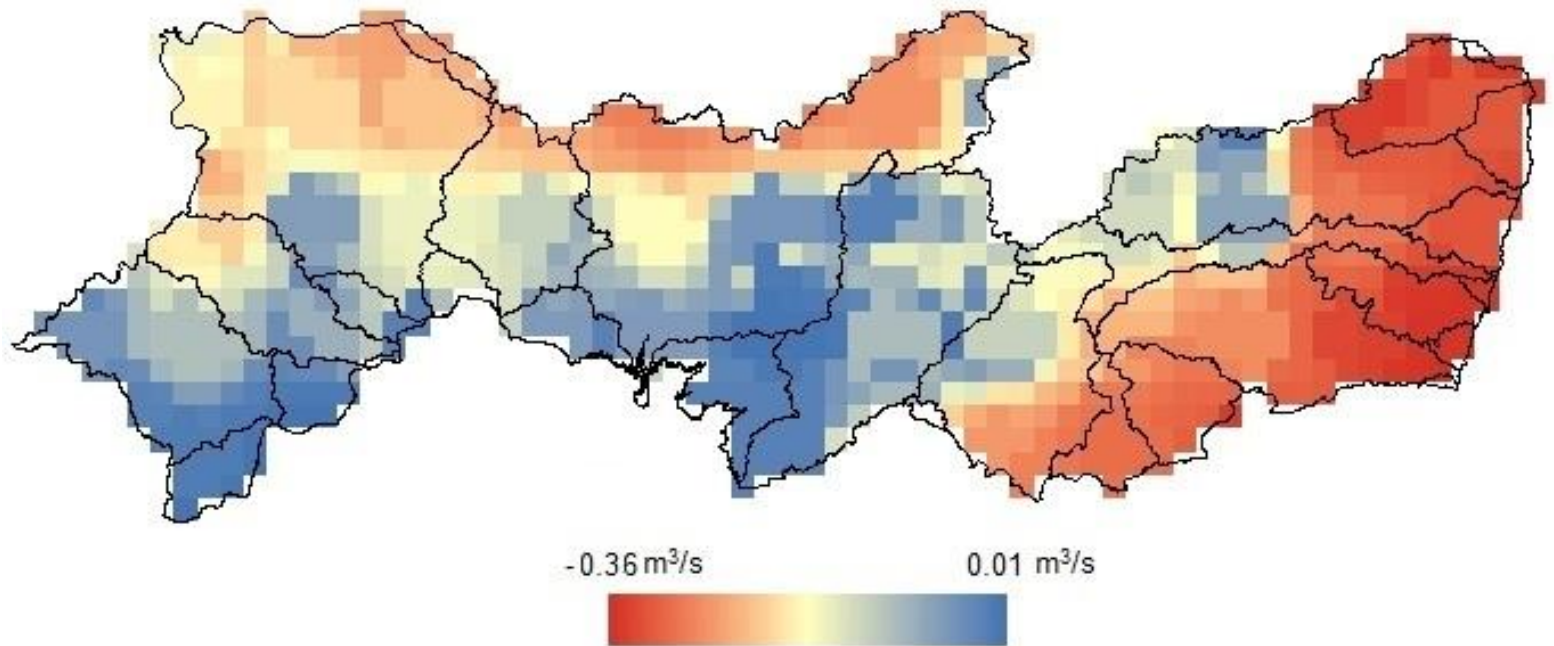
Results

Discharge in the baseline (1960-1990) and scenarios A2 and B2 (2070-2100)



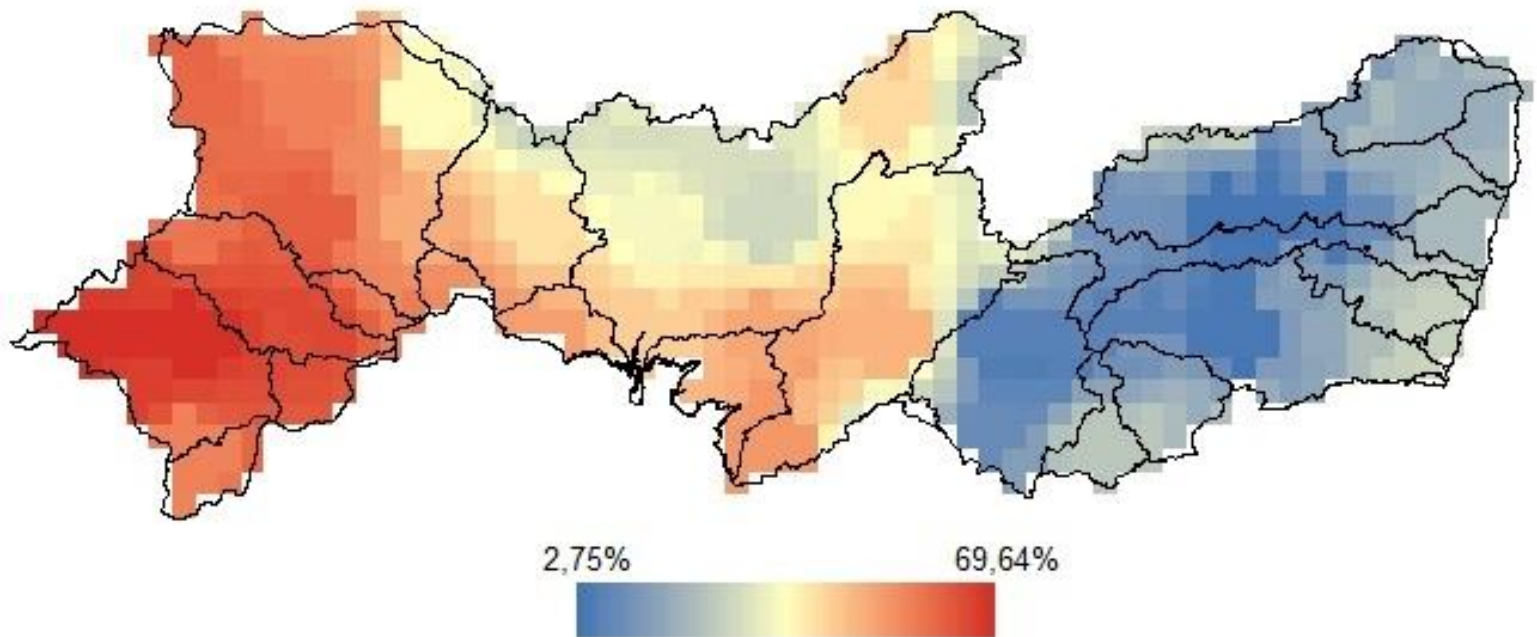
Results

Difference in the **surface runoff** between scenario A2 and baseline in m^3/s (negative values representing reduction of the runoff).



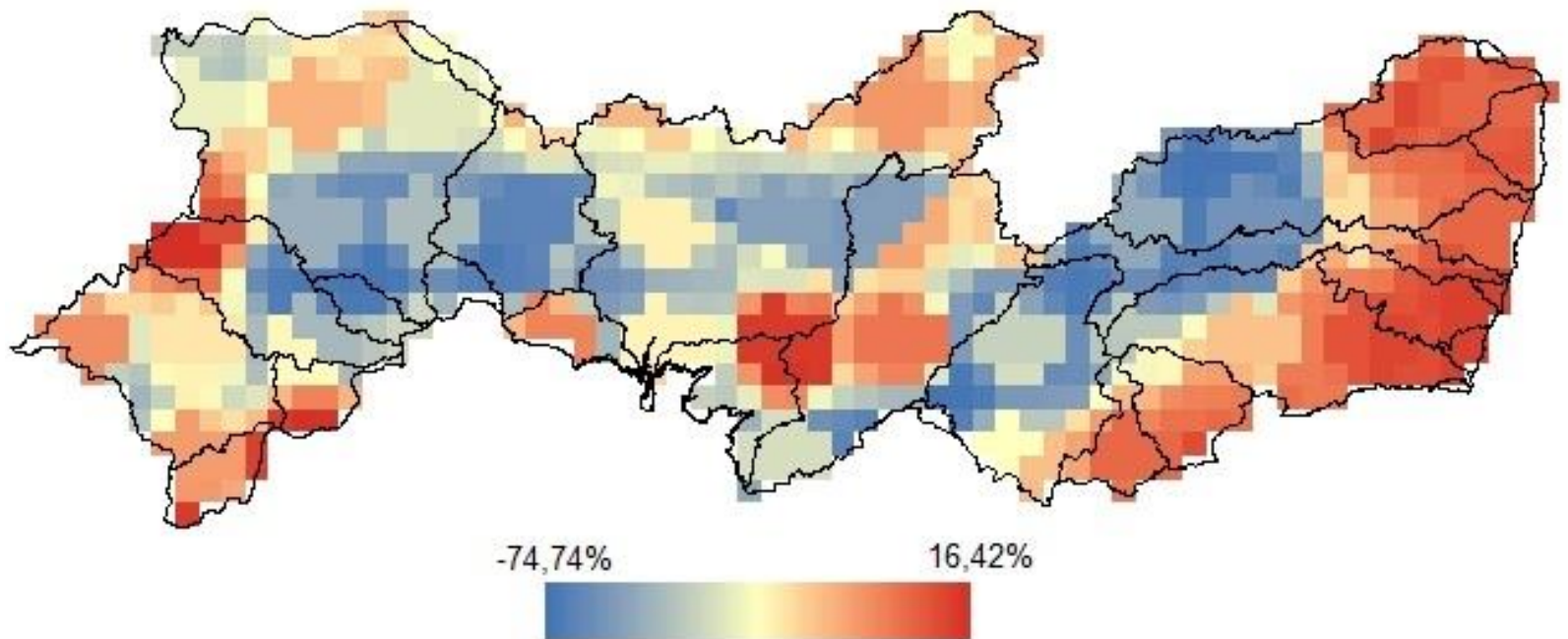
Results

Difference in the actual **evapotranspiration** between scenario A2 and baseline expressed as a percentage.

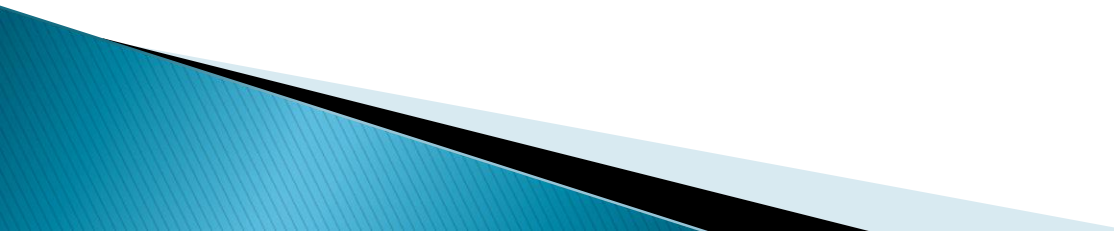


Results

Difference in the **soil moisture** between scenario A2 and baseline expressed as a percentage.



Results

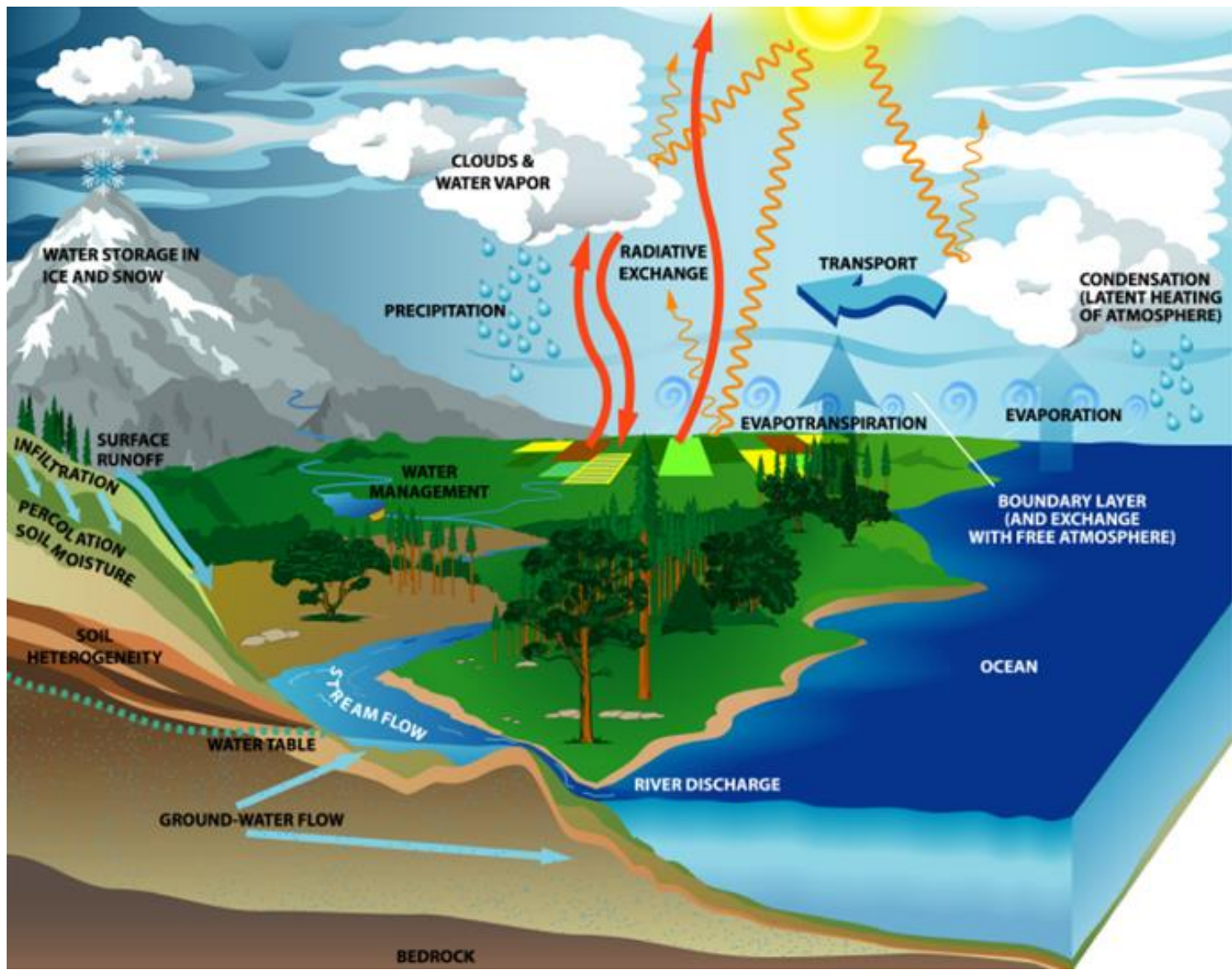
- Results in agreement with Milly (2005): surface runoff reduction of 20%
 - And disagreement with UK Met Office (2005) and Salati et al. (2008)
 - Thornthwaite-Mather is an alternative to complex hydrological models
- 

Hydrological Model

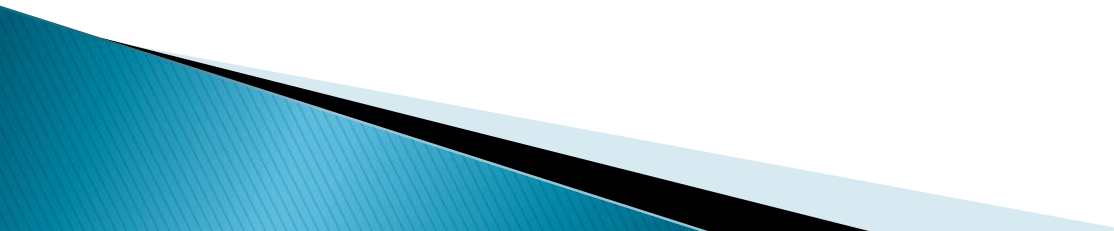


Concept

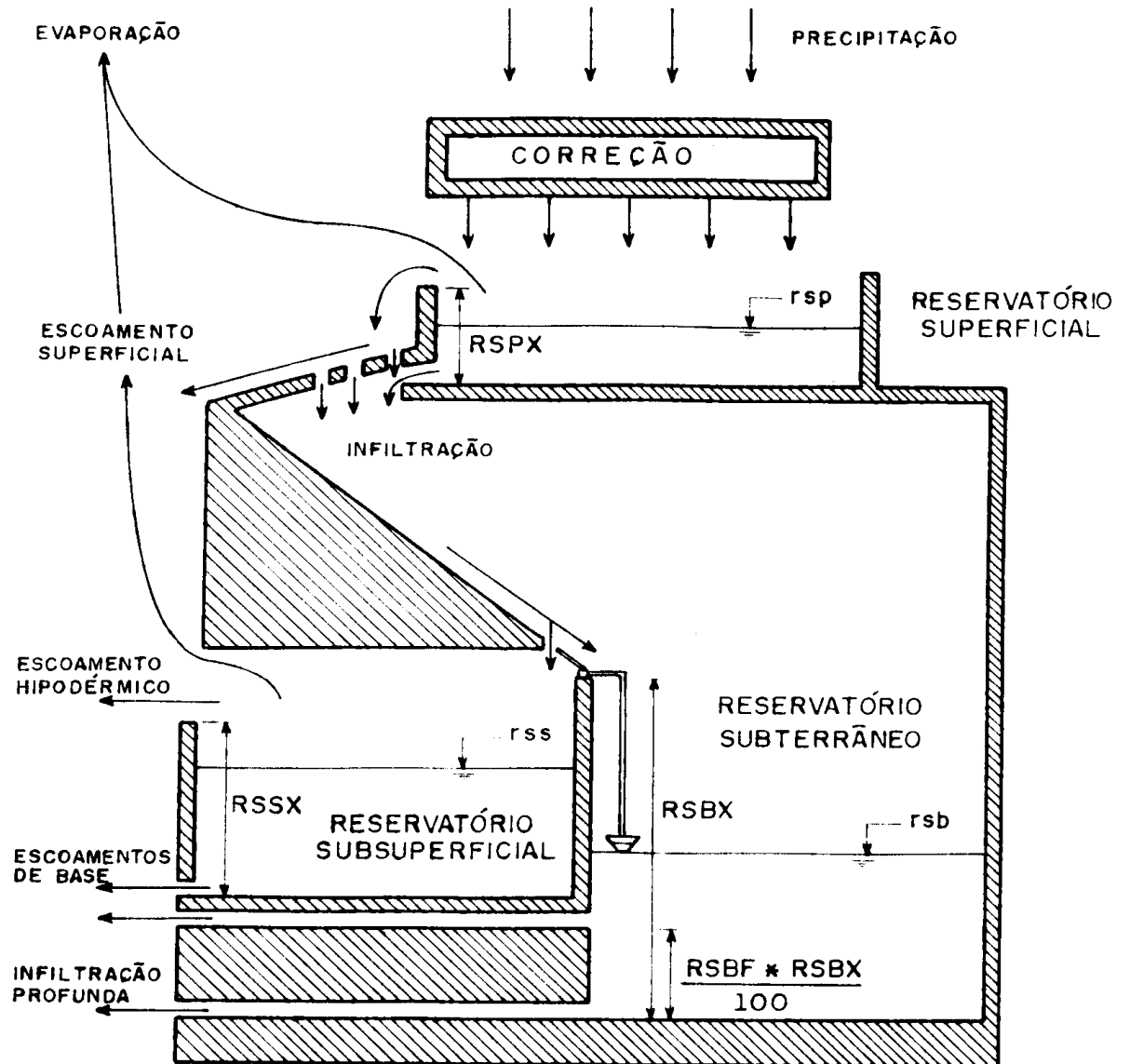
Rainfall–runoff models represent the part of hydrological cycle between the rainfall and the streamflow. They simulate the spatial distribution of rainfall, losses by interception, evaporation, depression in the soil, flow into the soil by the infiltration, percolation and groundwater, surface flow, interflow and the flow in the river.



MODHAC Model

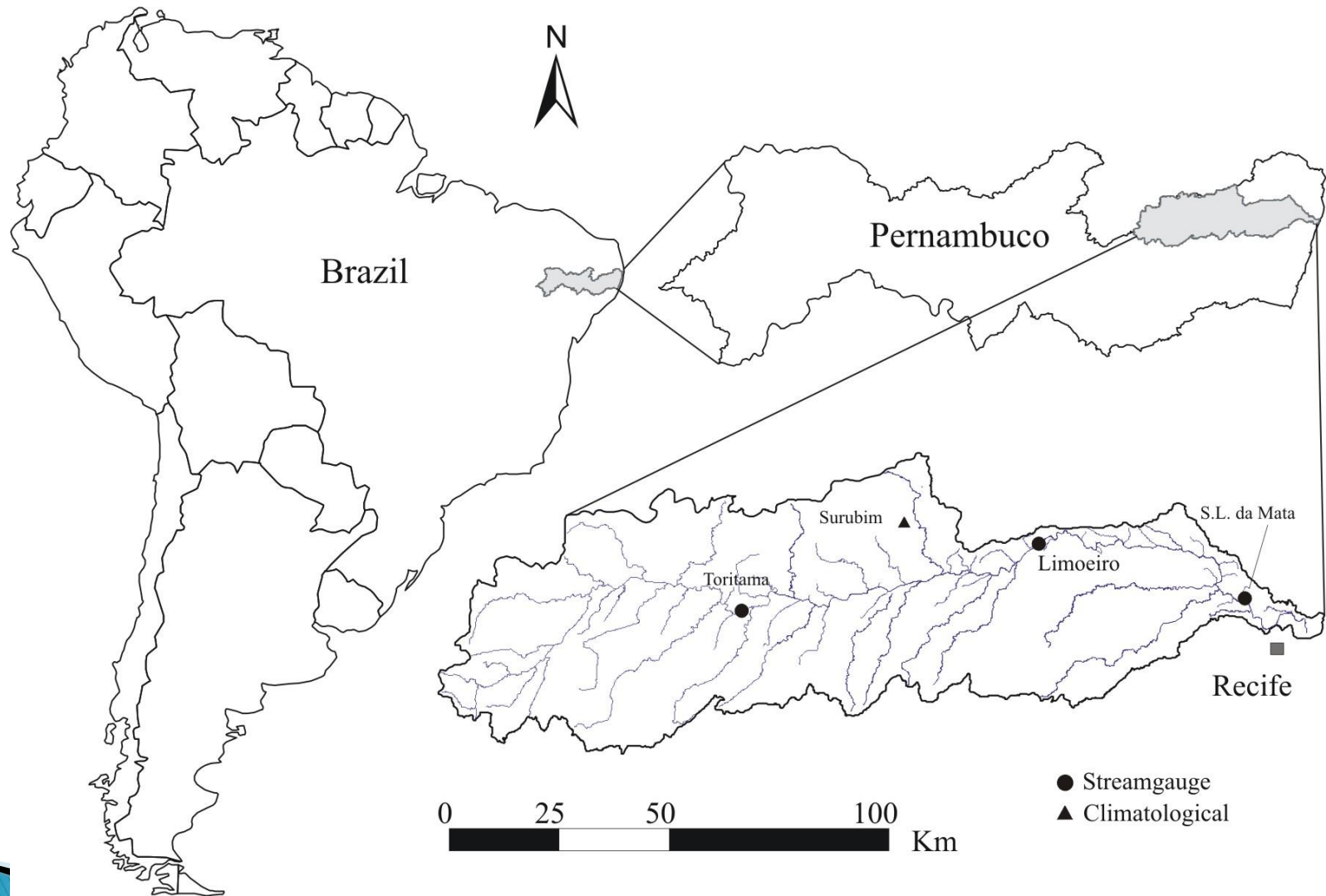
- ▶ Deterministic, conceptual, lumped
 - ▶ Daily and monthly time step
 - ▶ Input data
 - Precipitation
 - Observed streamflow
 - Potential evapotranspiration
 - ▶ Output: streamflow at the mouth of the basin
- 

MODHAC Model

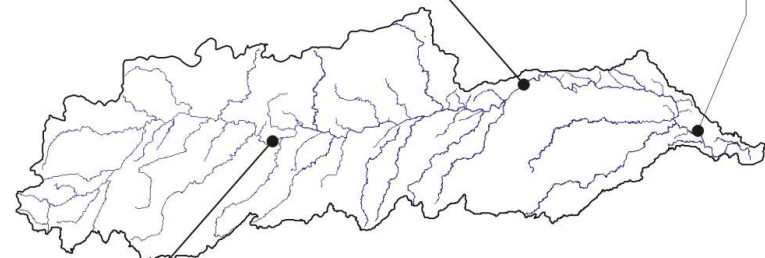
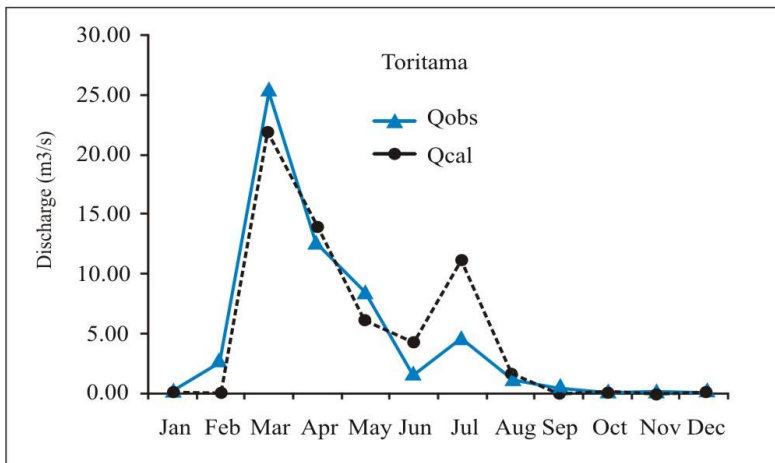
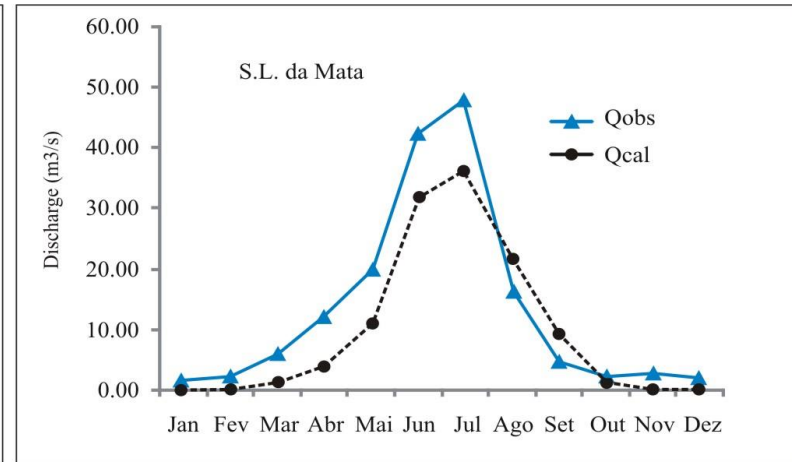
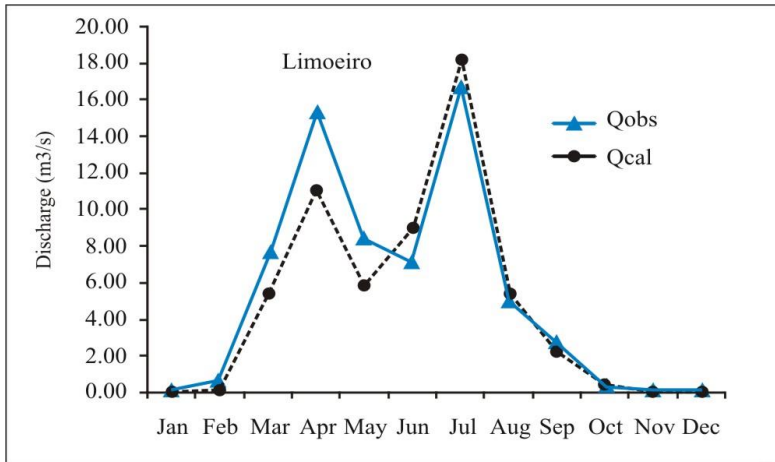


| Parameter | Description | Value | Mínimum | Máximum |
|-----------|--|--------|---------|---------|
| RSPX (mm) | capacidade máxima do reservatório superficial | 43,47 | 0 | 60 |
| RSSX (mm) | capacidade máxima do reservatório sub-superficial | 199,7 | 20 | 300 |
| RSBX (mm) | capacidade máxima do reservatório subterrâneo | 6,15 | 0 | 300 |
| RSBY | Efetivos no ajuste da curva de recessão do hidrograma | 0 | 0 | 100 |
| IMAX (mm) | permeabilidade do solo (infiltração máxima) | 121,4 | 20 | 100 |
| IMIN (mm) | infiltração mínima | 0,1278 | 0 | 10 |
| IDEC | coeficiente de infiltração | 0,1387 | 0 | 1 |
| ASP | Expoente da lei de esvaziamento do reservatório superficial | 0,4399 | 0 | 1 |
| ASS | Expoente da lei de esvaziamento do reservatório sub-superficial | 0,0262 | 0 | 1 |
| ASBX | Expoente da lei de esvaziamento do reservatório subterrâneo | 1 | 0,001 | 0,1 |
| ASBY | Efetivos no ajuste da curva de recessão do hidrograma | 1 | 0 | 1 |
| PRED | correção da precipitação | 999,9 | 0 | 0 |
| CEVA | parâmetro da lei de evapotranspiração do solo | 0,7 | 0 | 1 |
| CHET | fração da evapotranspiração potencial (evaporação direta da chuva) | 0,85 | 0 | 1 |

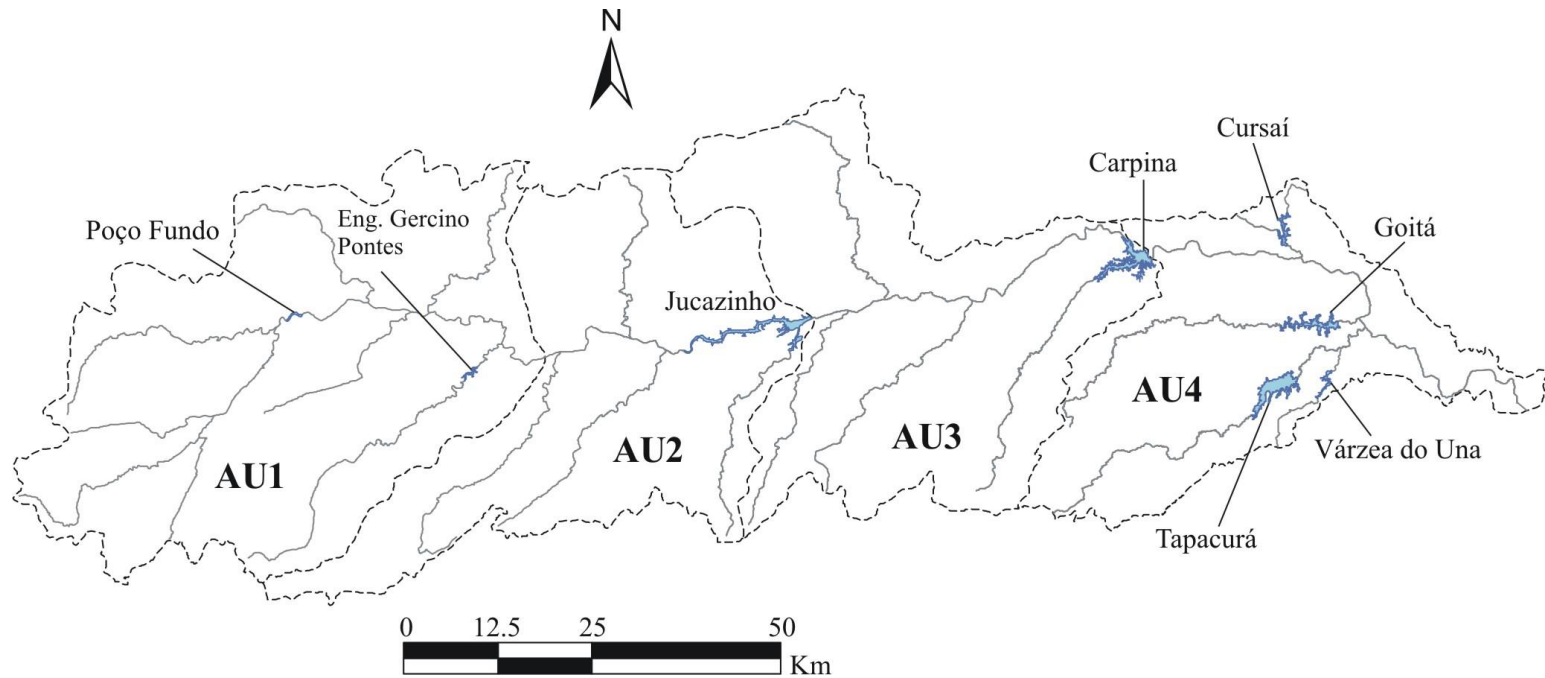
Application



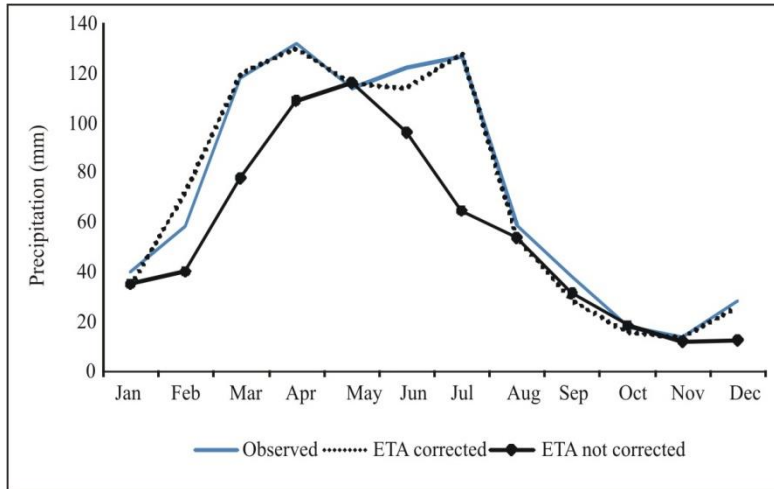
Calibration



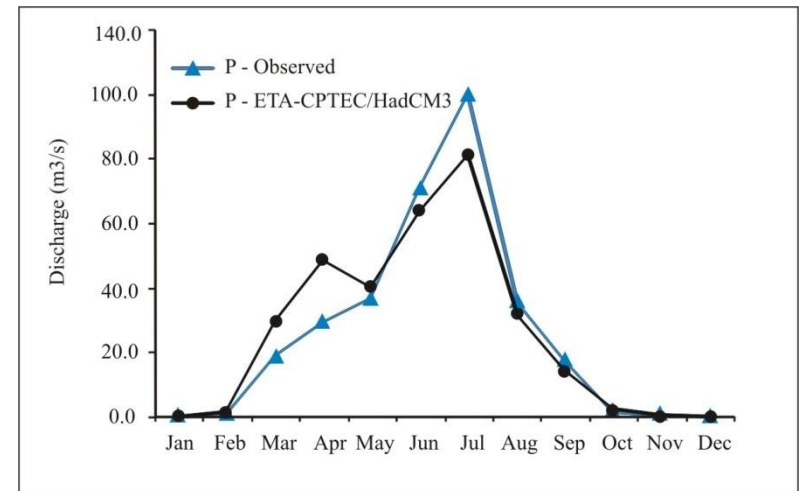
Simulation



Simulation

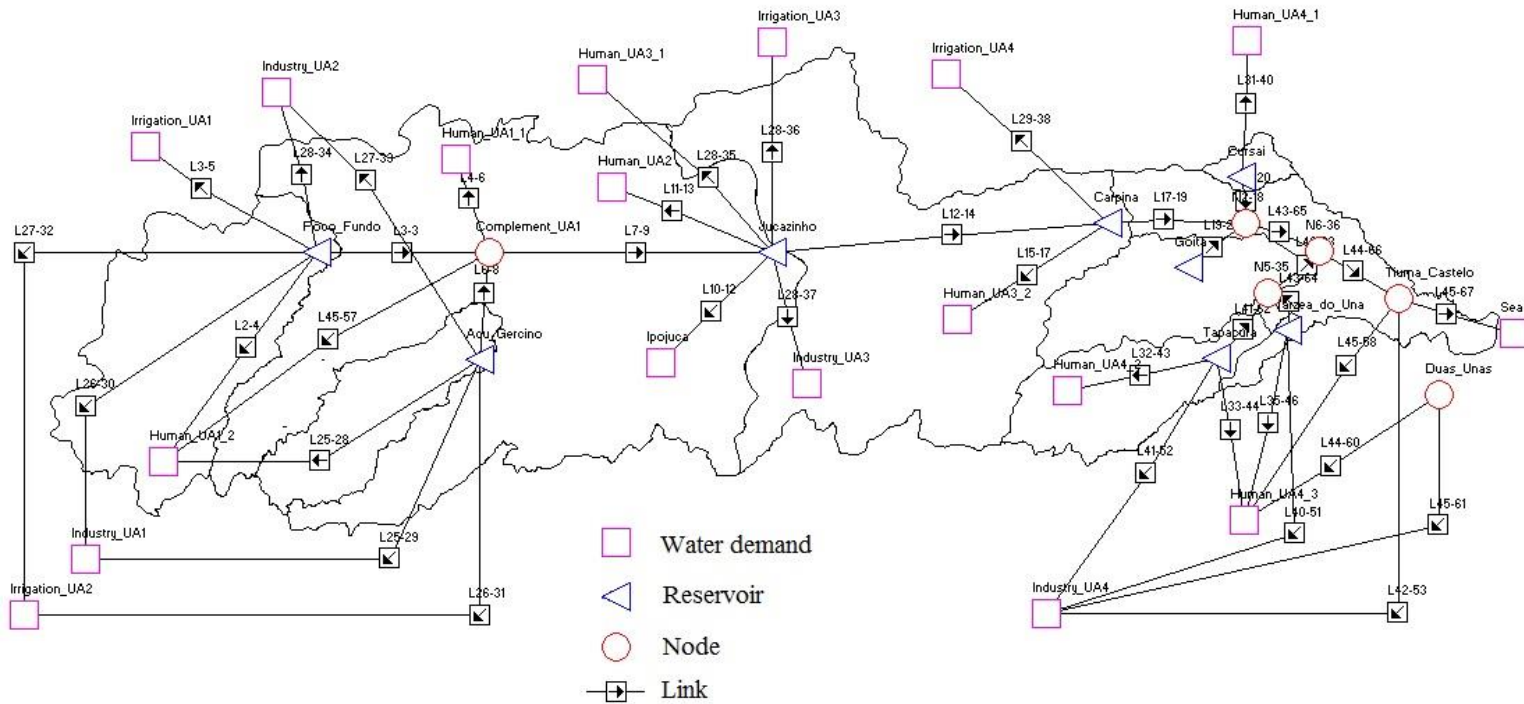


Mean precipitation in the Capibaribe river basin

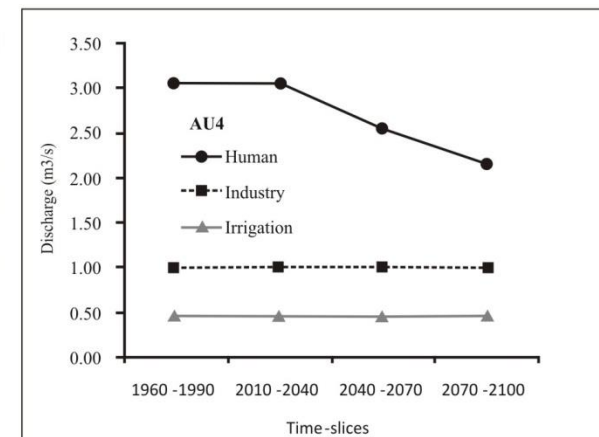
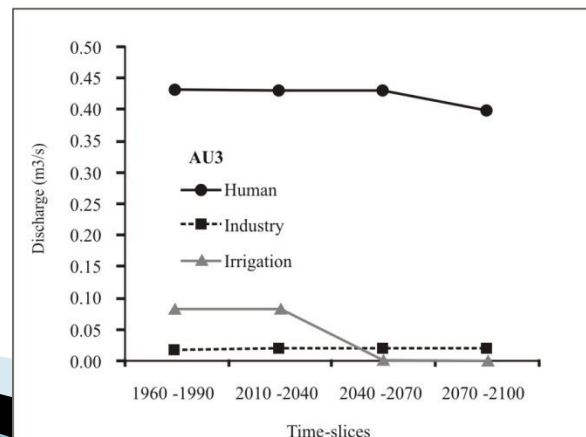
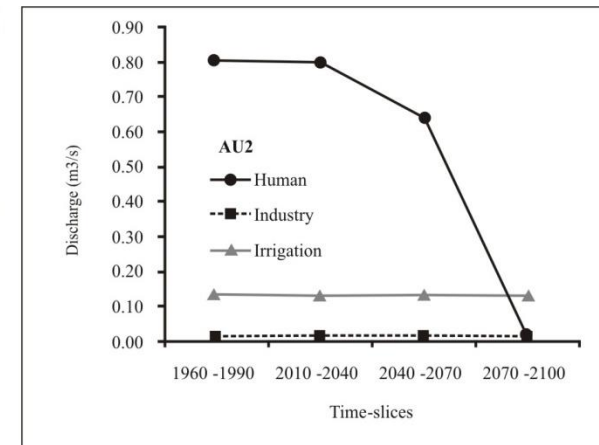
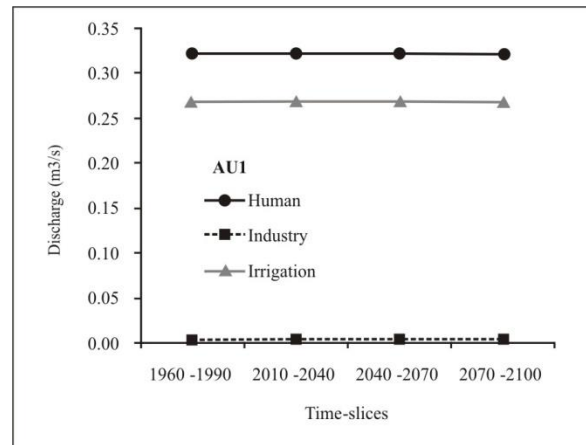
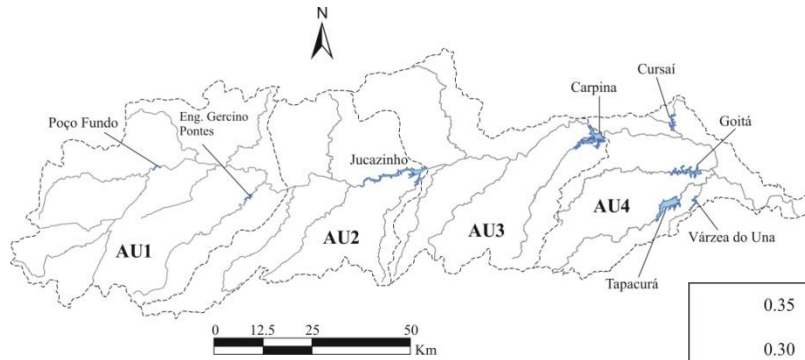


Streamflow at the outlet section of Capibaribe river calculated with MODHAC

Network Flow Model



Network Flow Model



Simulation and Optimization



System analysis techniques in water resources

Simulate real-world and optimize the decision processes that play a role on this reality.

Simulation

Modelling techniques used to represent the behavior of a system.

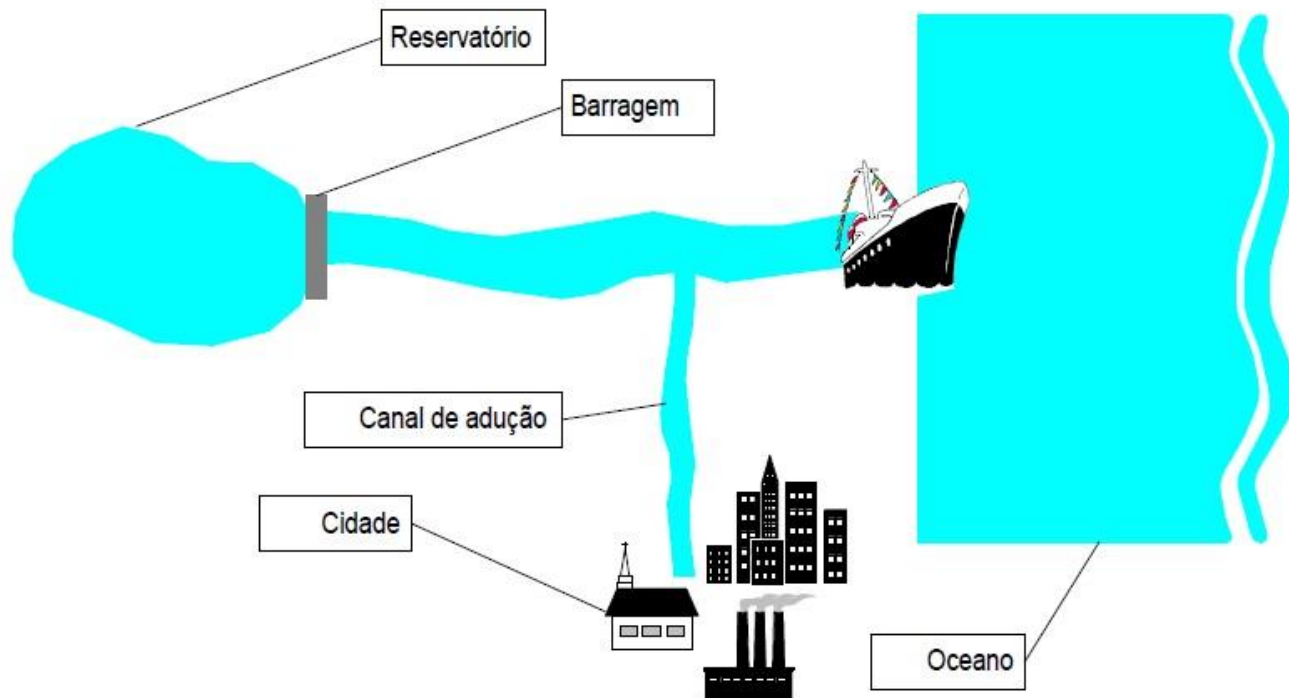
Optimization

Decision process according to a valuation established by the Objective-Function.



Simulation

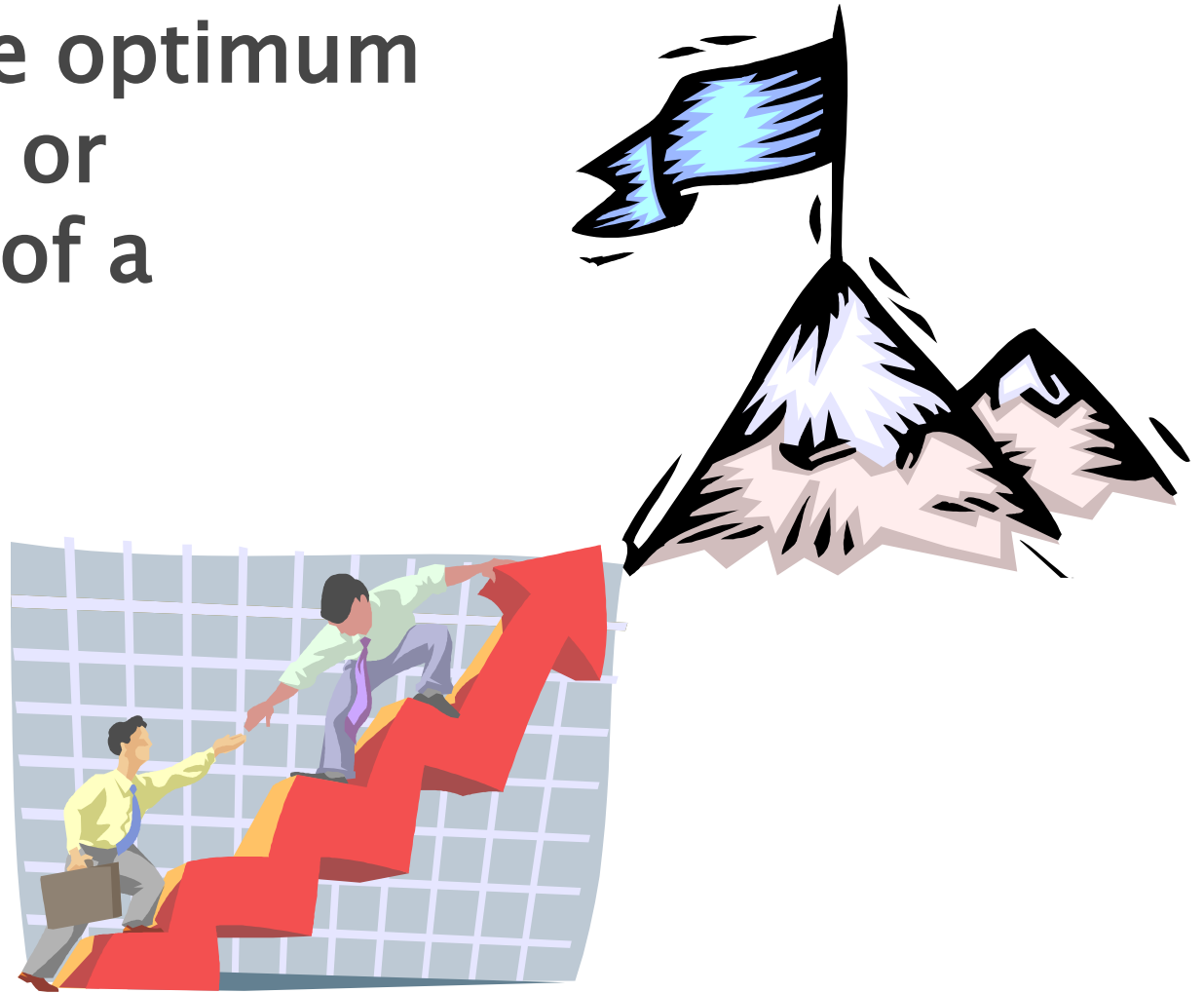
Reservoir yield that supplies water for one city.



Representação esquemática de um sistema de recursos hídricos

OPTIMIZATION:

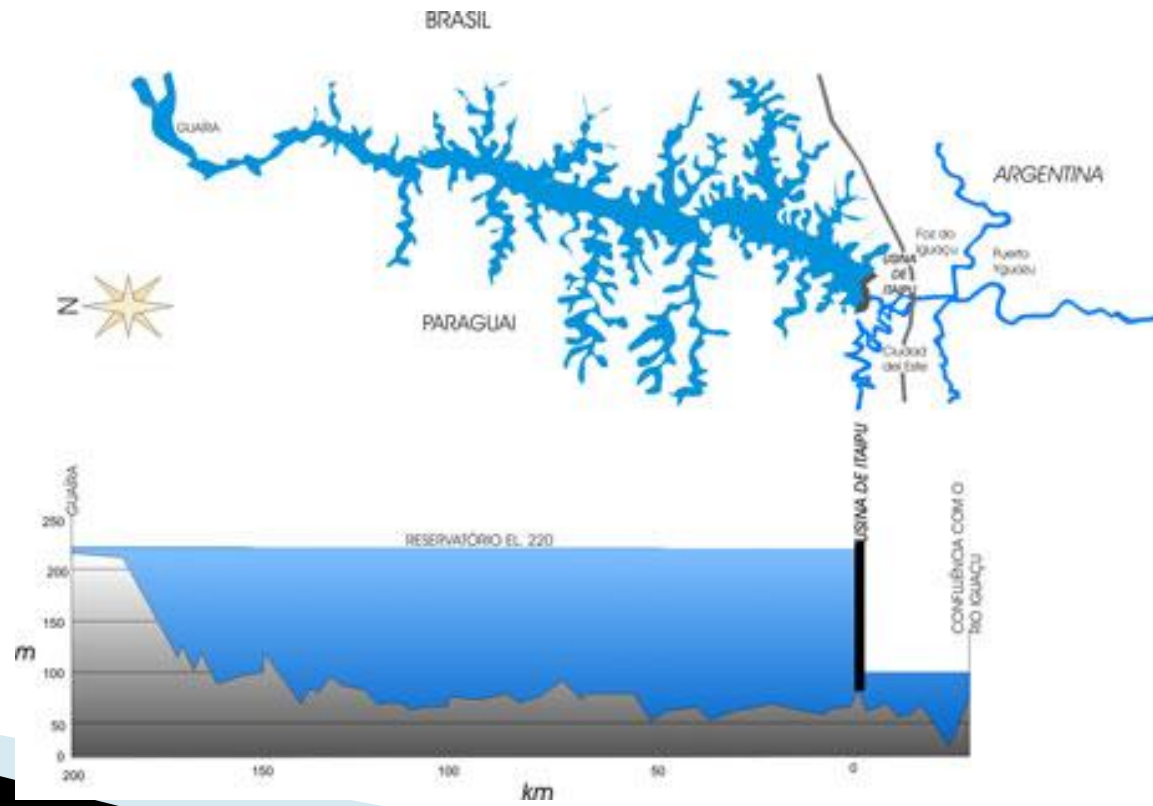
Achieve the optimum
(maximum or
minimum) of a
process.



Optimization

Use of mathematical techniques to:

-Design reservoir capacity



Optimization

Use of mathematical techniques to:

-Design canals



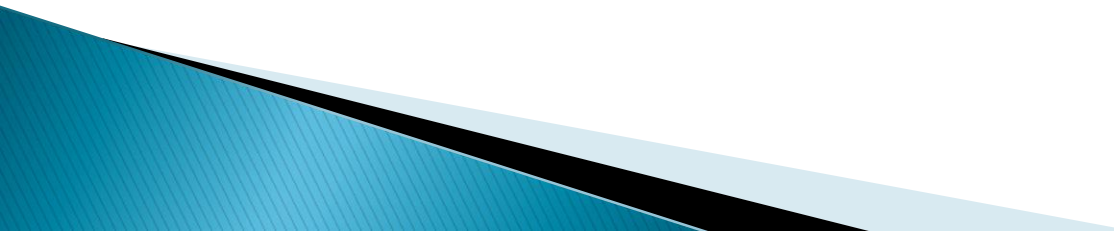
Optimization

Use of mathematical techniques to:

–Water allocation for several uses

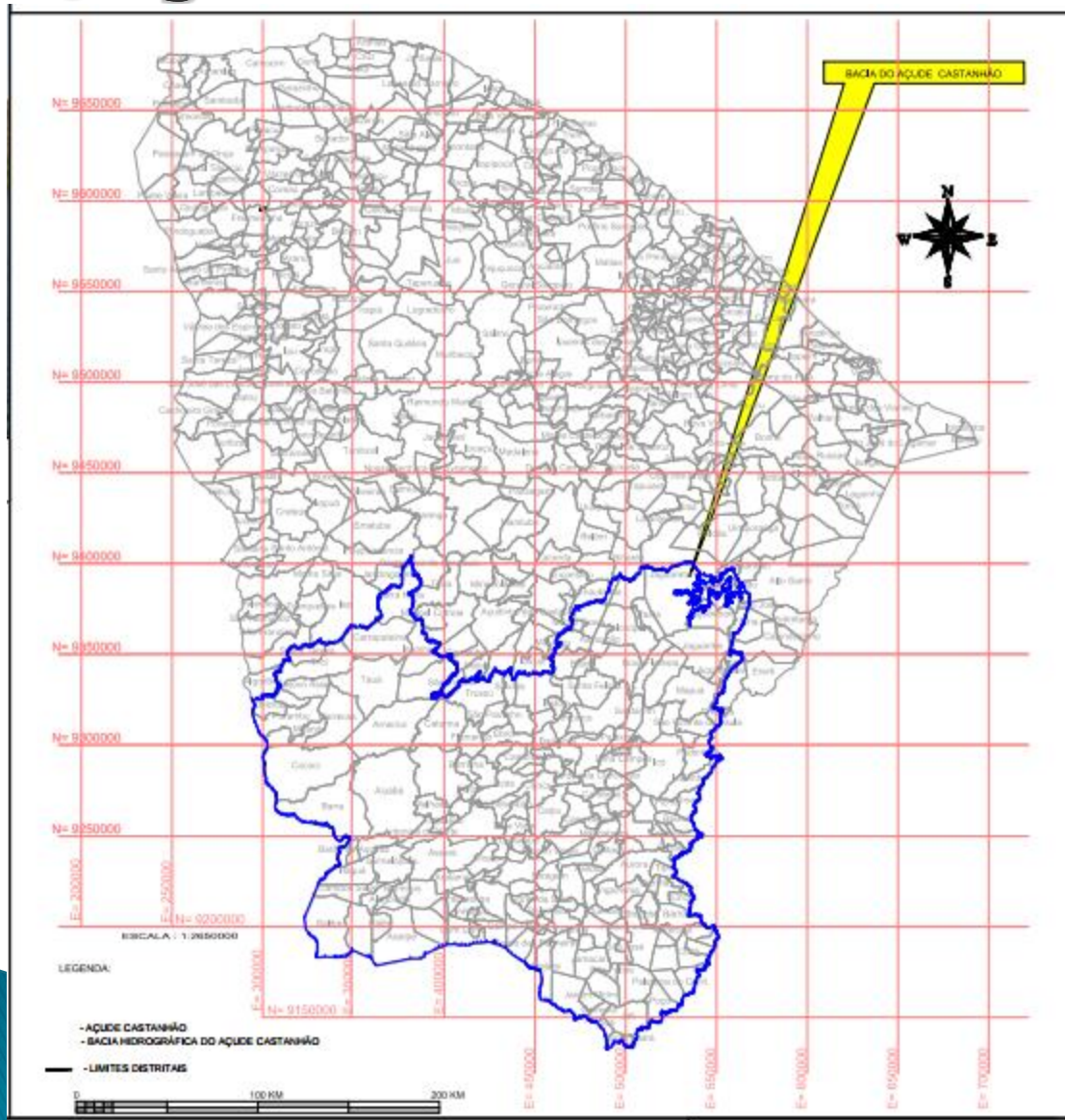


Optimization

- ▶ Decision Variables
 - ▶ Objective Functions
 - ▶ Constraints
 - ▶ State Variables
- 

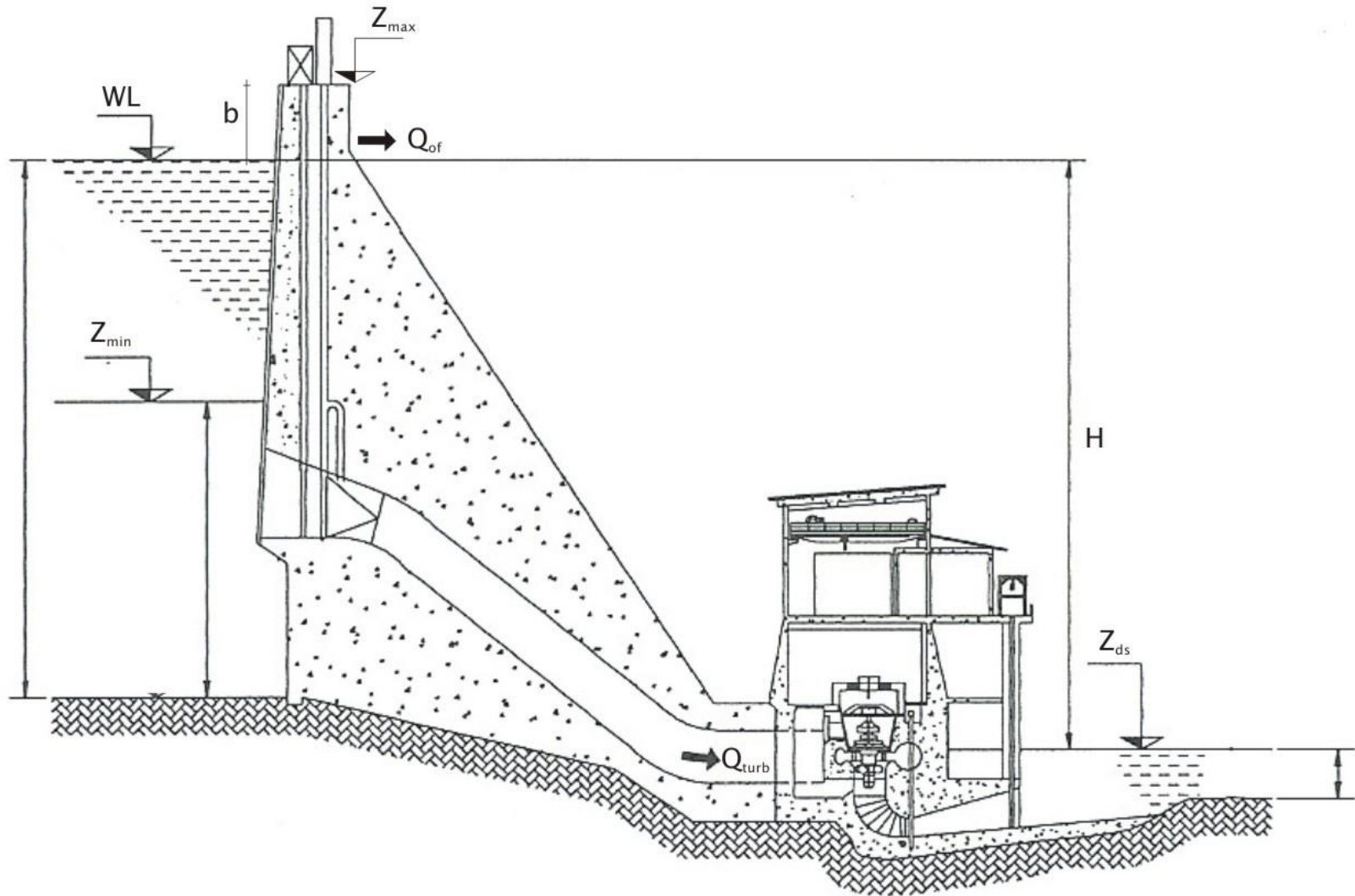
Exercise

Jaguaribe River–Castanhão reservoir



Ceará State in Northeast
Drainage area: 45,450 km²
Capacity: 6.7 billion m³

Multiple Use Reservoir



Multiple Use Reservoir

Decision Variables

- ▶ Maximum level of the dam: Z_{\max}
- ▶ Discharge for power generation: Q_{turb}
- ▶ Discharge for irrigation: Q_{irrig}

Multiple Use Reservoir

State Variables

- ▶ Operational water level: $Z = Z_{\max} - b$
- ▶ Storage: $S = a_s \cdot Z^3 + b_s \cdot Z^2 + c_s \cdot Z + d_s$
- ▶ Yield discharge: $Q_{yie} = a_y \cdot (Z - Z_0)^3 + b_y \cdot (Z - Z_0)^2 + c_y \cdot (Z - Z_0) + d_y$
- ▶ Overflow discharge: $Q_{of} = a_o \cdot e^{b_o \cdot (Z - Z_0)}$
- ▶ Total discharge: $Q_{tot} = Q_{turb} + Q_{irrig}$
- ▶ Maximum inundated area: $A = a_a \cdot Z^2 + b_a \cdot Z + c_a$

Multiple Use Reservoir

State Variables

- ▶ Volume for irrigation per year: $V_{\text{irrig}} = Q_{\text{irrig}} \cdot 86400 \cdot 365$
- ▶ Water head for power generation: $H = Z - Z_{\text{ds}}$
- ▶ Discharge returned to the river: $Q_{\text{ds}} = Q_{\text{turb}} + Q_{\text{of}}$
- ▶ Discharge supplied in the n hours of irrigation:
 $Q_n = Q_{\text{irrig}} \cdot 24 / n$
- ▶ Area that can be irrigated with Q_n : $A_{\text{irrig}} = Q_{\text{irrig}} \cdot 1000 / q_{\text{irrig}}$

Multiple Use Reservoir

Benefit–Cost

1. Capital costs

- ▶ Expropriated area: $A_e = a_e \cdot Z_{\max}^2 + b_e \cdot Z_{\max} + c_e$
- ▶ Total cost of expropriation: $C_e = CU_e \cdot A_e$
- ▶ Costs of execution: $C_{ex} = a_3 \cdot (Z - Z_1)^{b3}$
- ▶ Total capital cost: $C_c = C_e + C_{ex}$
- ▶ Capital recovery factor: $R = \frac{i \cdot (i + 1)^n}{(1 + i)^n - 1}$
- ▶ Annual amortization cost: $C_a = R \cdot C_c$

Multiple Use Reservoir

Benefit–Cost

2. Irrigation

▶ Operational cost: $C_{\text{irrig}} = A_{\text{irrig}} \cdot C_{\text{u}_{\text{irrig}}}$

▶ Price of selling of the production:

$$B_{\text{irrig}} = k_1 \cdot V_{\text{irrig}} + k_2 \cdot \ln(0.01 + 0.3 \cdot V_{\text{irrig}})$$

▶ Profit of irrigação: $\text{Prof}_{\text{irrig}} = B_{\text{irrig}} - C_{\text{irrig}}$

Multiple Use Reservoir

Benefit–Cost

3. Energy

- ▶ Energy generated: $E(\text{kwh}) = 85935.6 \times Q_{\text{turb}} \times H \times \eta$
- ▶ Benefit of energy: $Be = E \times 0.40$

Multiple Use Reservoir

Benefit–Cost

4. Navigation

- ▶ Benefit of navigation: $B_n = a_4 \cdot Q_{ds}^{b_4}$


Multiple Use Reservoir

Constraints

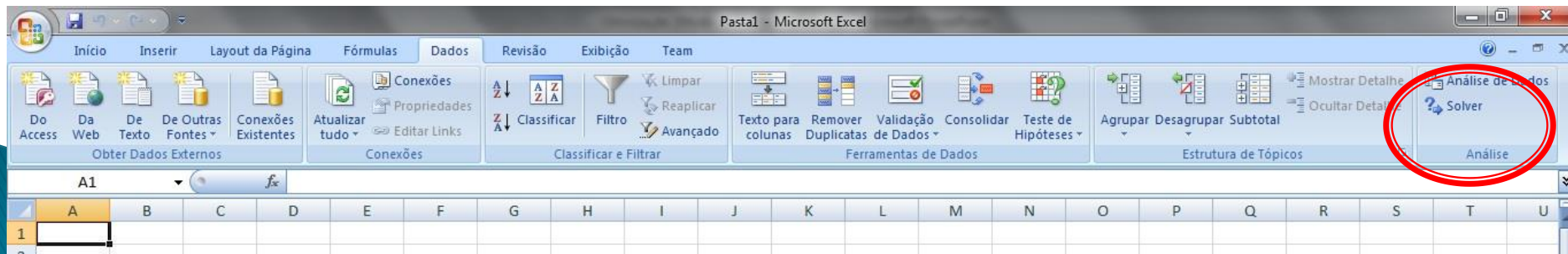
1. $Z_{\max} \geq 0$
2. $Q_{\text{turb}} \geq 0$
3. $Q_{\text{irrig}} \geq 0$
4. $Z_{\max} \leq \text{Maximum level possible}$
5. $Z_{\max} \geq \text{Minimum level possible}$
6. $A_{\text{irrig}} \leq A_{\text{max,irrig}}$
7. $Q_{\text{ds}} \geq Q_{\text{min}}$
8. $Q_{\text{tot}} \leq Q_{\text{yie}}$
9. $H \geq H_{\text{min}}$

Solver

Microsoft Office 2007

Click in “Buton Office”  and, after, “Excel Options”. Select “Supplement” and click in the buton “Go ...”. Select the Solver clicking in the box;

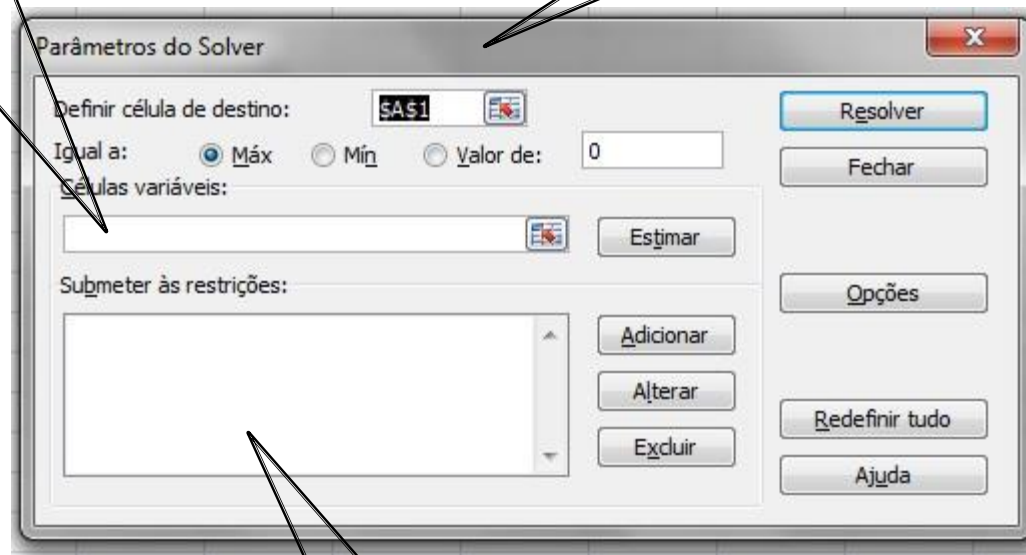
Solver is accessed in menu “Data”.



Solver

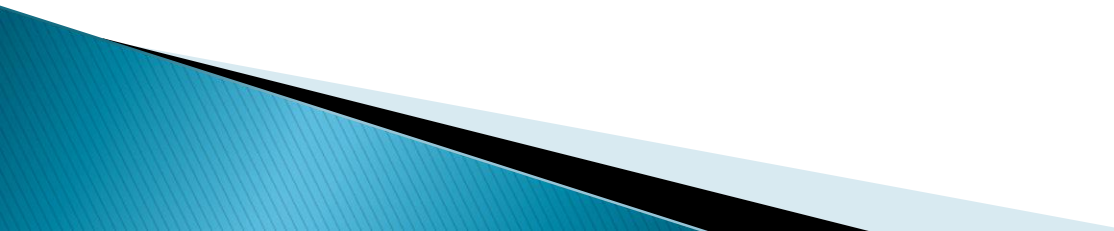
Decision Variables

Objective Function



Constraints

Summary

- ▶ Downscaling: dynamic and statistic
 - ▶ Bias correction
 - ▶ Hydrological model
 - ▶ Simulation/Optimization tools
 - ▶ Simulate IPCC scenarios
 - ▶ Evaluate adaptation actions
- 

References

- Bárdossy, A. and Pegram, G., 2011. Downscaling precipitation using regional climate models and circulation patterns toward hydrology. *Water Resources Research*, 47, 1–18.
- Chou, S.C., et al., 2012. Downscaling of South America present climate driven by 4-member HadCM3 runs. *Climate Dynamics*, 38 (3–4), 635–653.
- Ribeiro Neto, A., Montenegro, S.M.G.L., Silva, L.P.E., Cirilo, J.A. 2011. Impacts of Climate Change in Surface Runoff Using Regional Climate Model in Pernambuco State – Northeast of Brazil. Proceedings of XIV World Water Congress, Porto de Galinhas– Brazil, 2011.
- Wood, A.W., et al., 2002. Long-range experimental hydrologic forecasting for the eastern United States. *Journal of Geophysical Research*, 107 (D20), 6.1–6.15.

Thank you!

Gracias!

Obrigado!

ribeiront@gmail.com

