

SPSAS Climate Change

**São Paulo School of Advanced Science on Climate Change:
Scientific basis, adaptation, vulnerability and mitigation**

3-15 July 2017

Instituto de Astronomia, Geofísica e Ciências Atmosféricas,
Universidade de São Paulo, São Paulo, Brazil

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SPSAS Climate Change

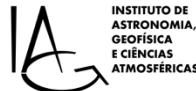
Information from paleoclimate archives

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EACH

School of Arts, Sciences and Humanities
University of São Paulo





- **Why study past climates?**
- **Main paleoclimate archives**
- **A selection of paleoclimate proxies**
- **Paleoclimate records: *the editor's cut***



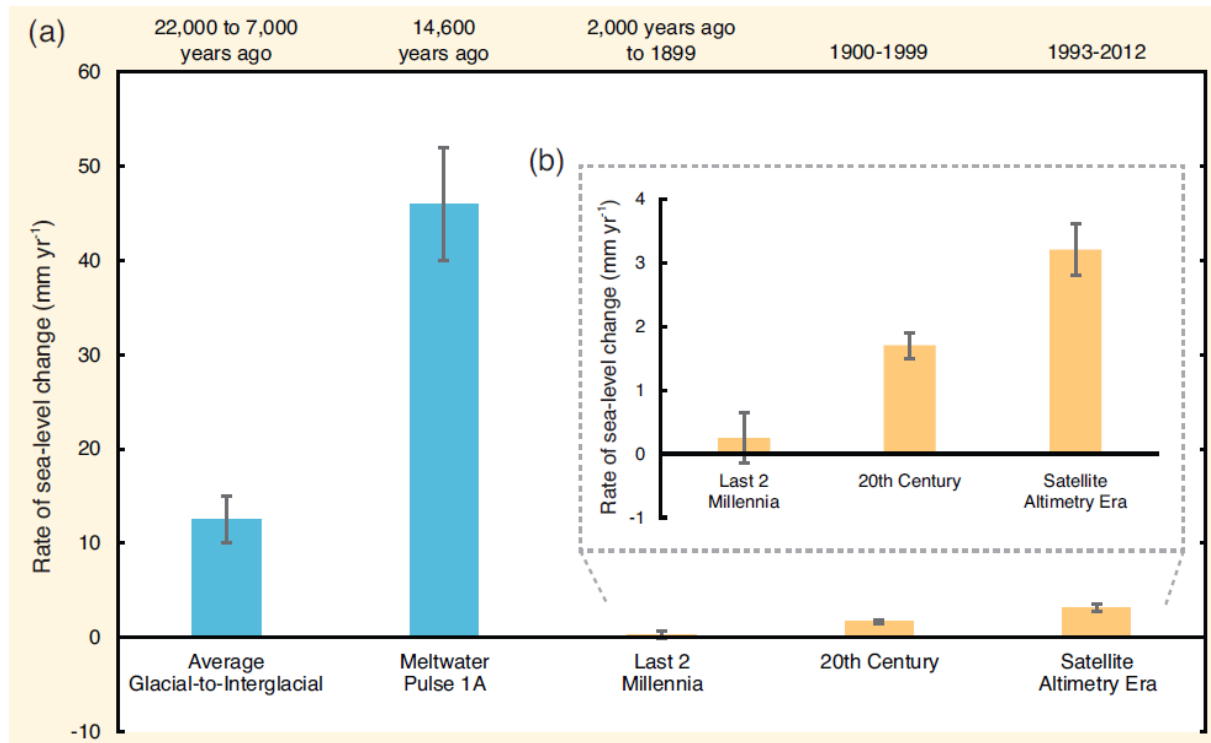
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Why study past climates?



1. Past climates provide a longer-term perspective than the instrumental record on understanding the controls, magnitudes, and spatial/temporal aspects of climate change



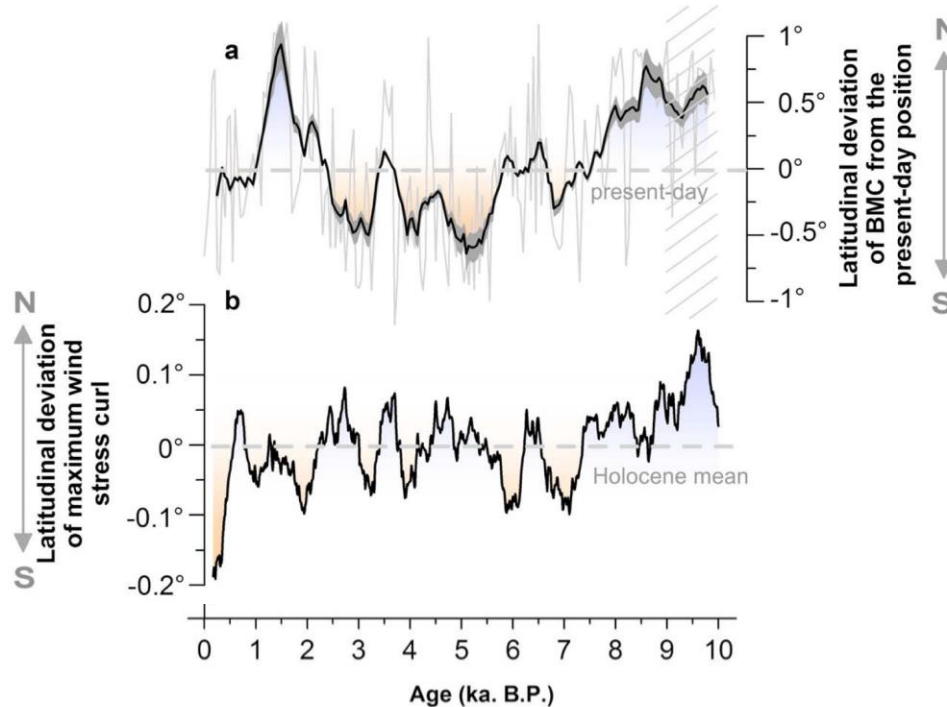
**Current rate of mean global sea level change is unusual relative to the last 2 kyr
Higher rates occurred, but especially during glacial-interglacial transitions**



Why study past climates?



2. Testing the accuracy (validating) of general circulation models by comparing hindcasts to paleoclimate archives



Mock (2007); Voigt et al. (2015)

Proxy-inferred and modeled SWW shifts compare qualitatively, but the model underestimates SWW variability by an order of magnitude

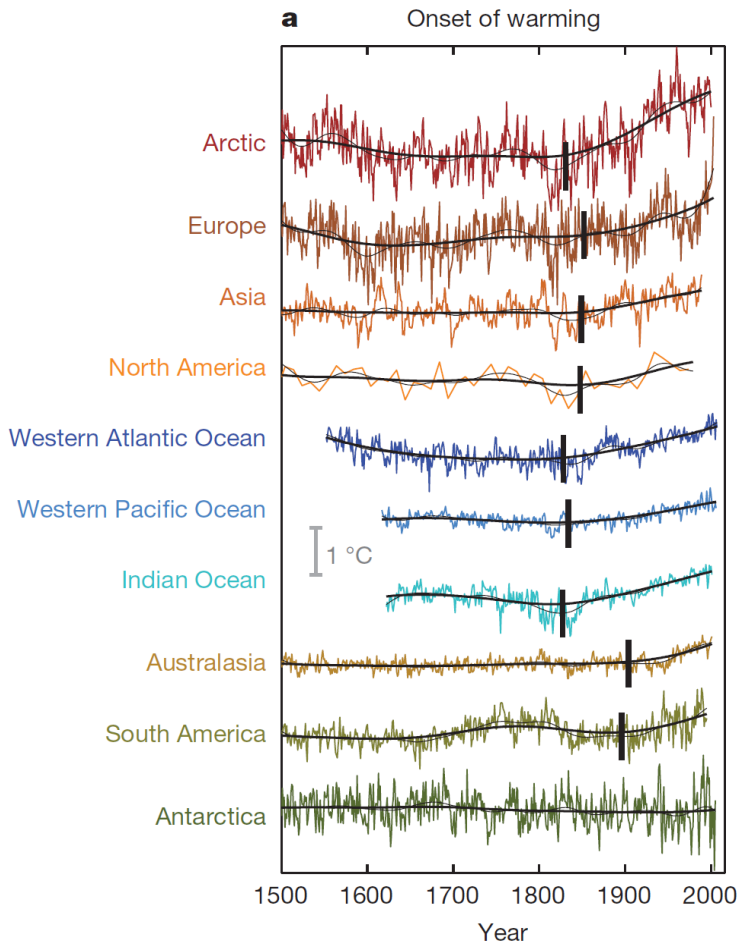
The underestimated natural variability implies a substantial uncertainty in mode projections of future SWW shifts



Why study past climates?



- 3. Past climates provide long records of natural climate variability that allow a clear understanding of natural climatic variability, and decoupling natural from anthropogenic variability



Instrumental records are too short to comprehensively assess anthropogenic climate change

In some regions, about 180 years of industrial-era warming has already caused surface temperatures to emerge above pre-industrial values

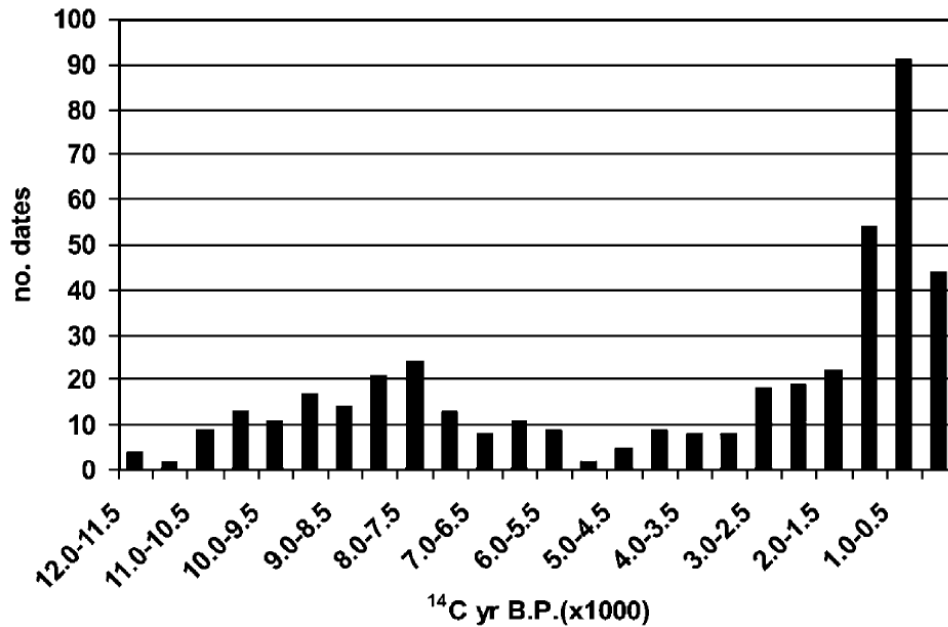


Why study past climates?

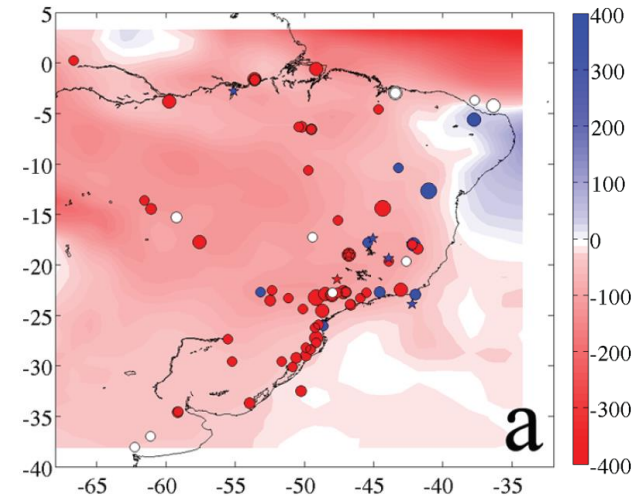


- 4. Past climates provide detailed records to understand the role of climate change on ecosystems and society

Archaeological ages for central Brazil



MH-PI precipitation anomalies from PMIP3/CMIP5 and paleoclimate archives



Mid Holocene depopulation in central Brazil is most likely related to a negative precipitation anomaly



Why study past climates?



nature
geoscience

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NATURE GEOSCIENCE | COMMENTARY

Climate change narratives

[Richard D. Pancost](#)

Nature Geoscience **10**, 466–468 (2017) | doi:10.1038/ngeo2981

Published online 26 June 2017

Reconstructions of Earth's past are much more than benchmarks for climate models

They also help us comprehend risk by providing concrete narratives for diverse climates



Why study past climates?



Paleoclimatology and paleoceanography are interdisciplinary fields of research that involve a multiscale (spatial and temporal) approach

Archaeology

Climatology

Ecology

Environmental chemistry

Geomorphology

Glaciology

Limnology

Paleontology

Palynology

Oceanography

Pedology

Sedimentology

Vulcanology



- Why study past climates?
- Main paleoclimate archives**
- A selection of paleoclimate proxies
- Paleoclimate records: *the editor's cut*



Main paleoclimate archives



Sediment cores



Ice cores



Tree rings



Corals



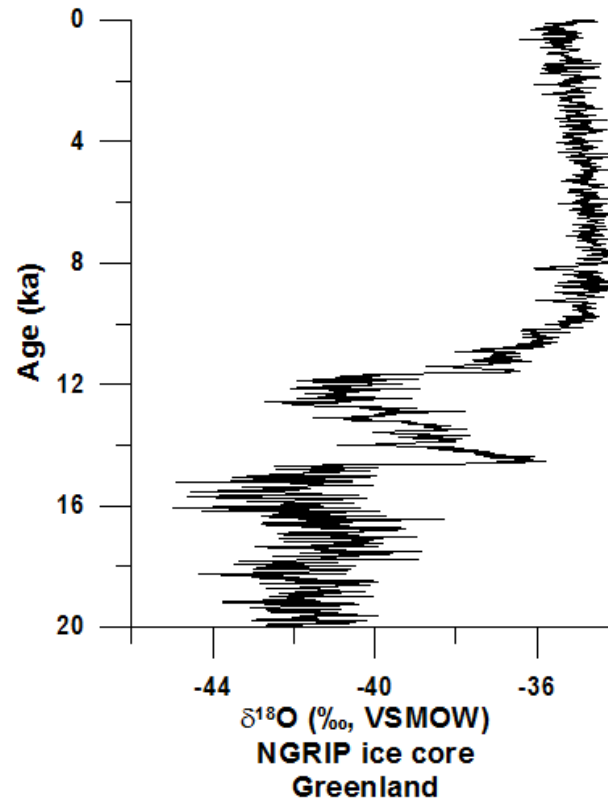
Speleothems





A good paleoclimate archive should meet certain criteria

1. Its physical, chemical or biological properties must represent environmental conditions





A good paleoclimate archive should meet certain criteria

2. It must be possible to determine the age of deposition/formation of the paleoclimate archive

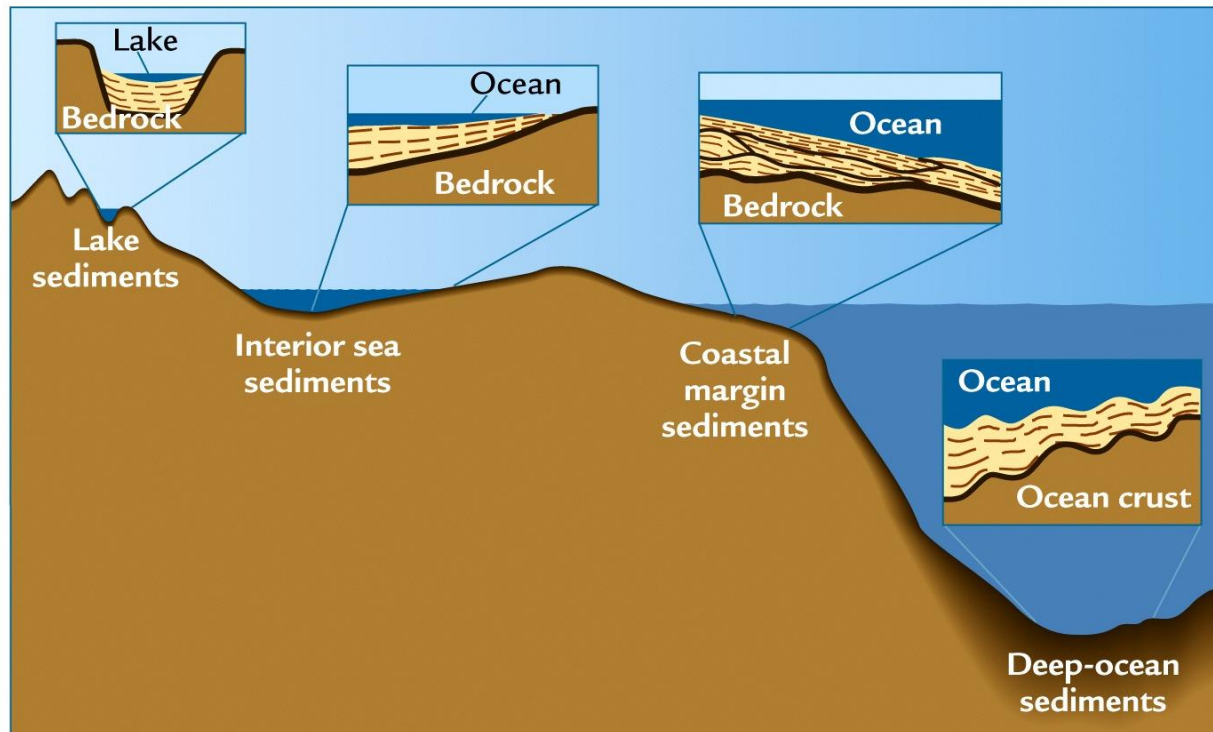




Main paleoclimate archives



Sediments are transported as particulate or dissolved matter and deposited in sedimentary basins (continental or marine) forming potential paleoclimate archives

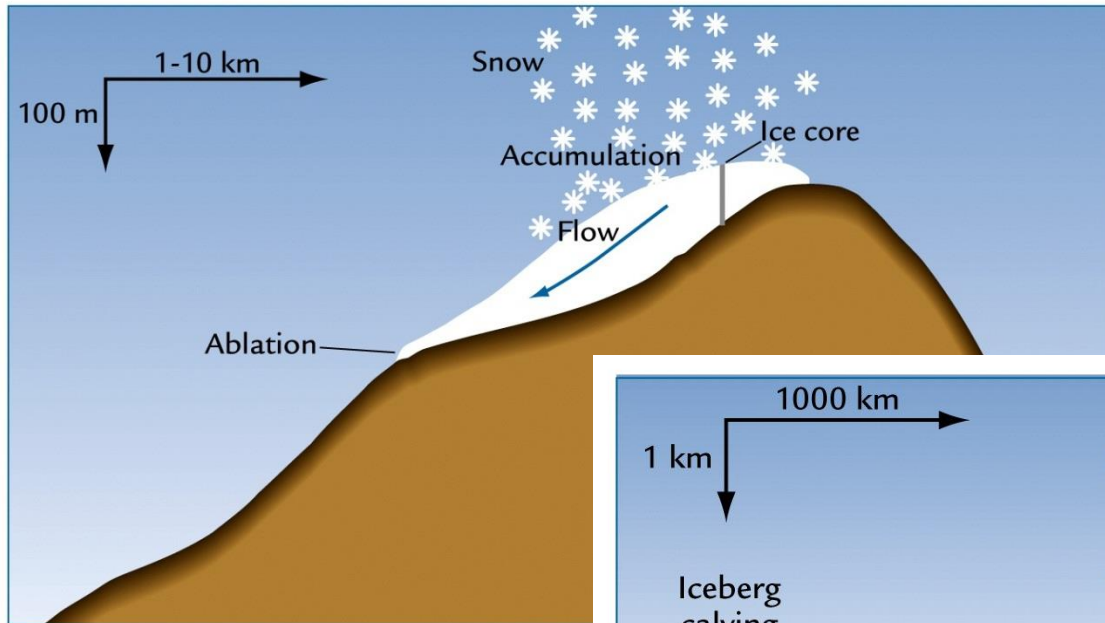




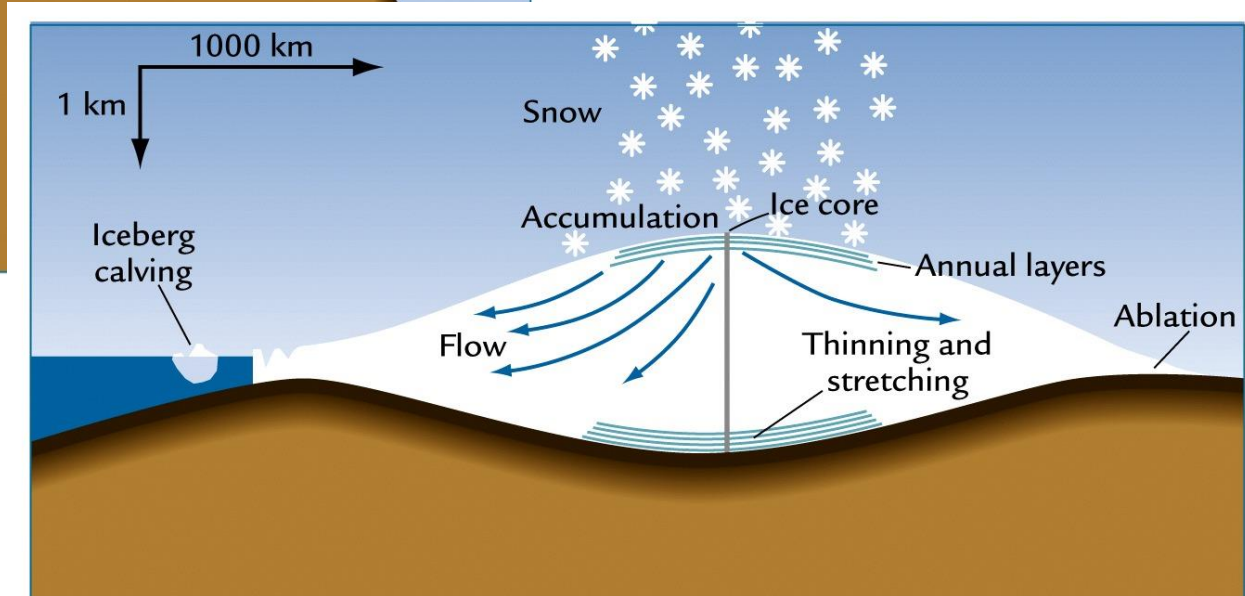
Main paleoclimate archives



Snow accumulates in cold regions to form ice and accumulation mainly occurs in high latitudes or altitudes potentially giving rise of paleoclimate archives



Mountain glaciers



Continental glaciers



Main paleoclimate archives



Tropical and subtropical clear waters may be suitable for the growth of corals and some build layered skeletons suitable as paleoclimate archives





Main paleoclimate archives



Trees growth by adding rings to the trunk, and mid-latitude trees under strong seasonality show great potential as paleoclimate archives

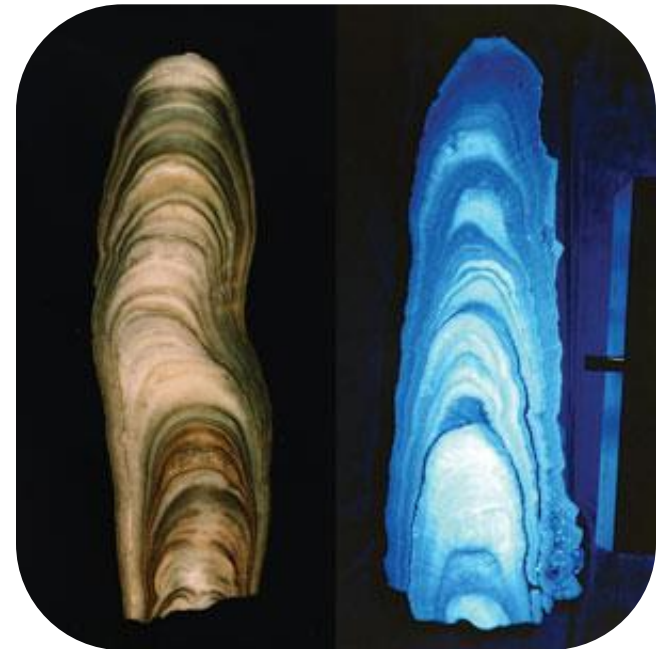




Main paleoclimate archives



Stalagmites growth by the deposition of layers; in caves with stable environmental conditions, the chemical composition of these layers may record regional climate parameters constituting good paleoclimate archives





Main paleoclimate archives – *an insight into marine cores*





Gravity corer

- Up to ~12m
- No water depth limitation
- “Easy” operation
- High efficiency
- Large diameter (~12cm)
- Short ship time





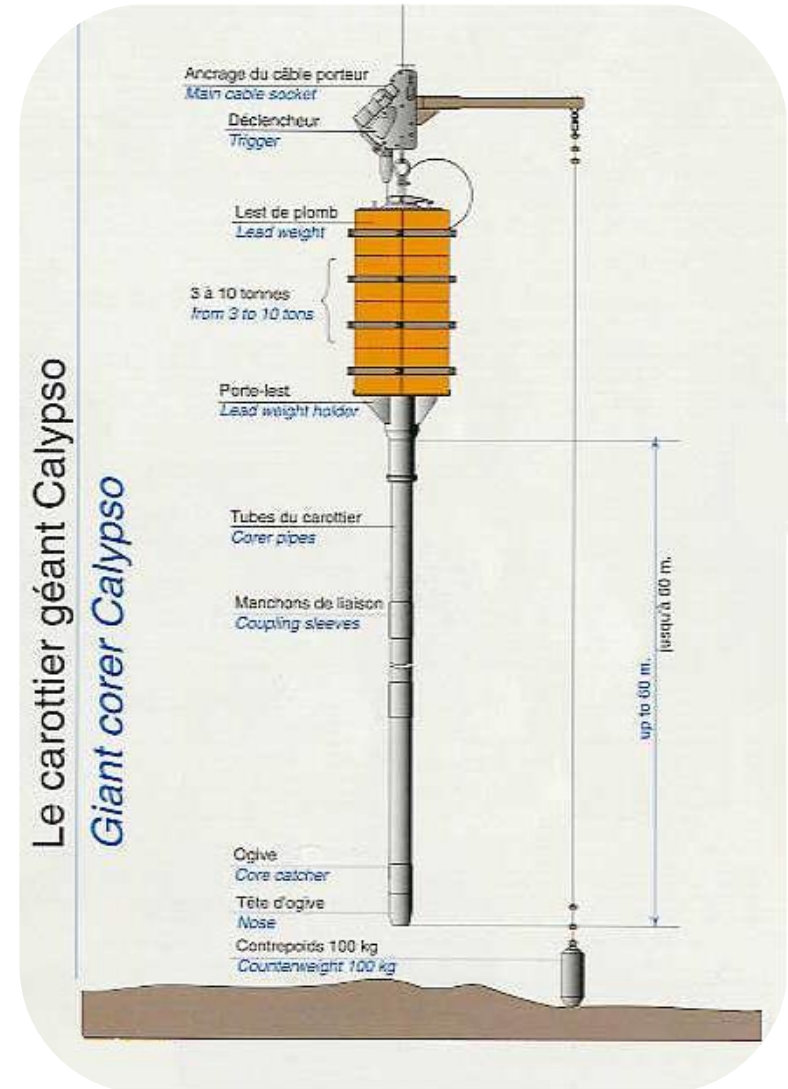
Main paleoclimate archives – *an insight into marine cores*





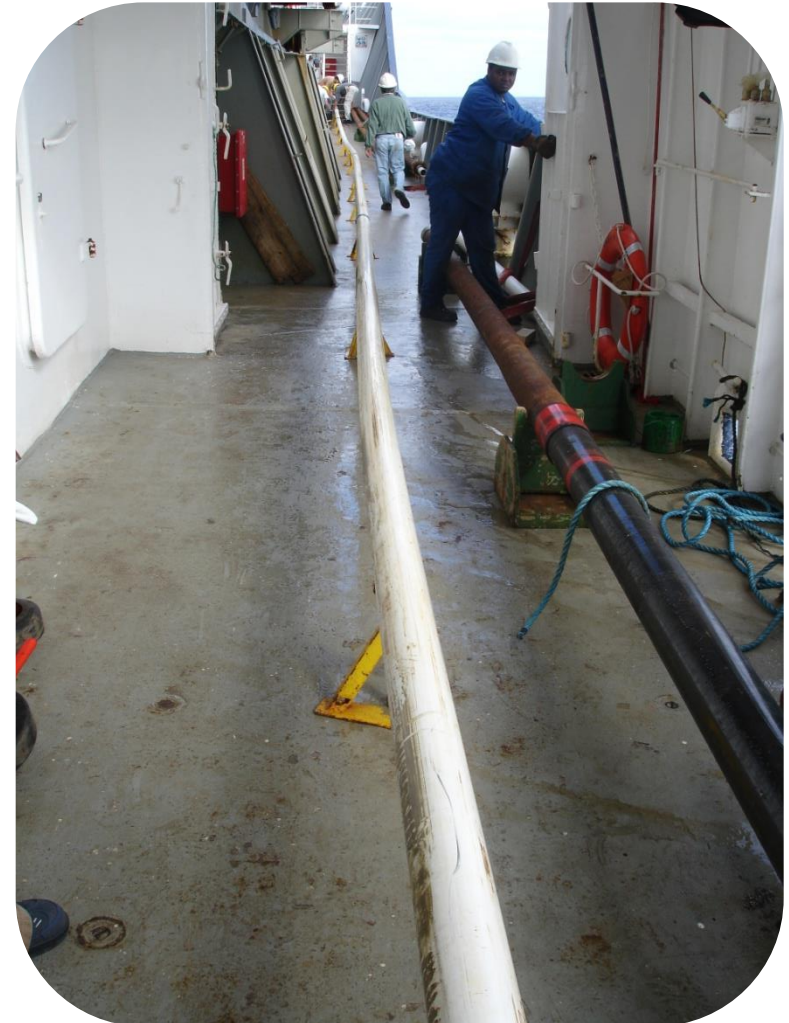
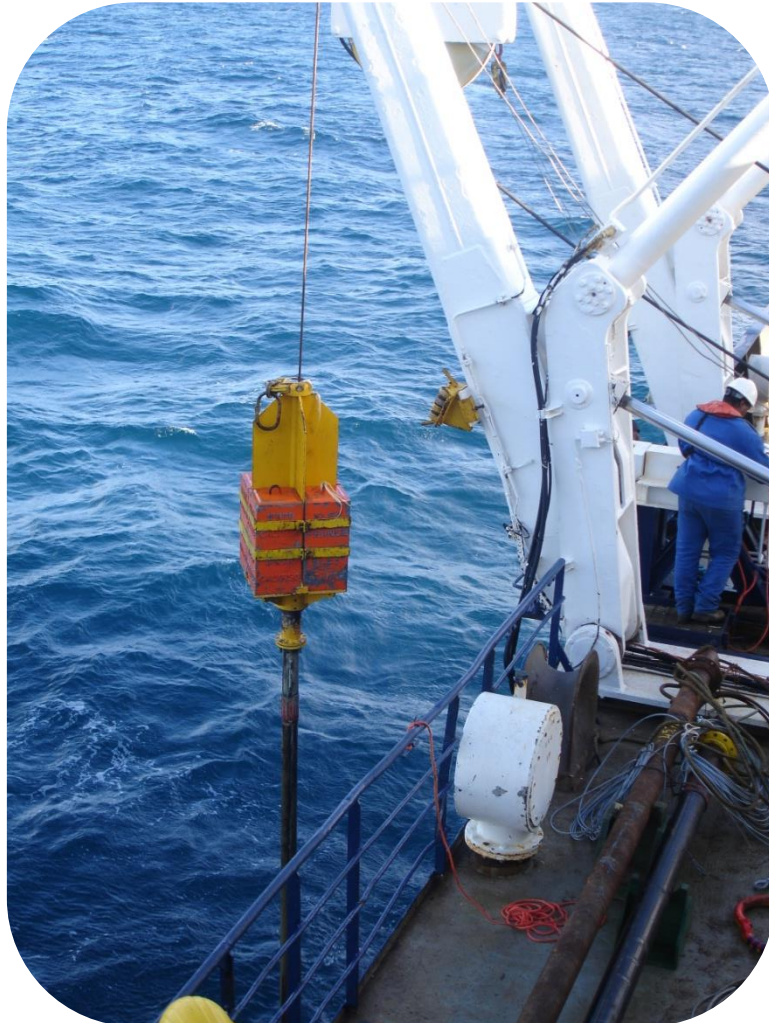
Giant piston corer (Calypso)

- Up to ~65m
- No water depth limitation
- Relatively complex operation
- Medium to high efficiency
(gaps)
- Large diameter (~12cm)
- Short ship time





Main paleoclimate archives – *an insight into marine cores*





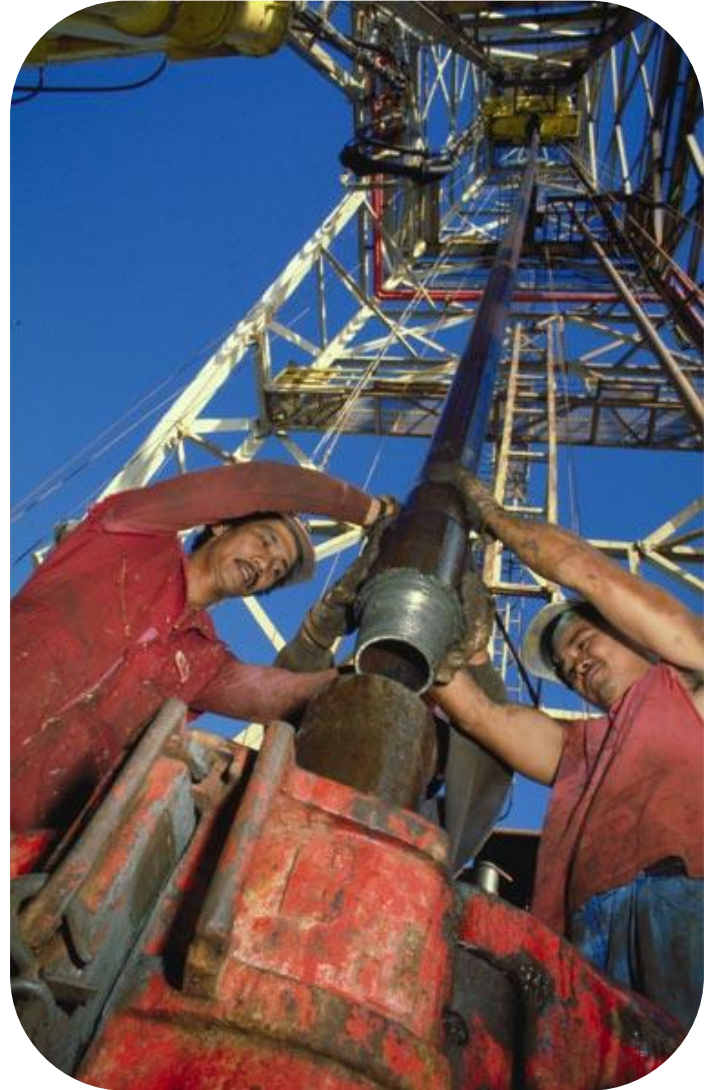
Main paleoclimate archives – *an insight into marine cores*





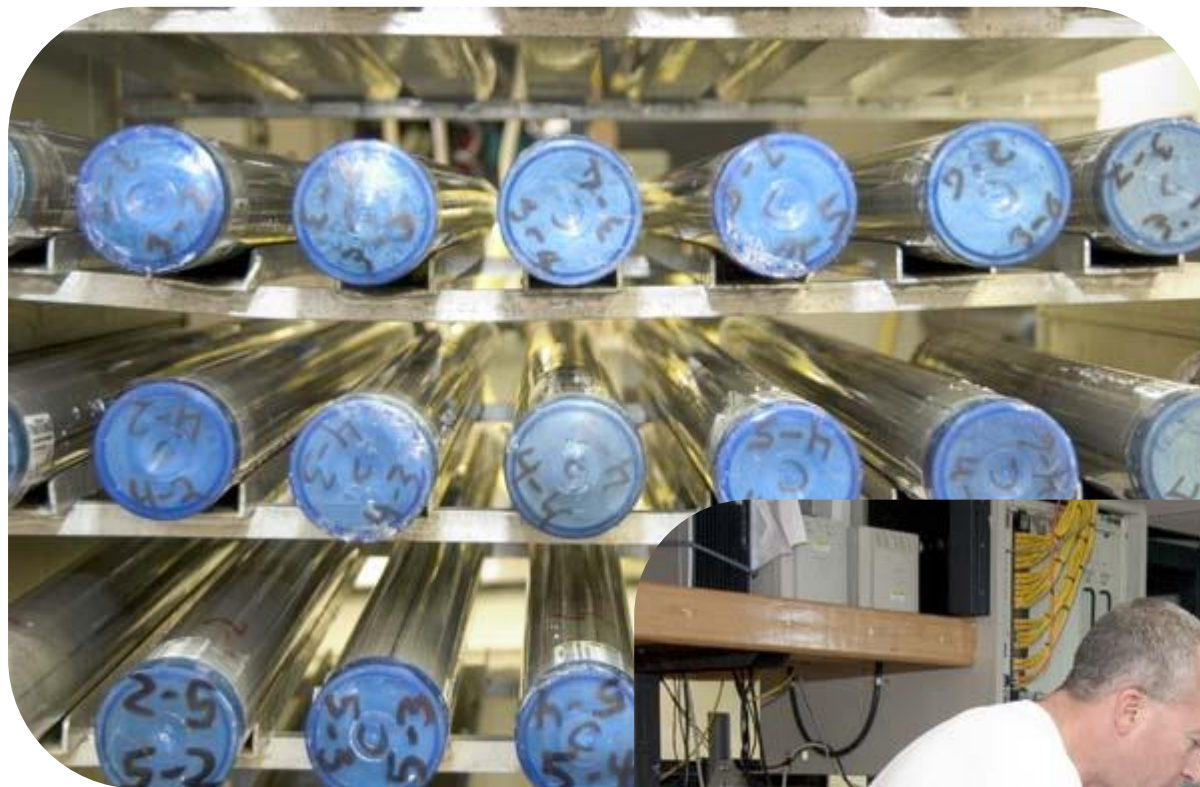
Drilling device (ship based)

- Hundreds of m
- No water depth limitation
- Highly complex operation
- Medium to high efficiency
- Small diameter (~6cm)
- Long ship time





Main paleoclimate archives – *an insight into marine cores*

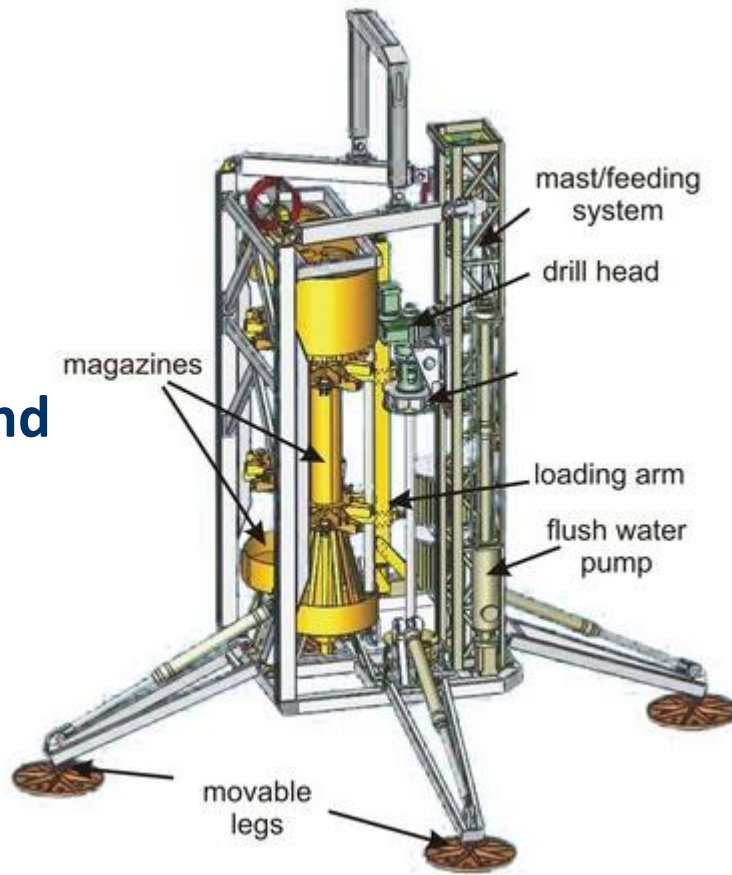




Main paleoclimate archives – *an insight into marine cores*

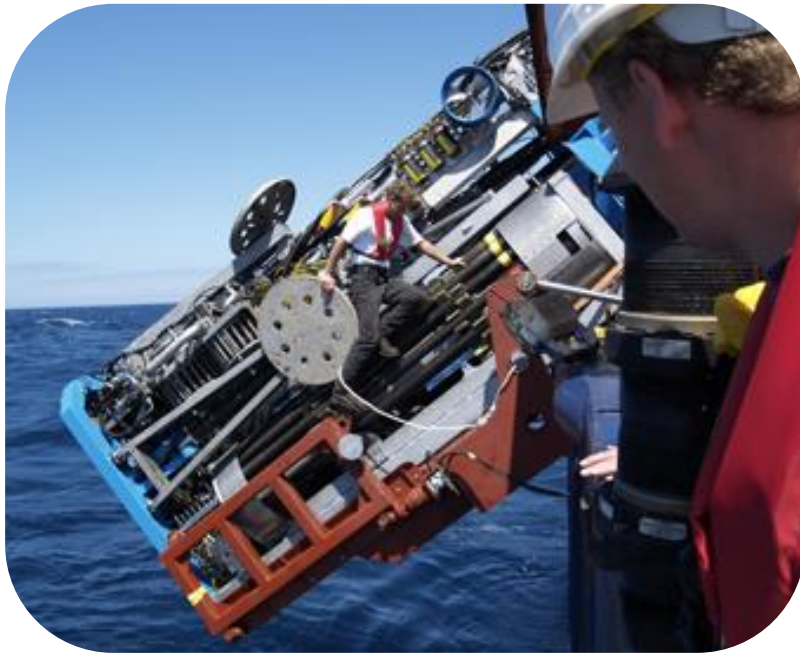
Drilling device (in situ)

- Up to 70 m
- Operation between 200 and 2000m
- Medium to high complex operation
- Medium to high efficiency
- Small diameter
- Long ship time





Main paleoclimate archives – *an insight into marine cores*



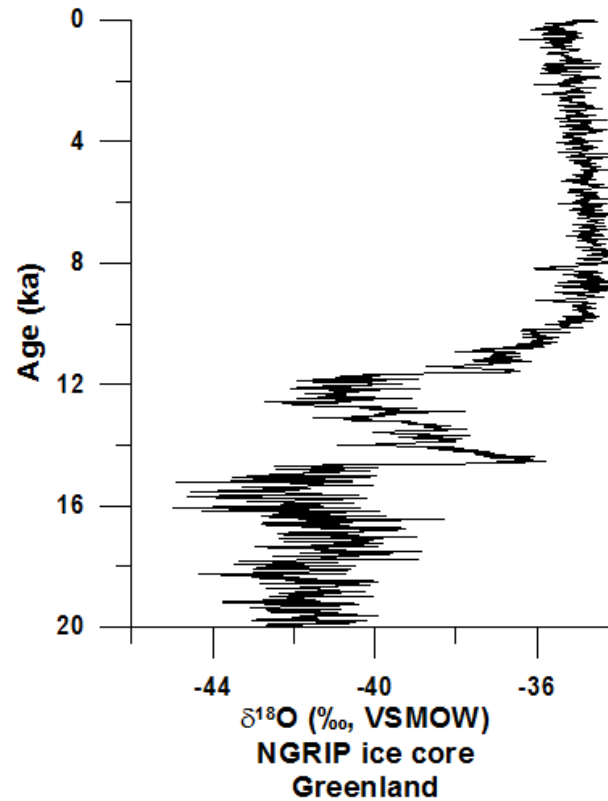


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- **A selection of paleoclimate proxies**
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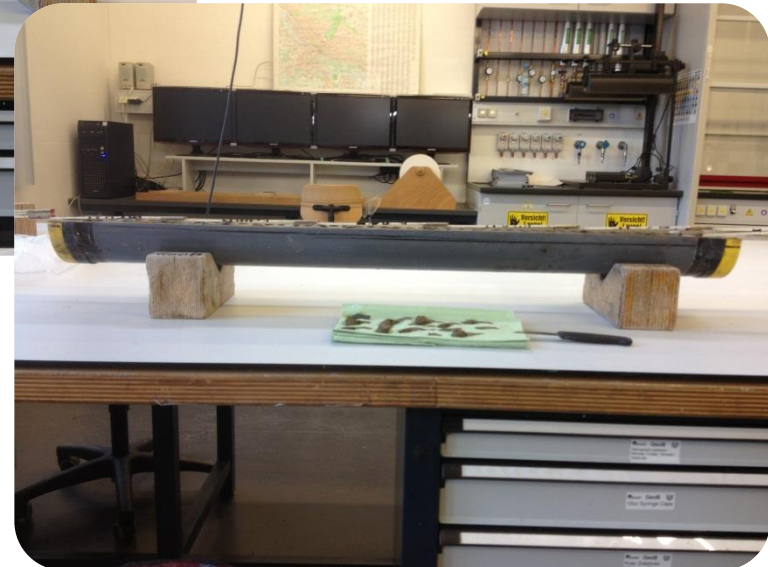
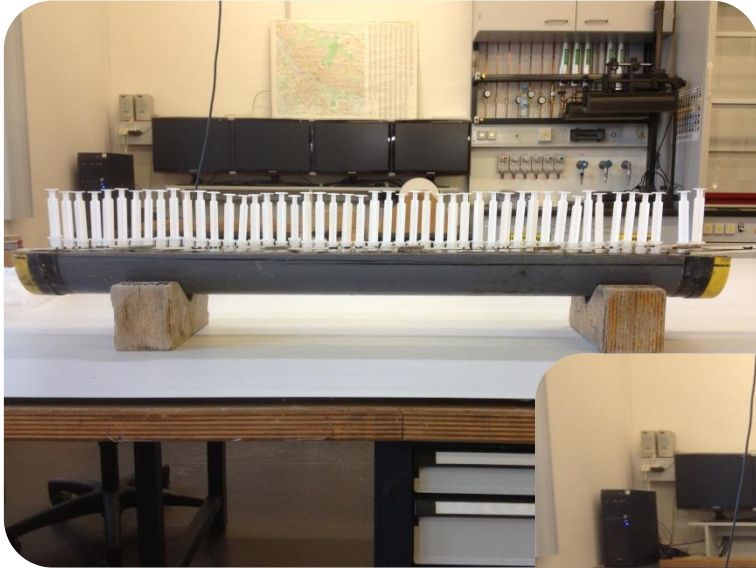
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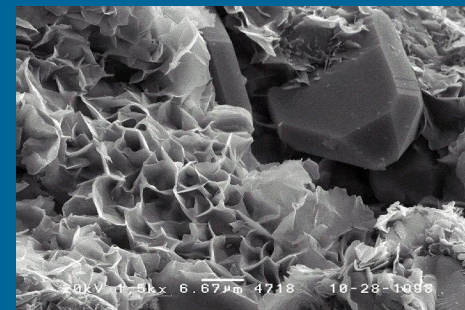
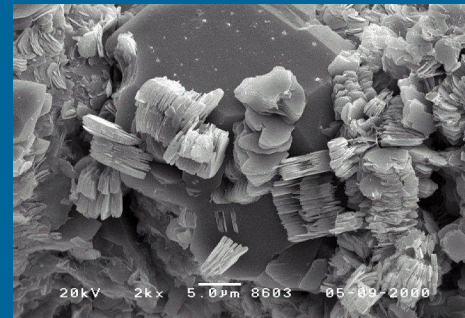
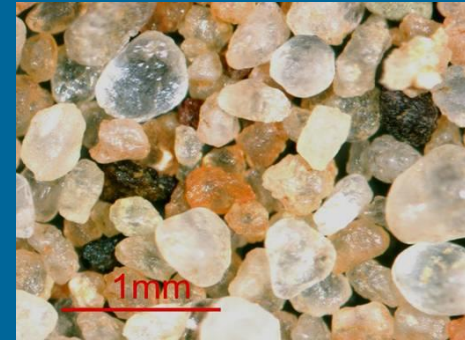
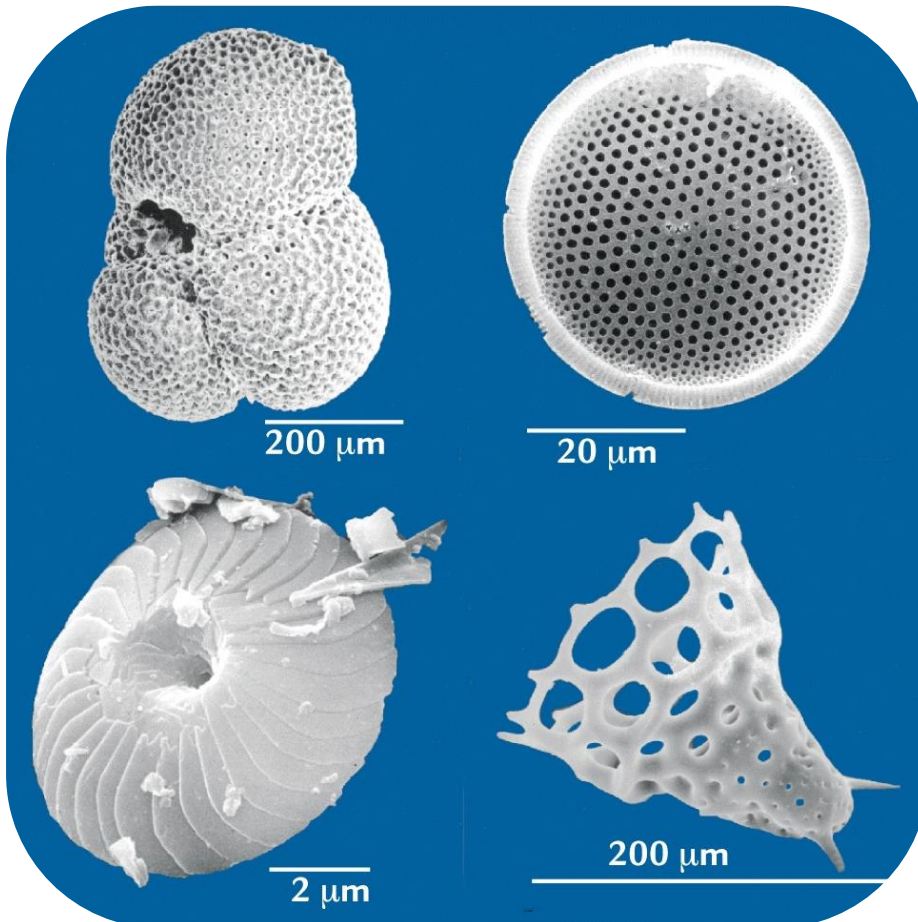


A selection of paleoclimate proxies



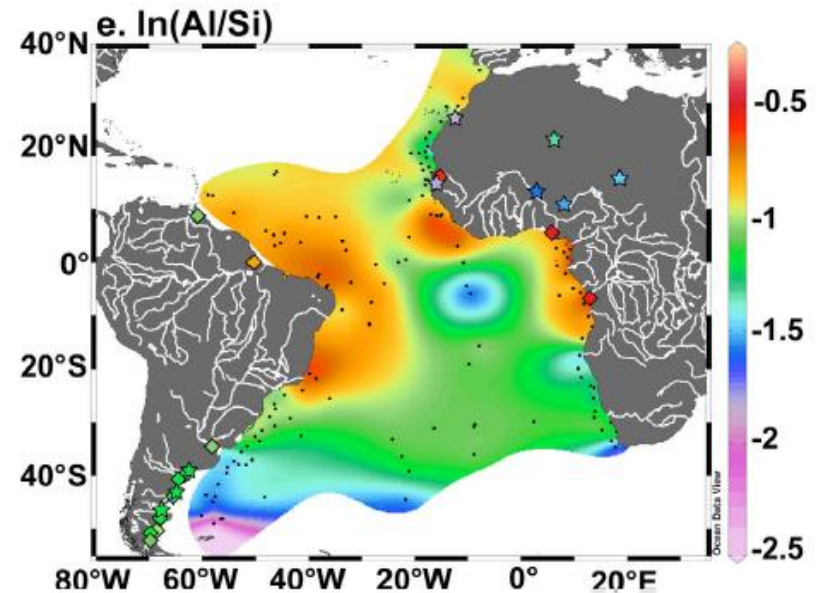
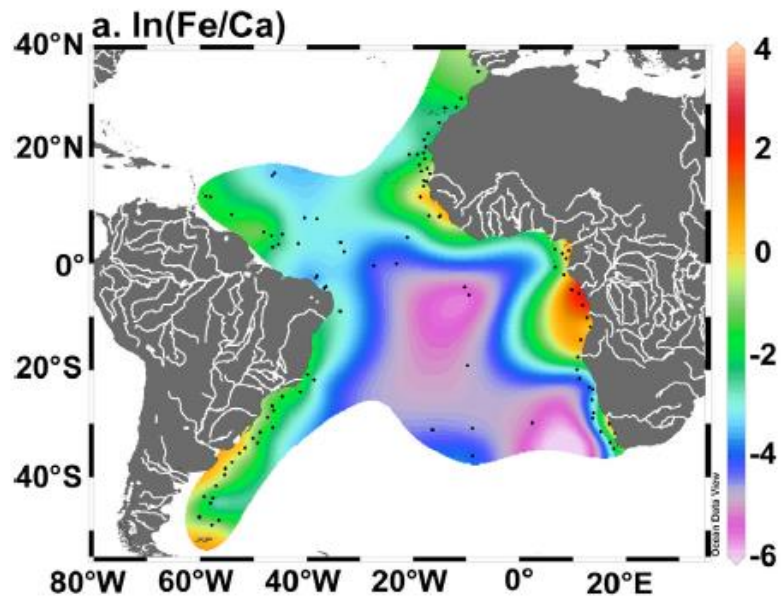


Autochthonous vs. alloctonous sediments



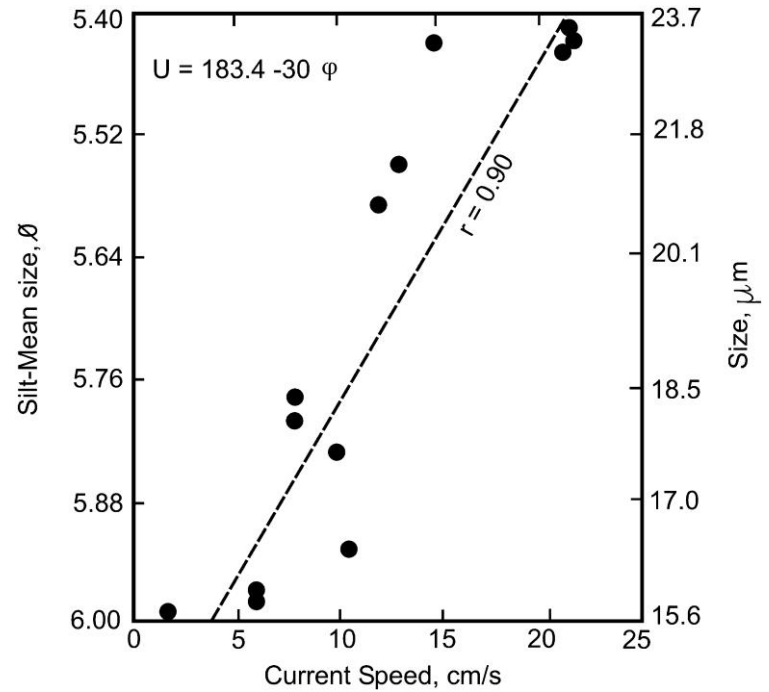
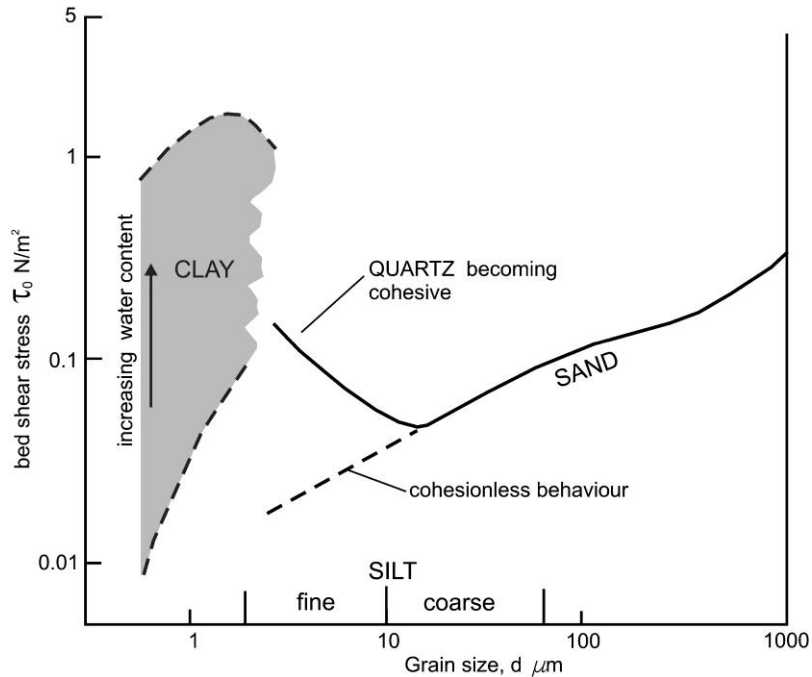


- Major elements in bulk sediment





• Grain size distribution of the terrigenous fraction

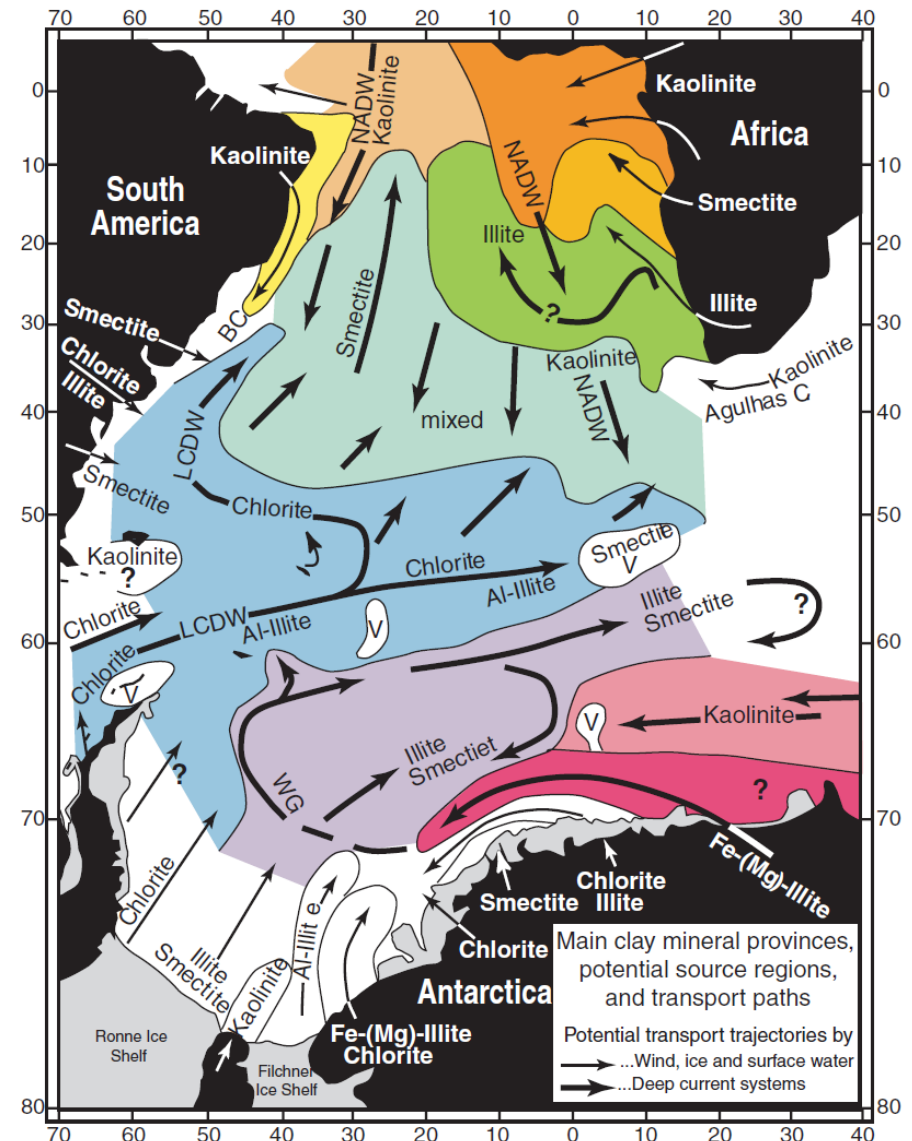
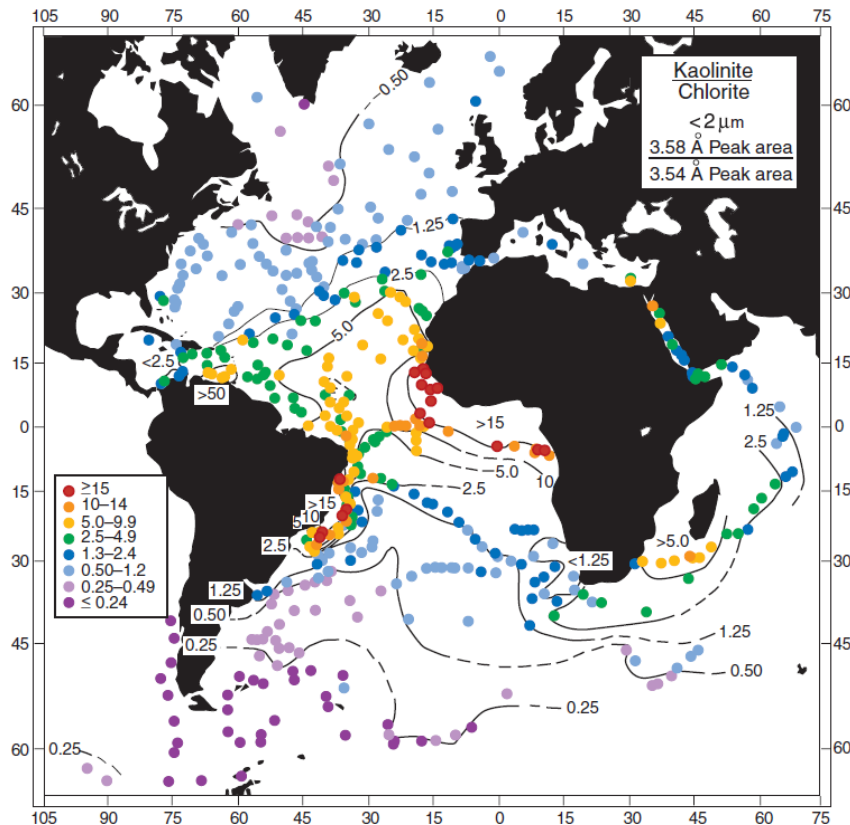




A selection of paleoclimate proxies



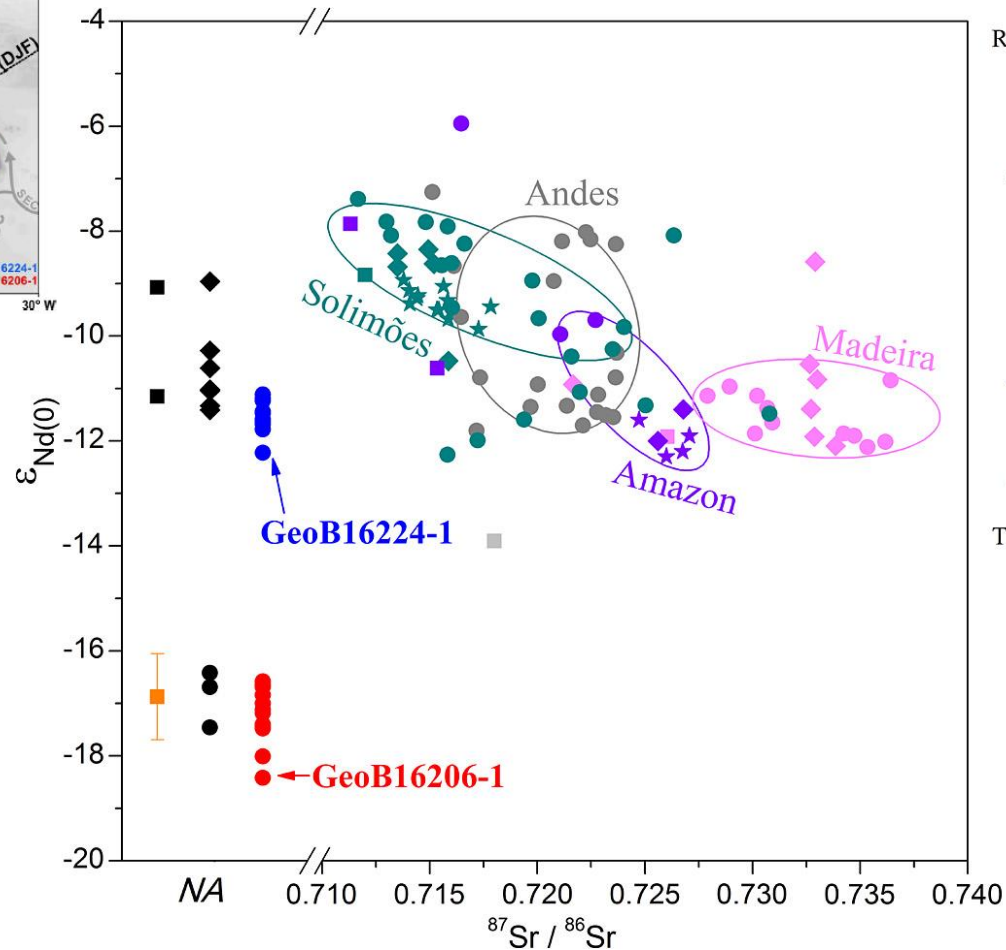
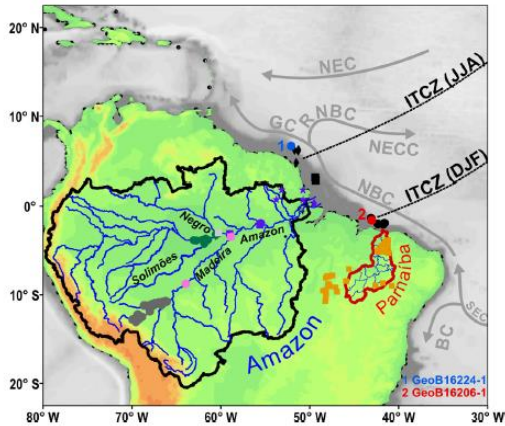
•Clay minerals assemblage



Biscaye (1965), Petschick et al. (1996)



•Nd isotopes in the terrigenous fraction



Reference terrestrial samples

- Andes (Basu et al., 1990)
- Solimões (Allègre et al., 1996)
- ★ Solimões (Viers et al., 2008)
- ◆ Solimões (Bouchez et al., 2011)
- Solimões (Horbe et al., 2014)
- Negro (Allègre et al., 1996)
- Amazon (Allègre et al., 1996)
- Amazon (Bouchez et al., 2011)
- ★ Amazon Plateau (Parra et al., 1997)
- ◆ Amazon Delta (Parra et al., 1997)
- Madeira (Allègre et al., 1996)
- Madeira (Viers et al., 2008)
- ◆ Madeira (Bouchez et al., 2011)

This study

Near-core-top sediments

- off the Parnaíba River mouth
- off the Amazon River mouth
- ◆ off French Guiana

Geological sediments

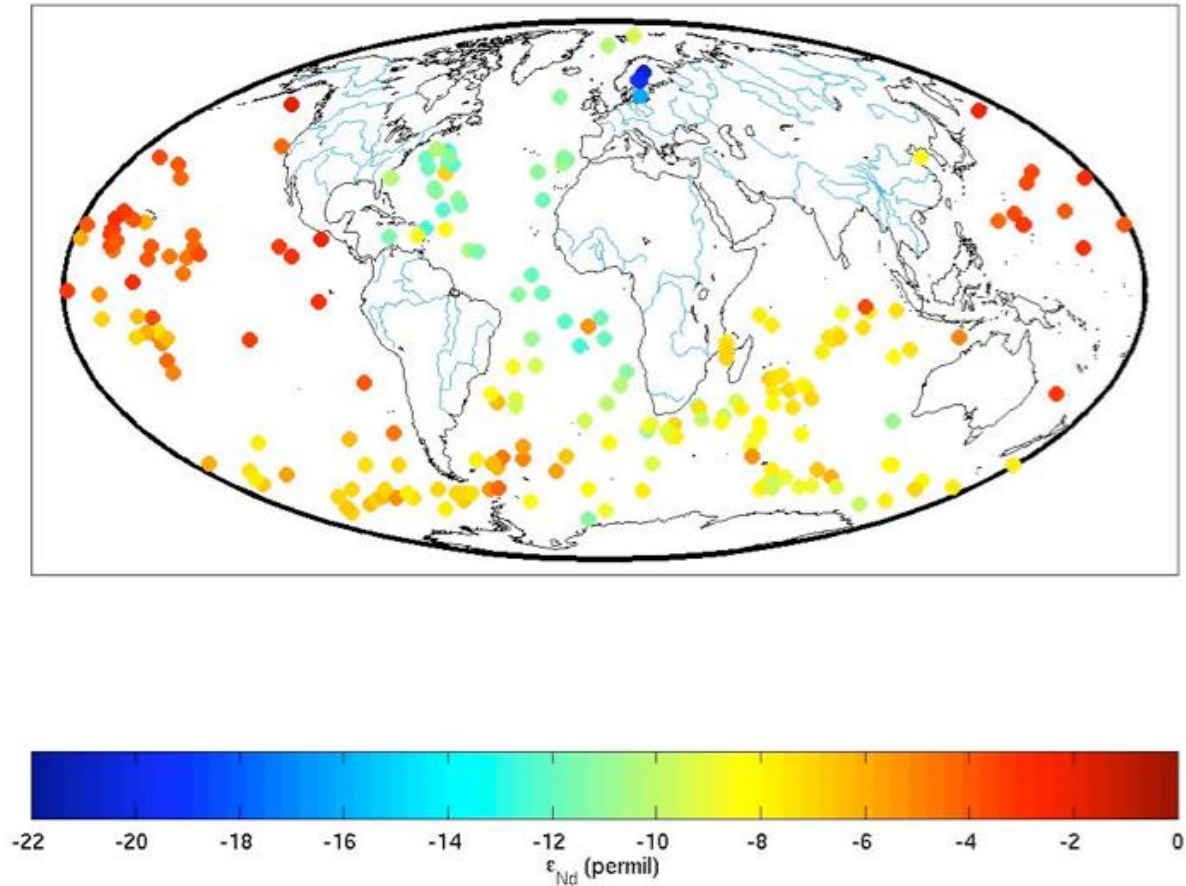
- Parnaíba River drainage

Downcore sediments

- GeoB16224-1
- GeoB16206-1



- Nd isotopes in autochthonous fractions

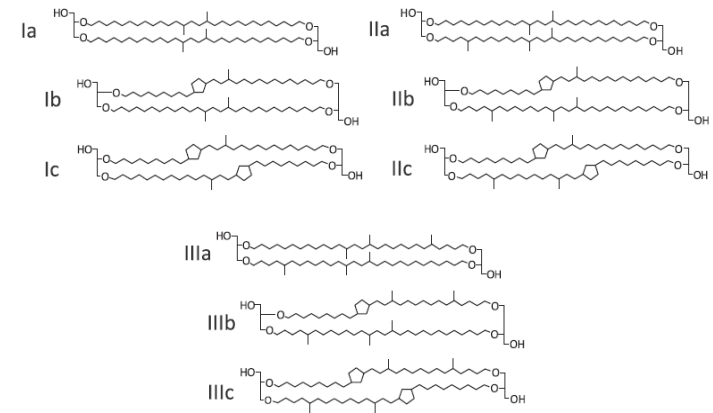
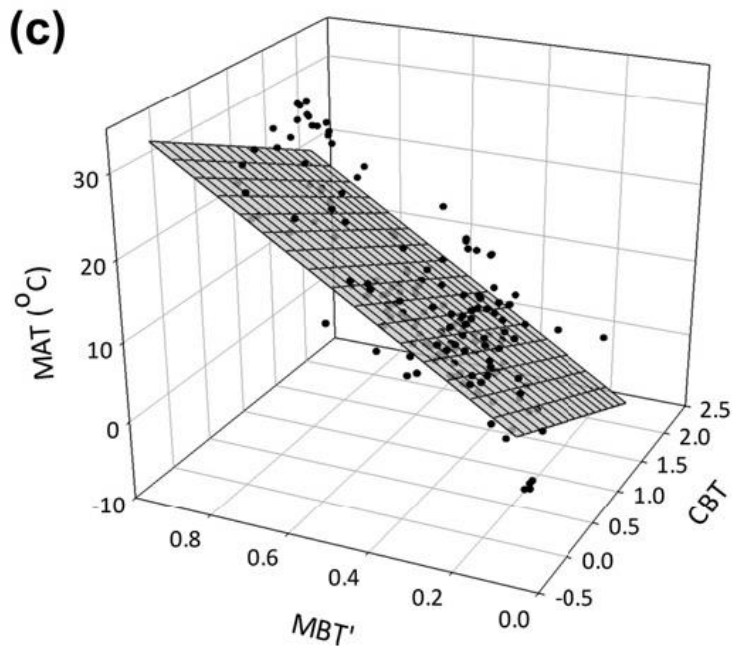




• Biomarkers

MATs reconstructed via branched glycerol dialkyl glycerol tetraethers (GDGTs) analyses

Global soil calibration



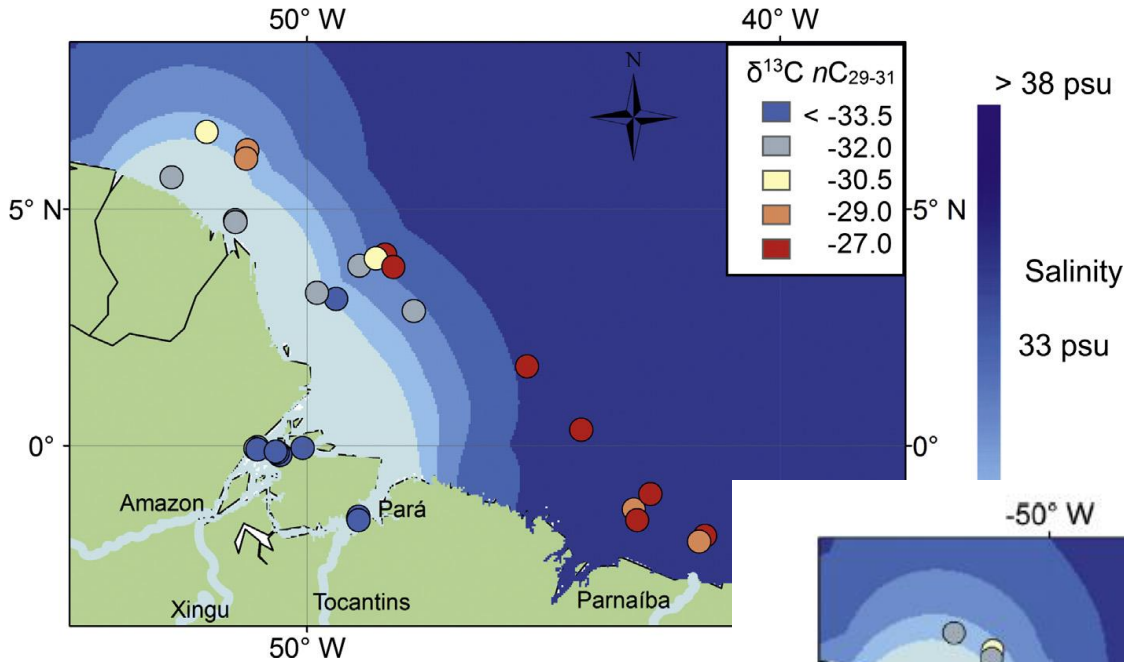
Branched GDGTs



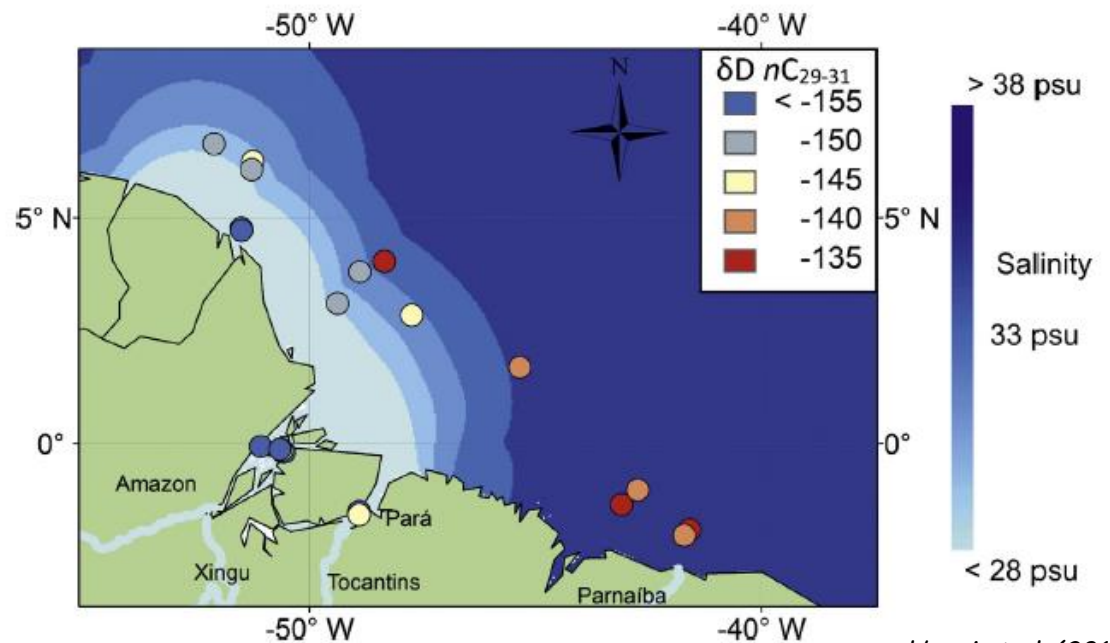
A selection of paleoclimate proxies



• Biomarkers



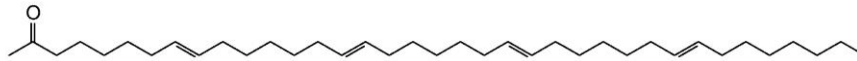
*Long-chain n-alkane
(plant wax) $\delta^{13}\text{C}$ and
 δD*



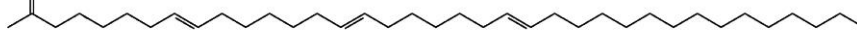


•Biomarkers

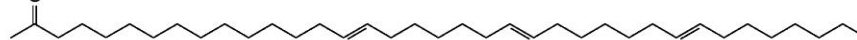
A) Alkenones:



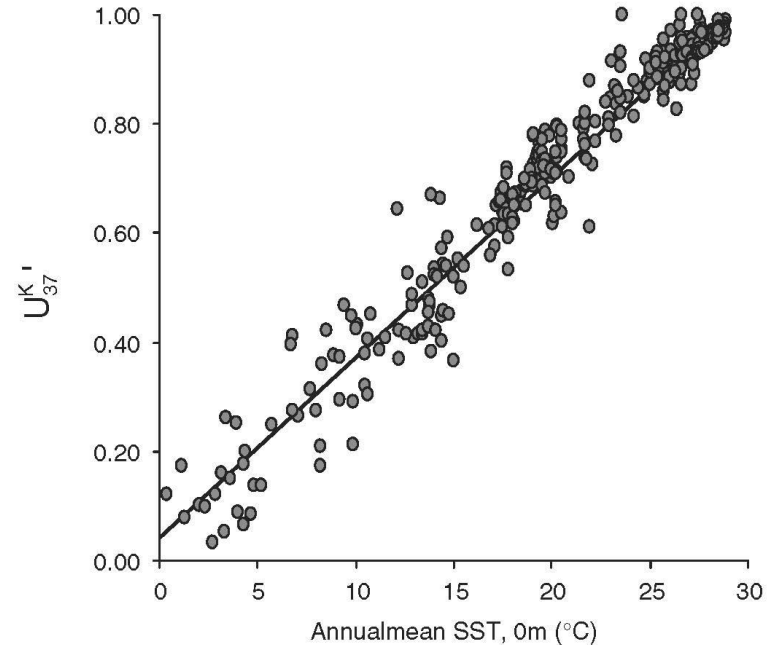
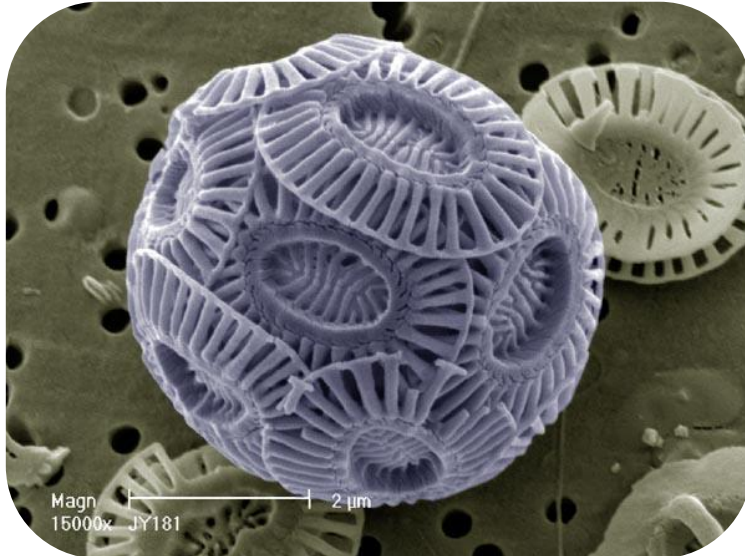
heptatriaconta-8E,15E,22E,29E-tetraen-2-one *or* C_{37:4}Me



heptatriaconta-8E,15E,22E-trien-2-one *or* C_{37:3}Me

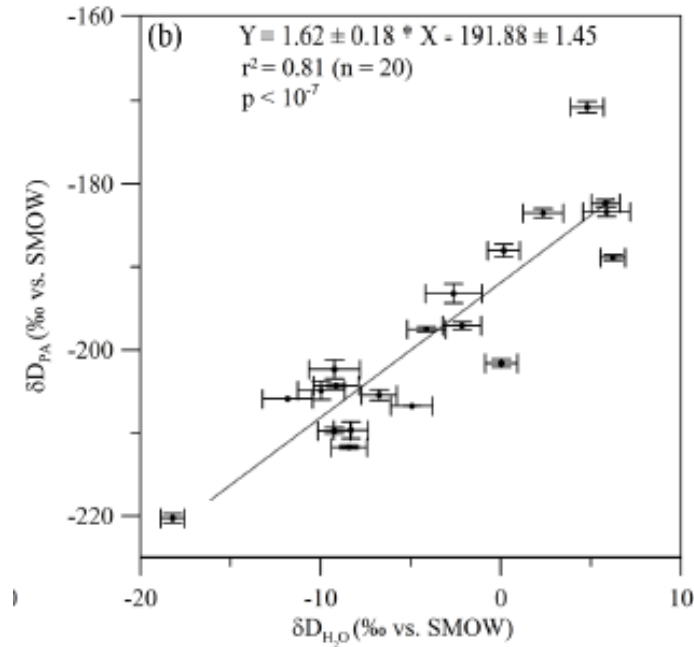


heptatriaconta-15E,22E-dien-2-one *or* C_{37:2}Me

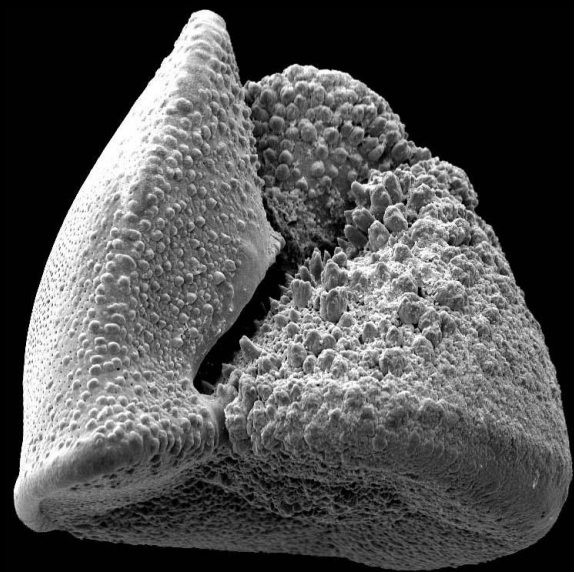
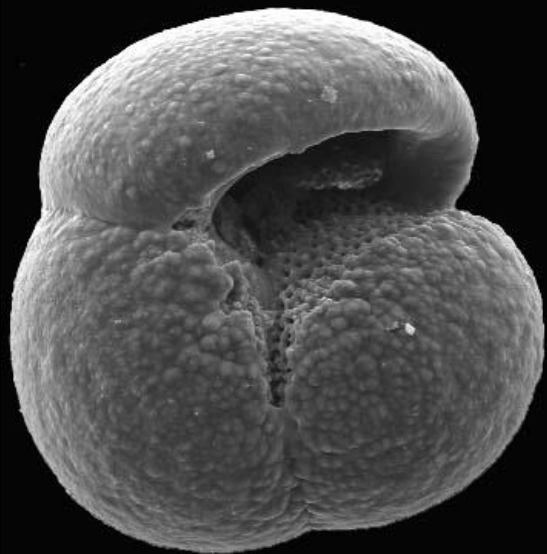
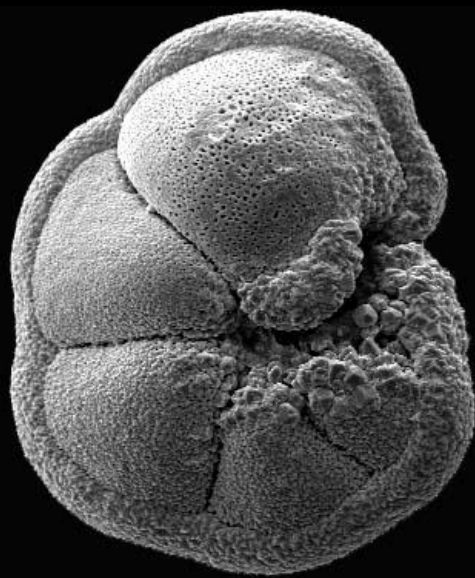
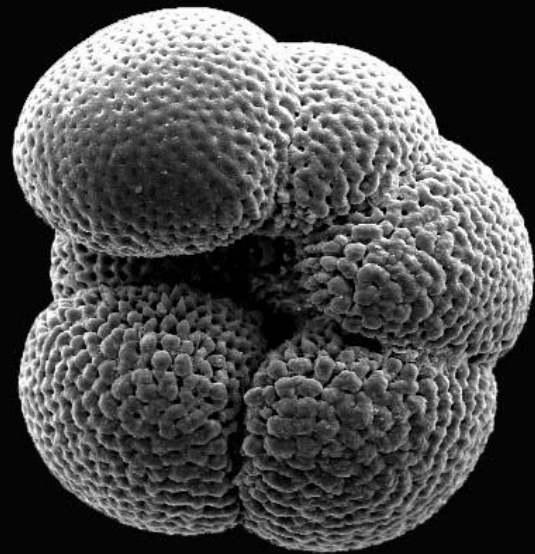
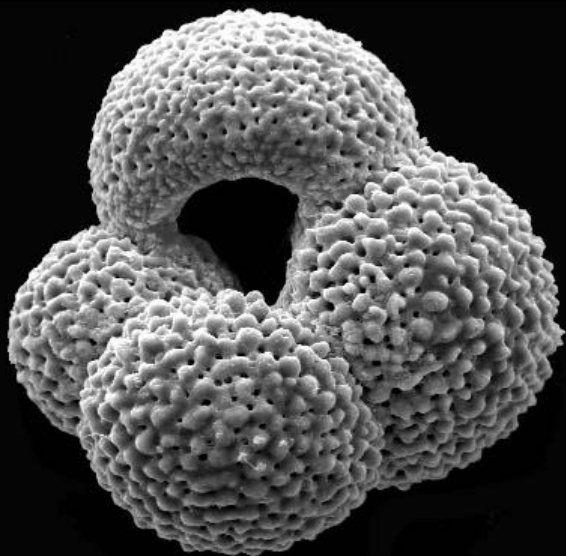
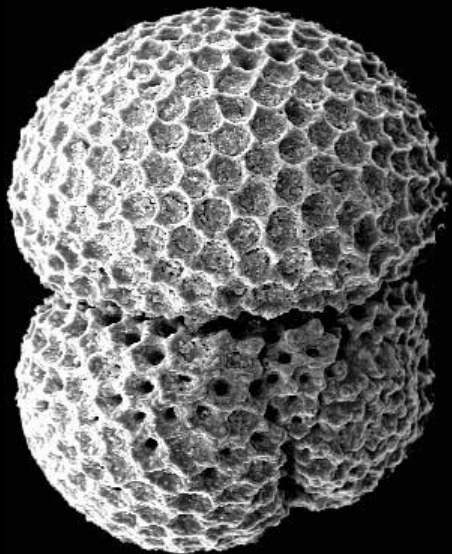




•Biomarkers



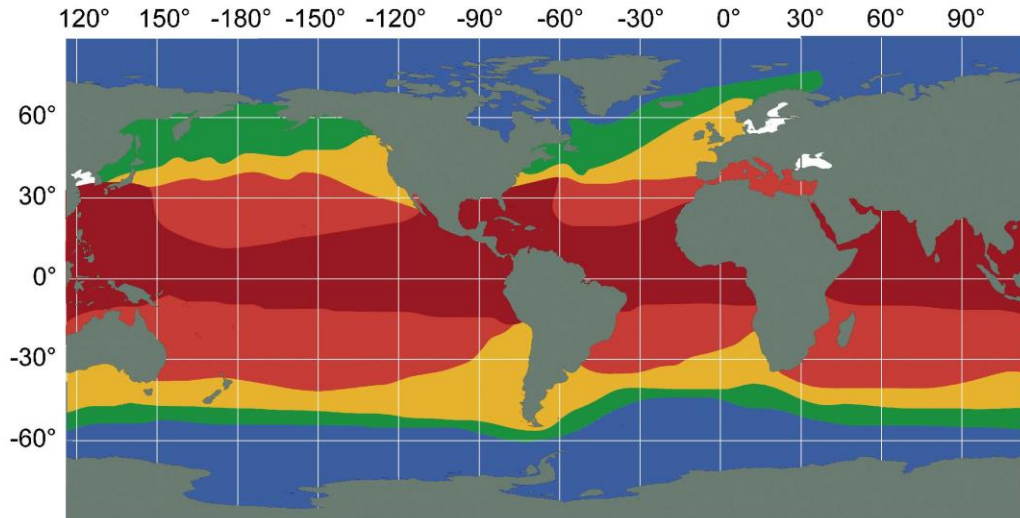
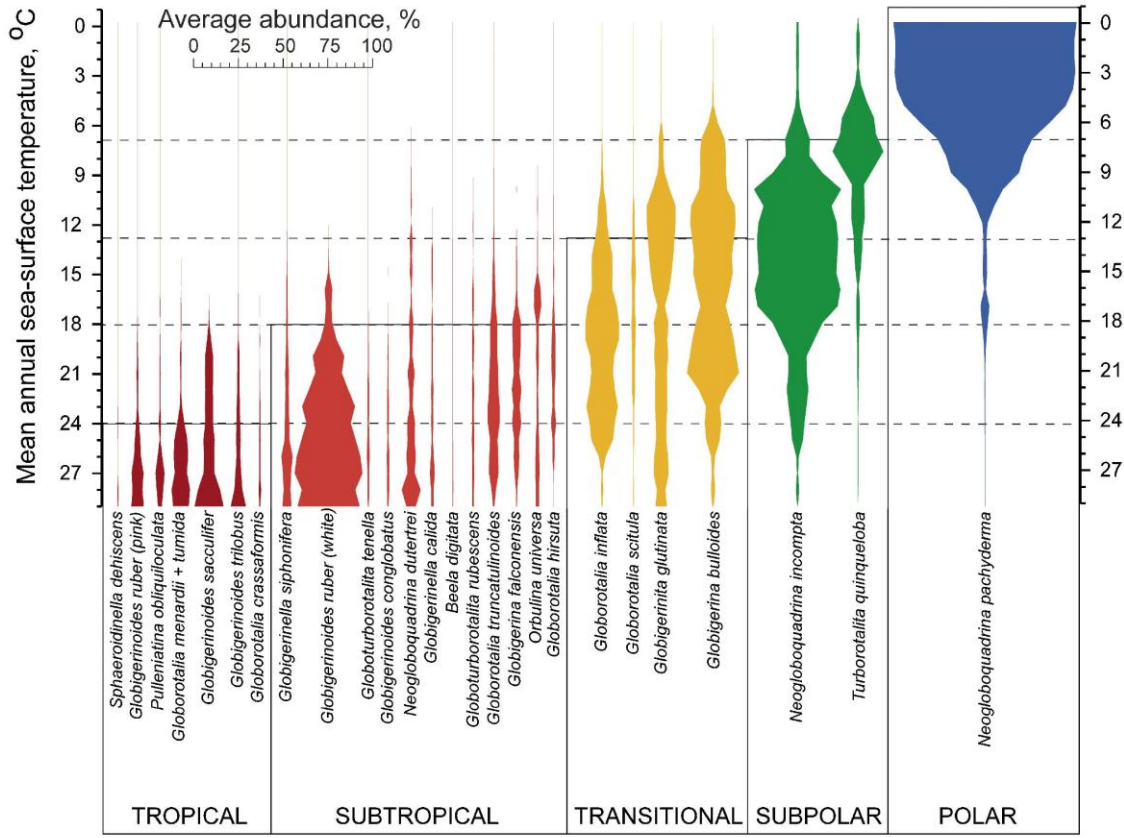
*Palmitic acid (an
unsaturated fatty acid
highly abundant in
most aquatic
environments) δD*



100 μm



•Planktonic foraminifera



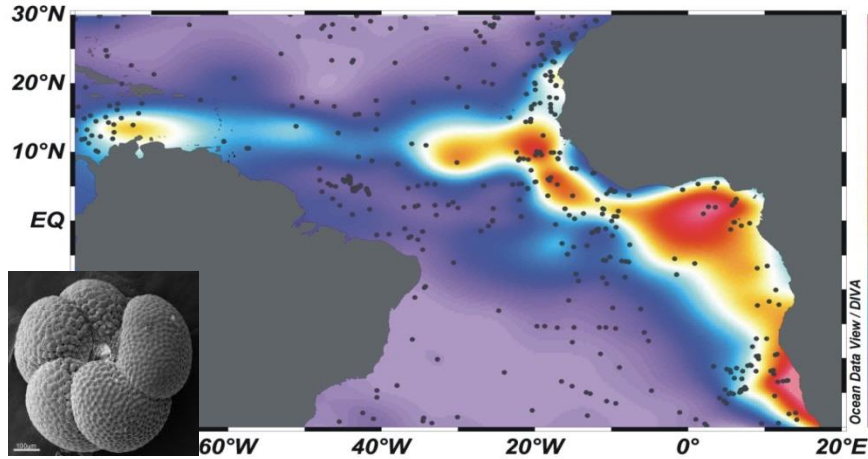


A selection of paleoclimate proxies

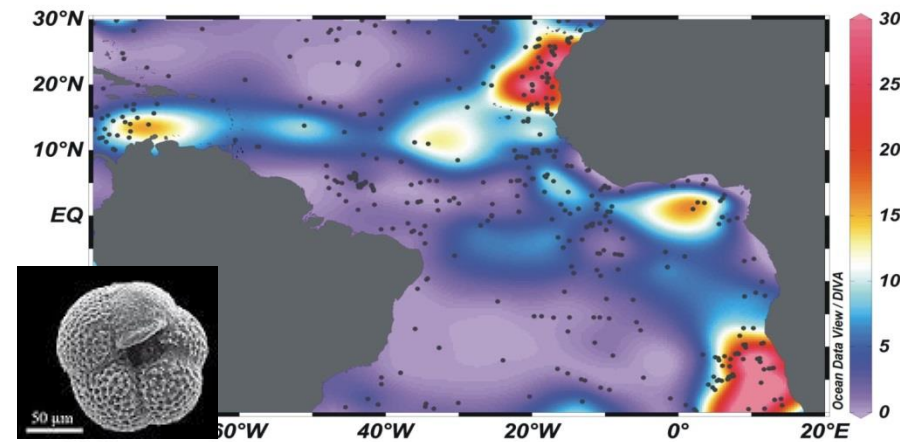


• Planktonic foraminifera

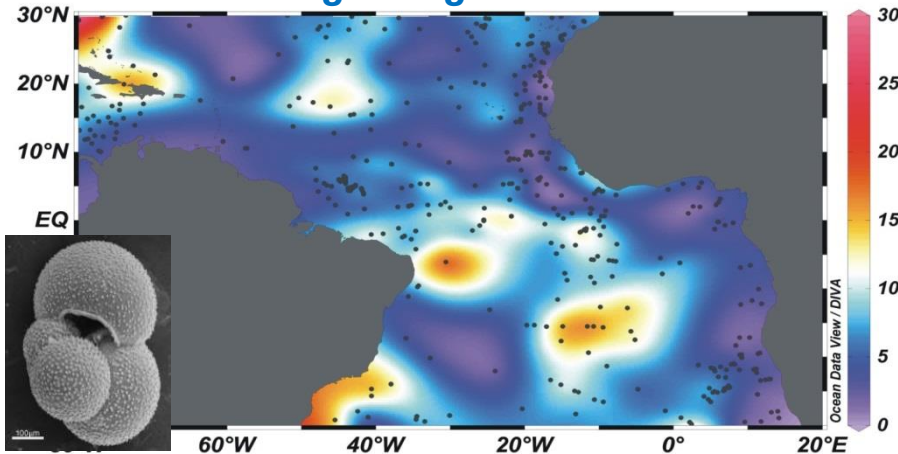
Neogloboquadrina dutertrei



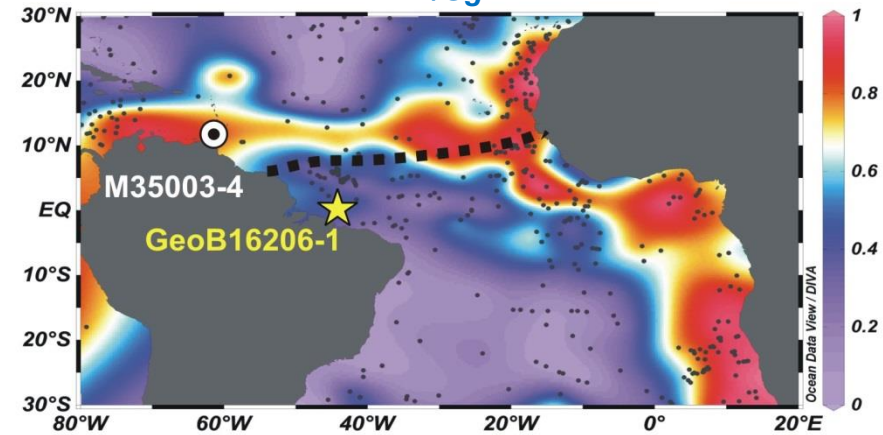
Neogloboquadrina incompta



Globigerina glutinata



$R_{N/Gg}$



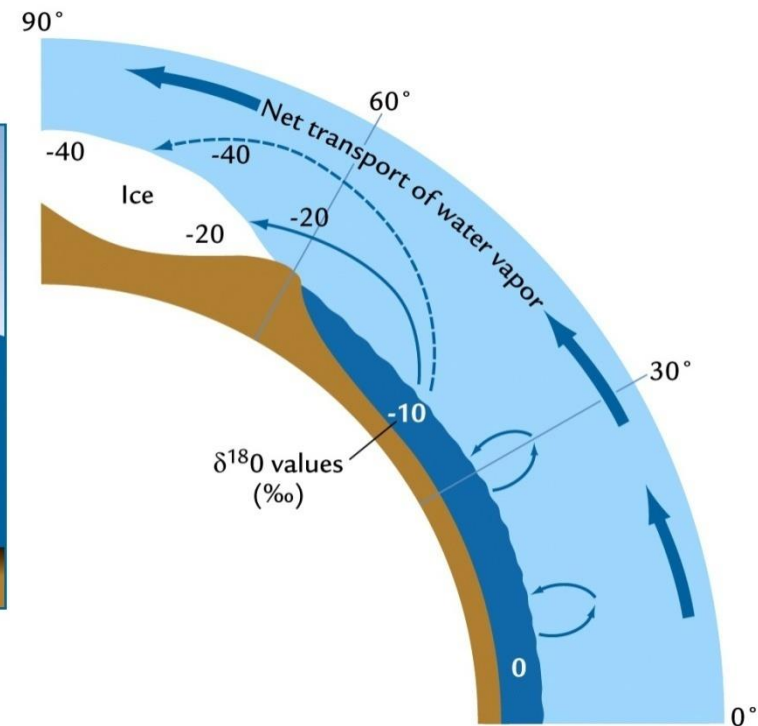
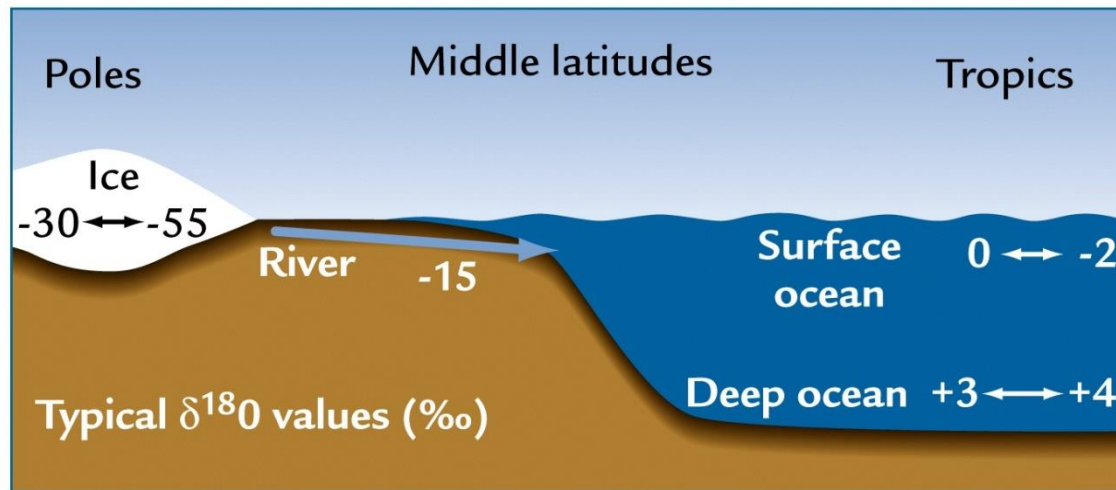
$$R_{N/Gg} = \%Neogloboquadrina / (\%Neogloboquadrina + \%G. glutinata)$$



A selection of paleoclimate proxies

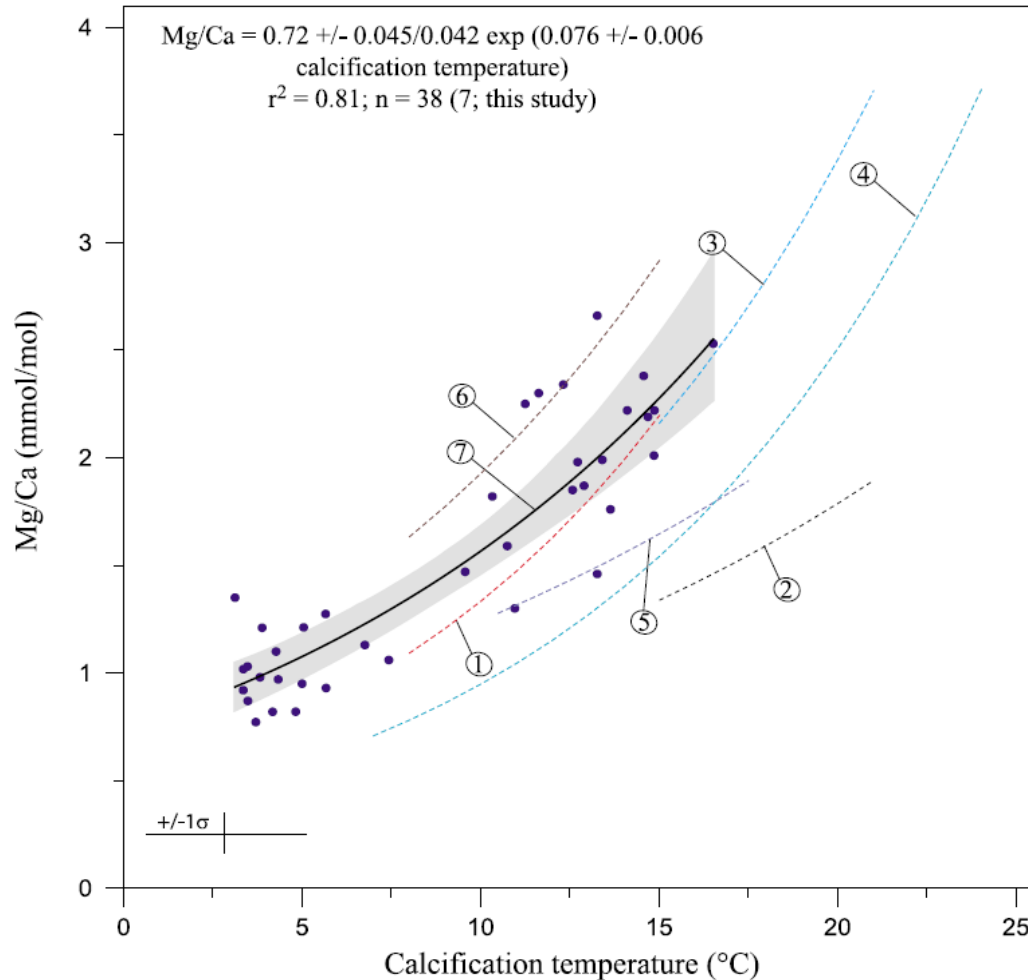


- Oxygen stable isotopes in marine carbonates (e.g., foraminifera) but also ice cores and continental carbonates (e.g., stalagmites)



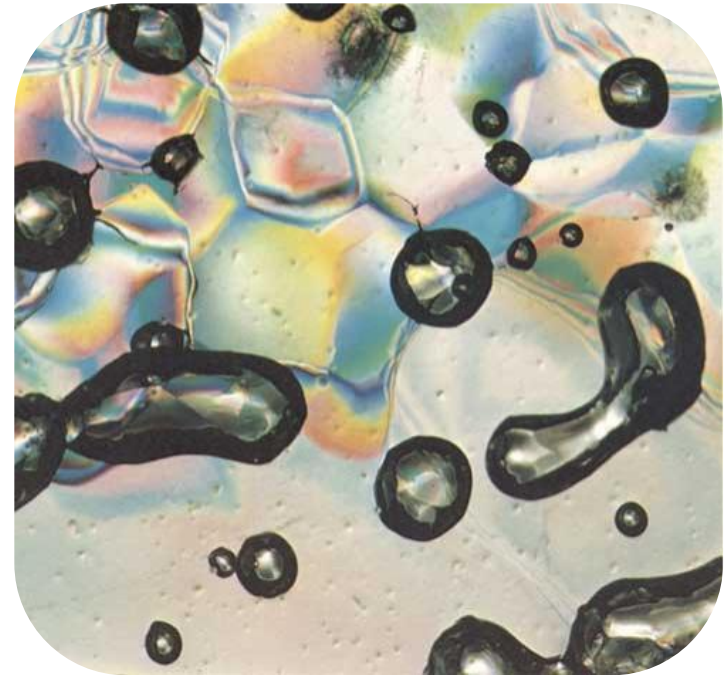
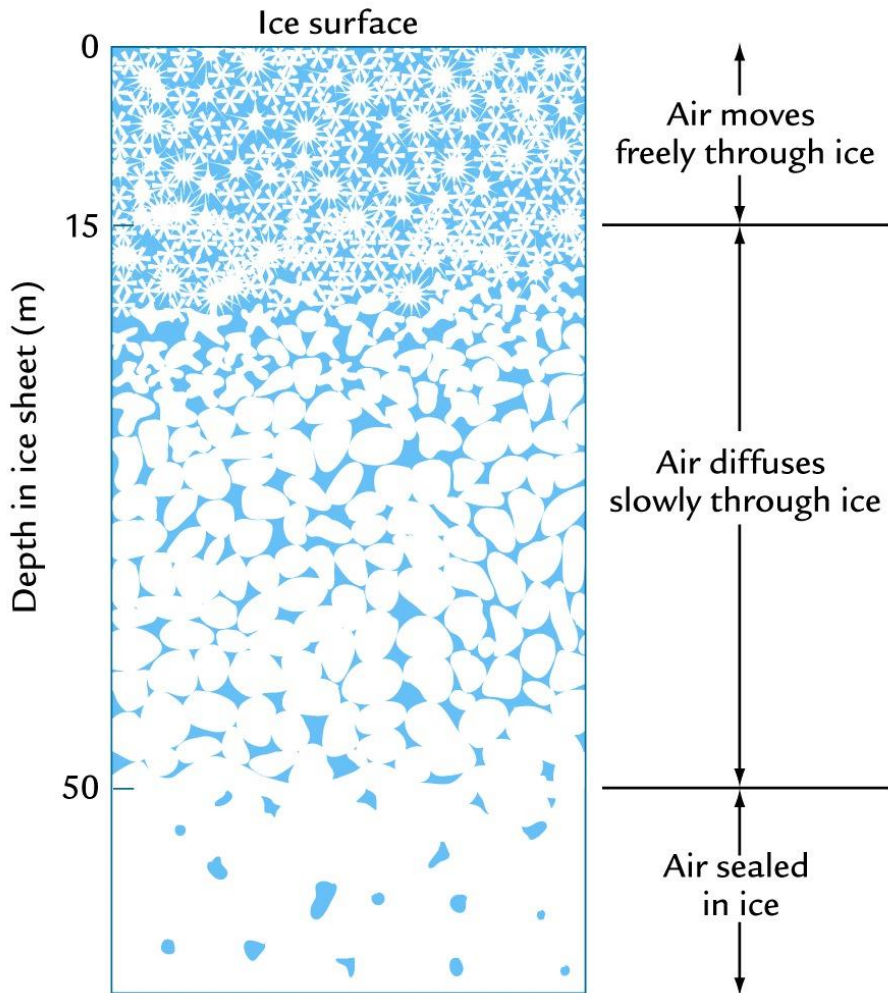


•Mg/Ca in planktonic foraminifera





• Air trapped in ice

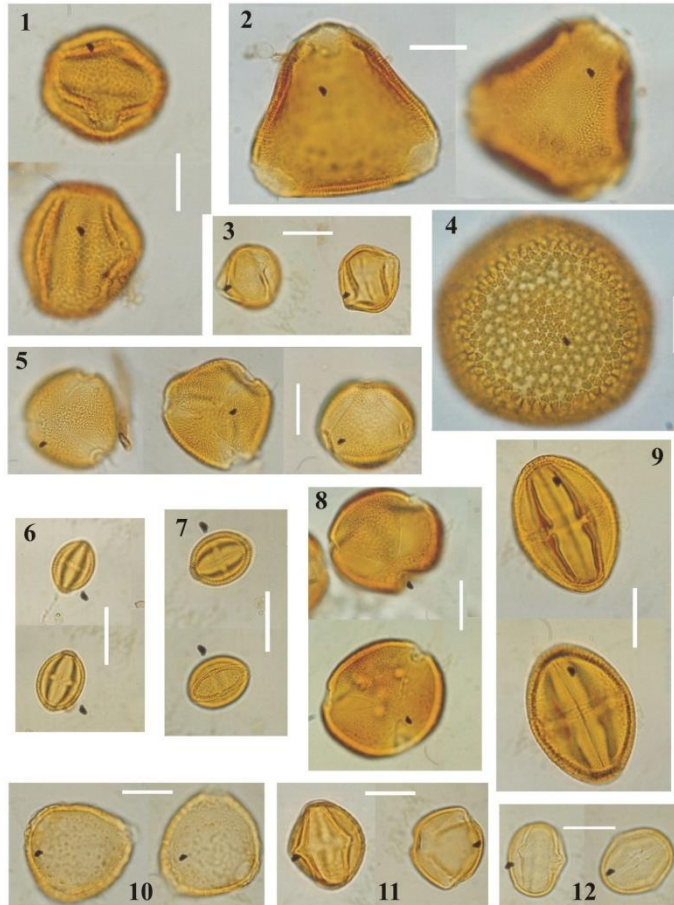




A selection of paleoclimate proxies



•Pollen in continental and marine sediments





A good paleoclimate archive should meet certain criteria

2. It must be possible to determine the age of deposition/formation of the paleoclimate archive

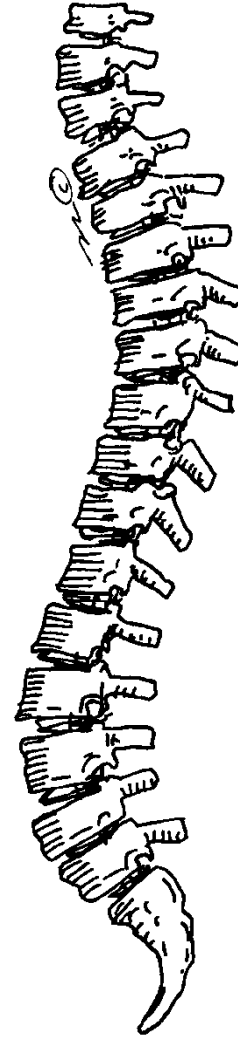




A selection of paleoclimate proxies



*Dating is the
backbone of
paleoclimatology and
paleoceanography*

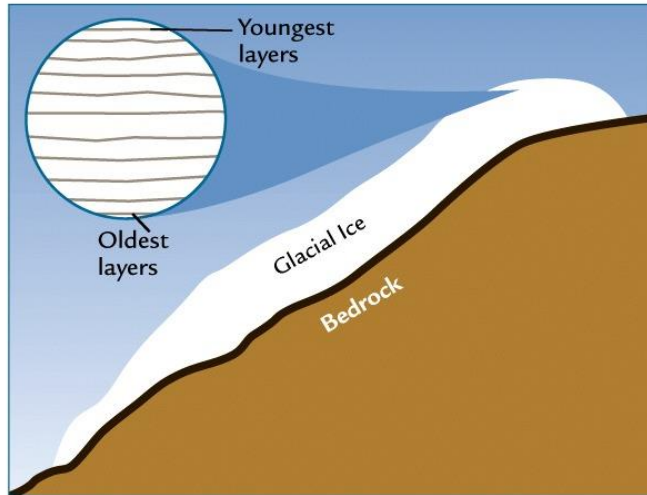




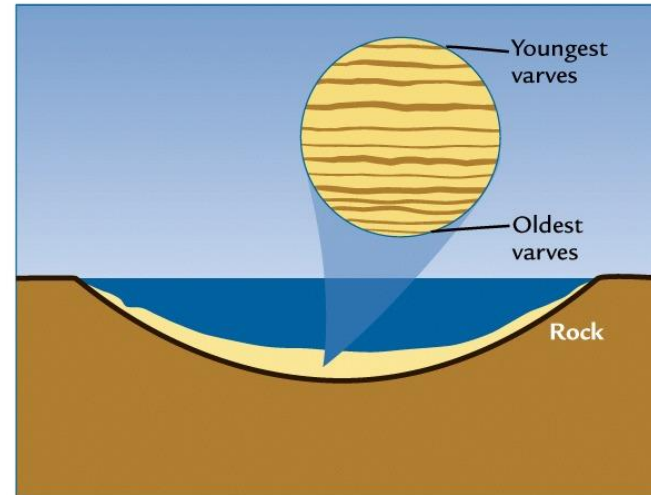
A selection of paleoclimate proxies



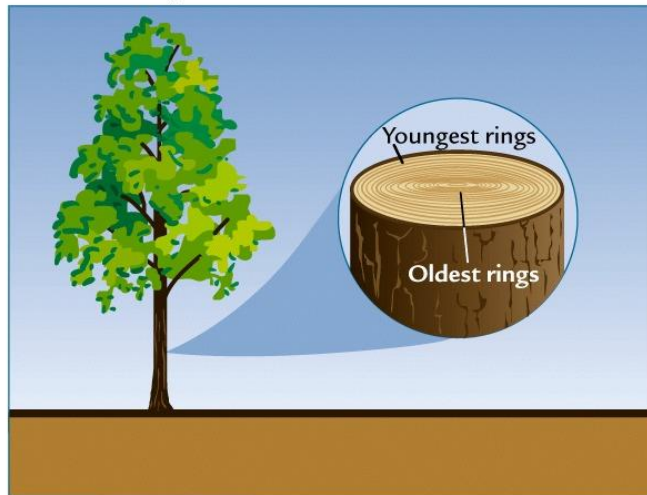
1. Layer counting



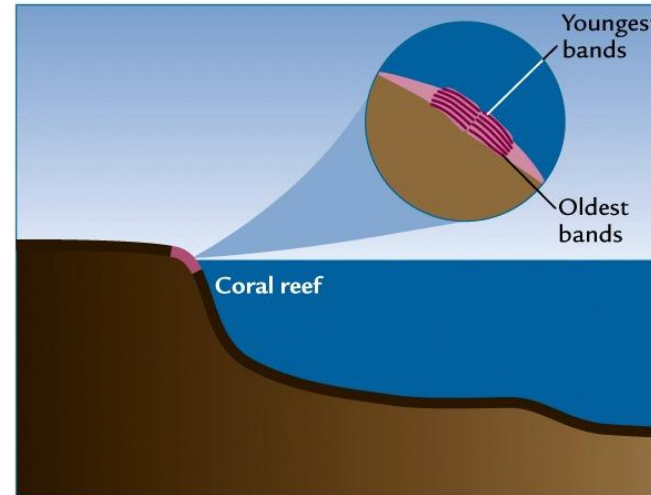
A Annual ice layers



B Annual sediment varves



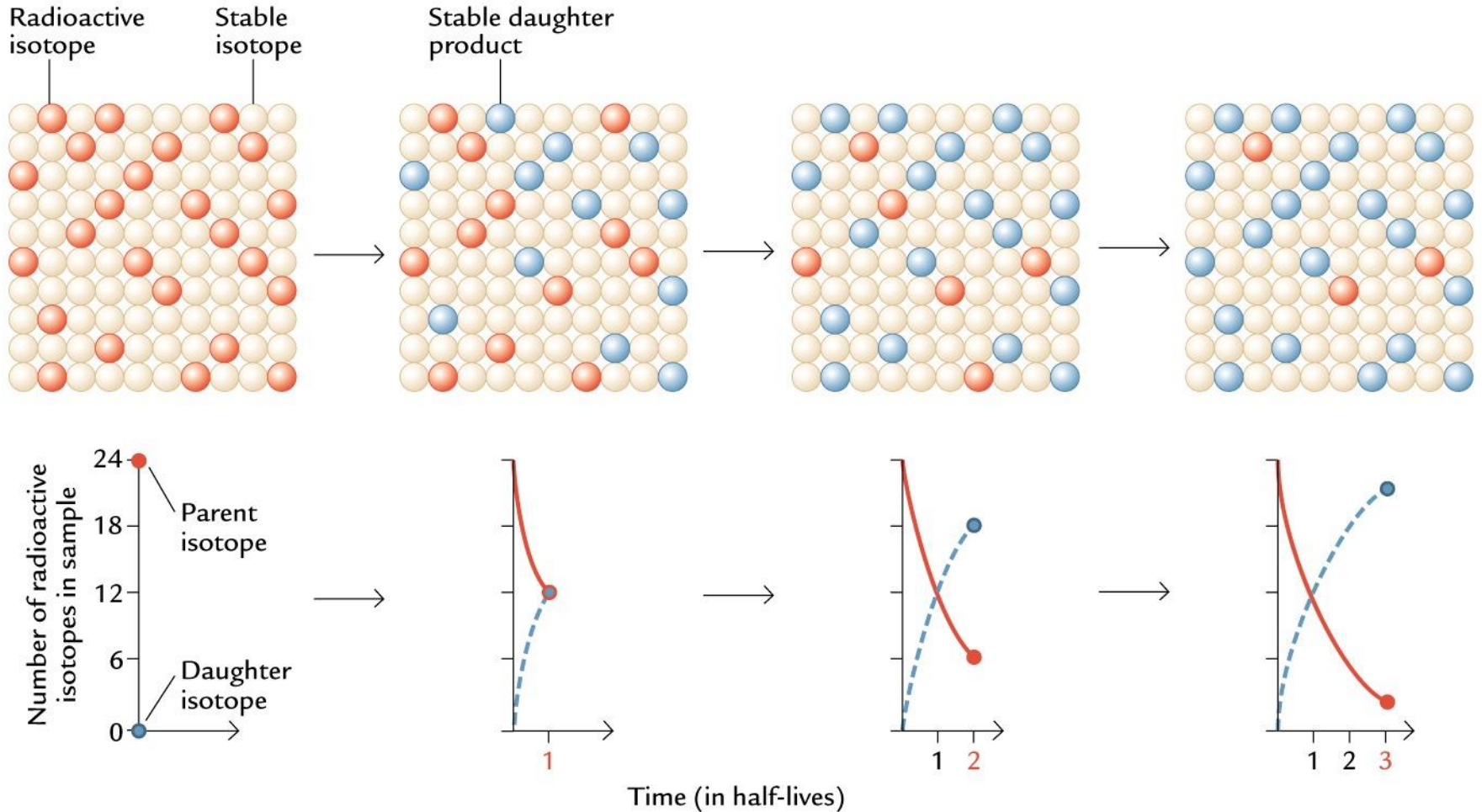
C Annual tree rings



D Annual coral bands



2. Radiometric dating





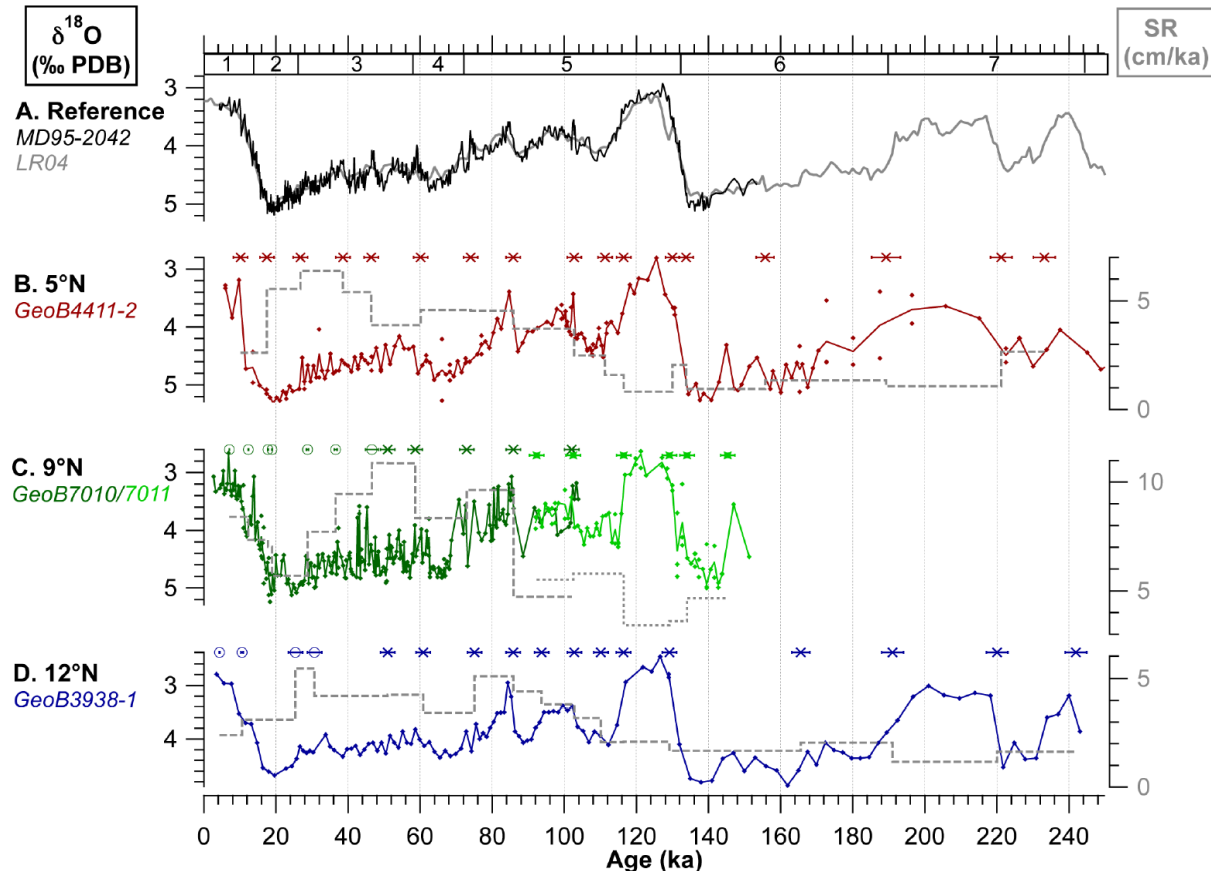
2. Radiometric dating

Given the concentration of the daughter and parent isotope as well as the half-life it is possible to calculate the age

| Parent isotope | Daughter isotope | Half-life | Useful for ages |
|-------------------|-------------------|-----------|-----------------|
| ^{230}Th | ^{226}Ra | 75 kyr | < 600 kyr |
| ^{14}C | ^{14}N | 5.8 kyr | < 50 kyr |



3. Stratigraphic tuning



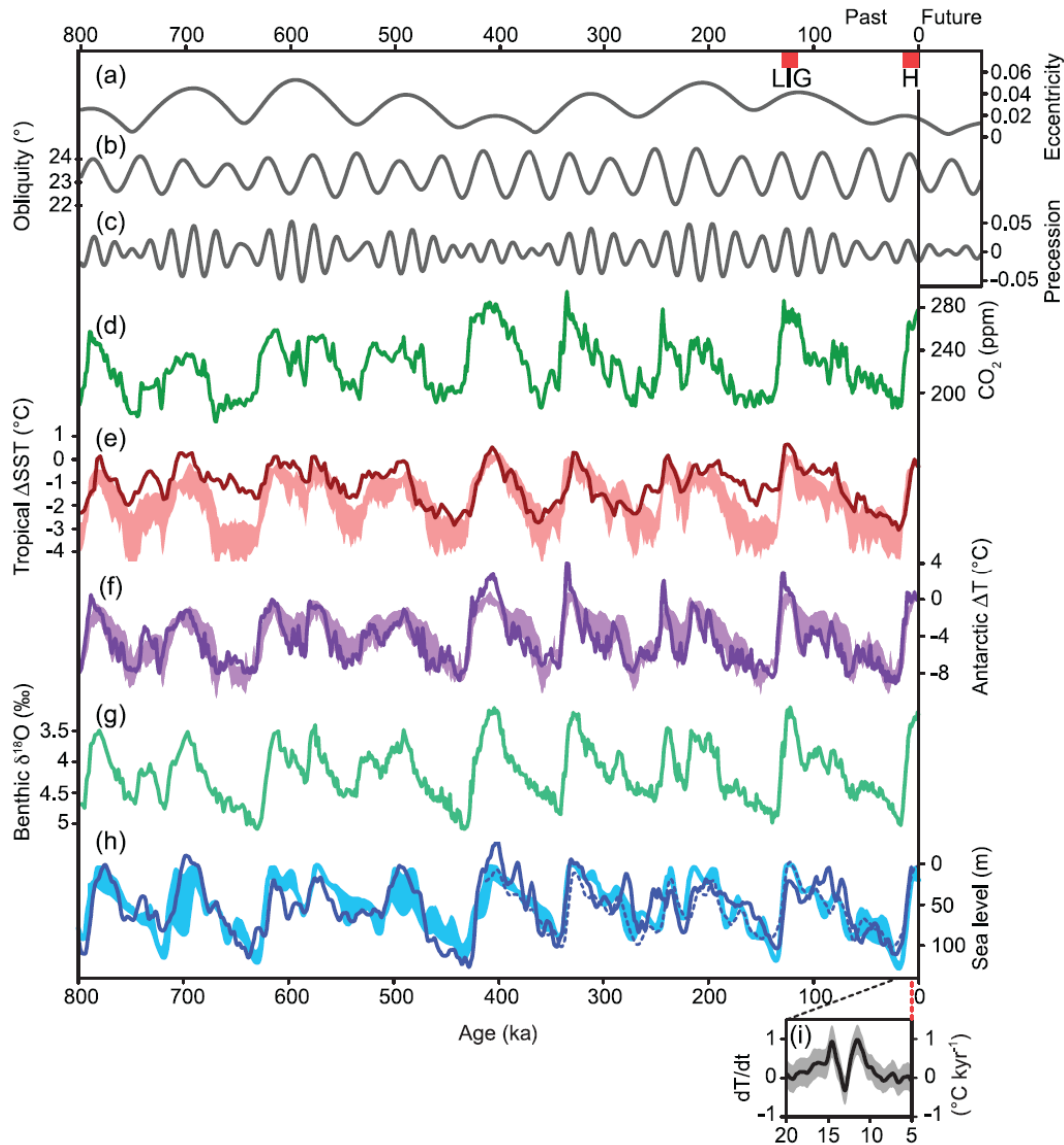
The age model from a different core is imported to the core without an age model by correlating a specific record of both cores that are mechanically related



- Why study past climates?
- Main paleoclimate archives
- A selection of paleoclimate proxies
- Paleoclimate records: *the editor's cut***



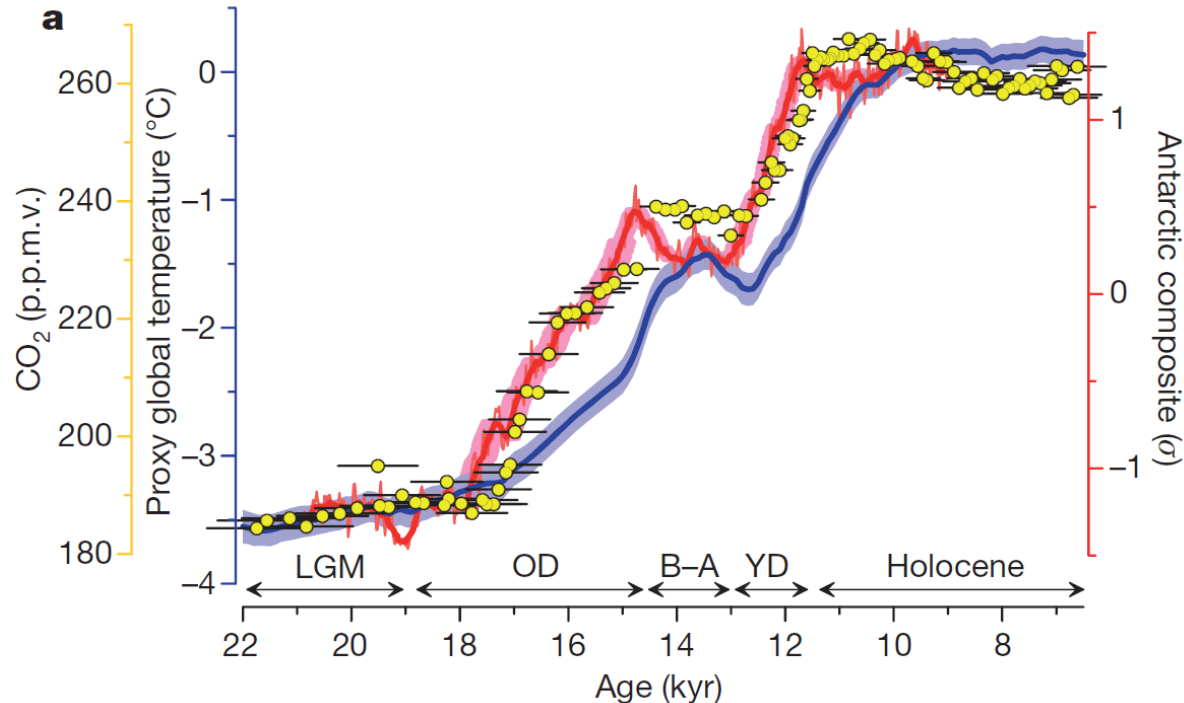
Greenhouse-gas variations and past climate responses



It is a fact that present-day (2011) concentrations of the atmospheric greenhouse gases CO₂, CH₄ and N₂O exceed the range of concentrations recorded in ice cores during the past 800,000 years



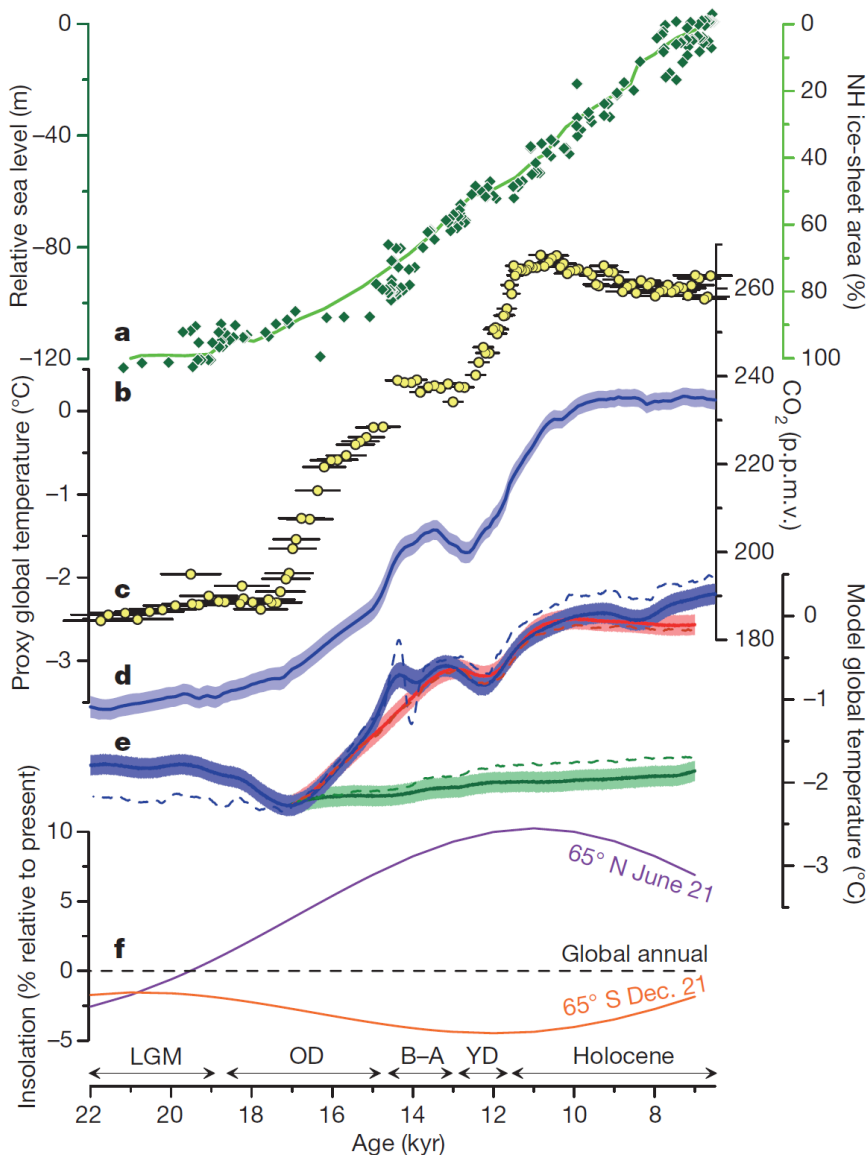
Greenhouse-gas variations and past climate responses



There is high confidence that changes in atmospheric CO₂ concentration play an important role in glacial–interglacial cycles



Greenhouse-gas variations and past climate responses



During the last deglaciation, it is very likely that global mean temperature increased by 3 to 8°C; while the mean rate of global warming was very likely 0.3 to 0.8°C per thousand years, two periods were marked by faster warming rates (likely 1-1.5°C per thousand years), although regionally and on shorter time scales higher rates may have occurred

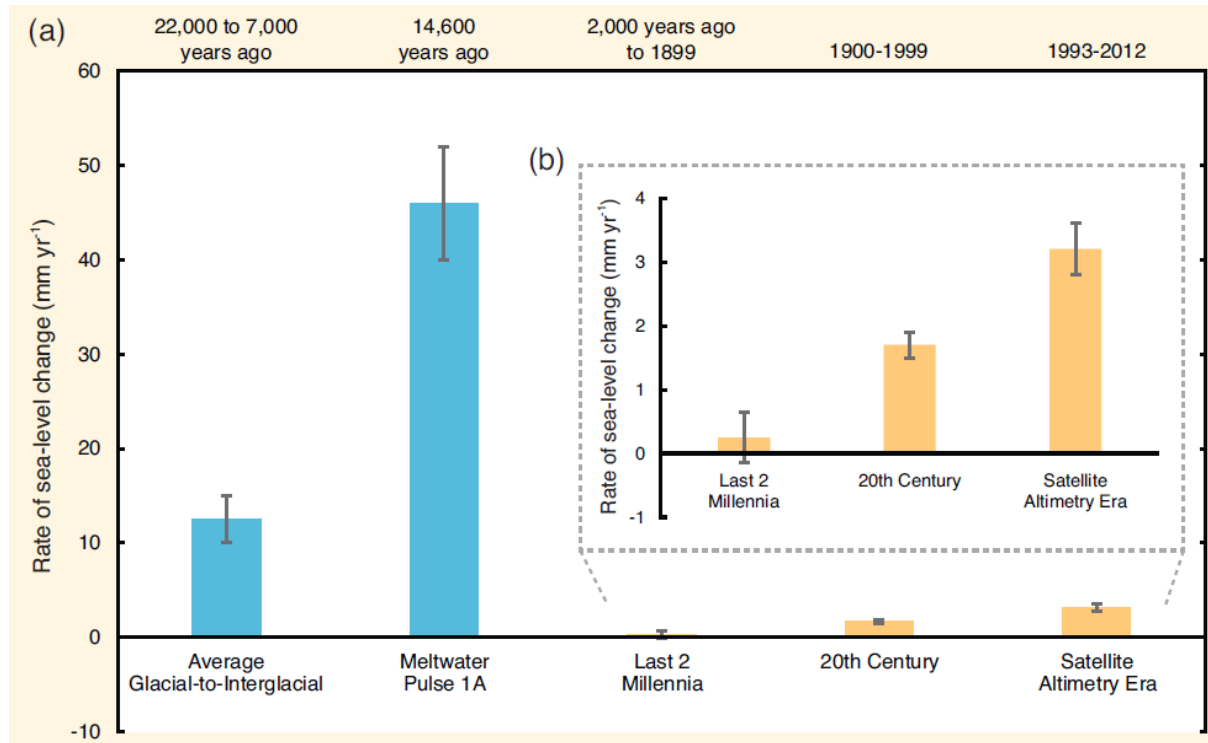


Greenhouse-gas variations and past climate responses

With medium confidence, global mean surface temperature was significantly above pre-industrial levels during several past periods characterized by high atmospheric CO₂ concentrations



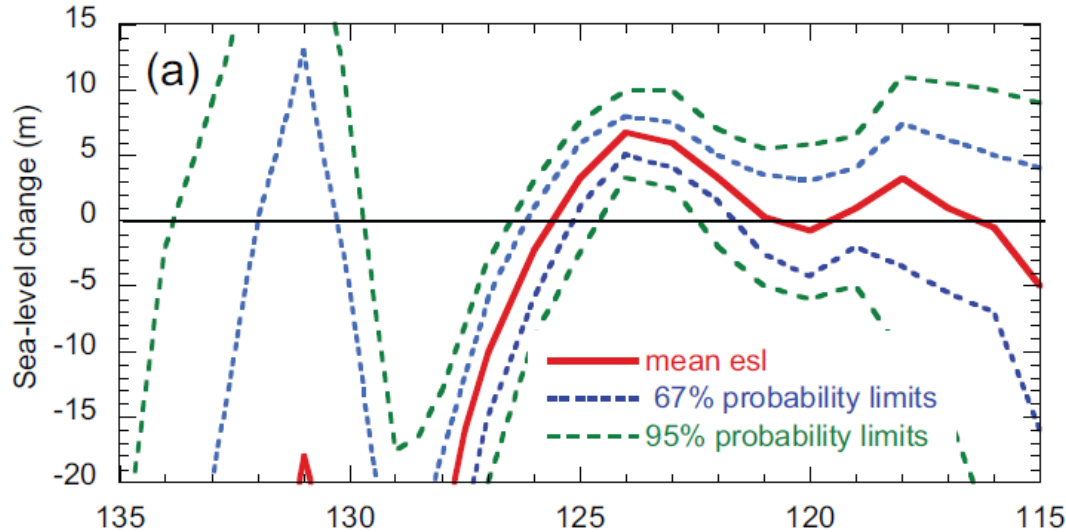
Global sea level changes during past warm periods



The current rate of global mean sea level change, starting in the late 19th-early 20th century, is, with medium confidence, unusually high in the context of centennial-scale variations of the last two millennia



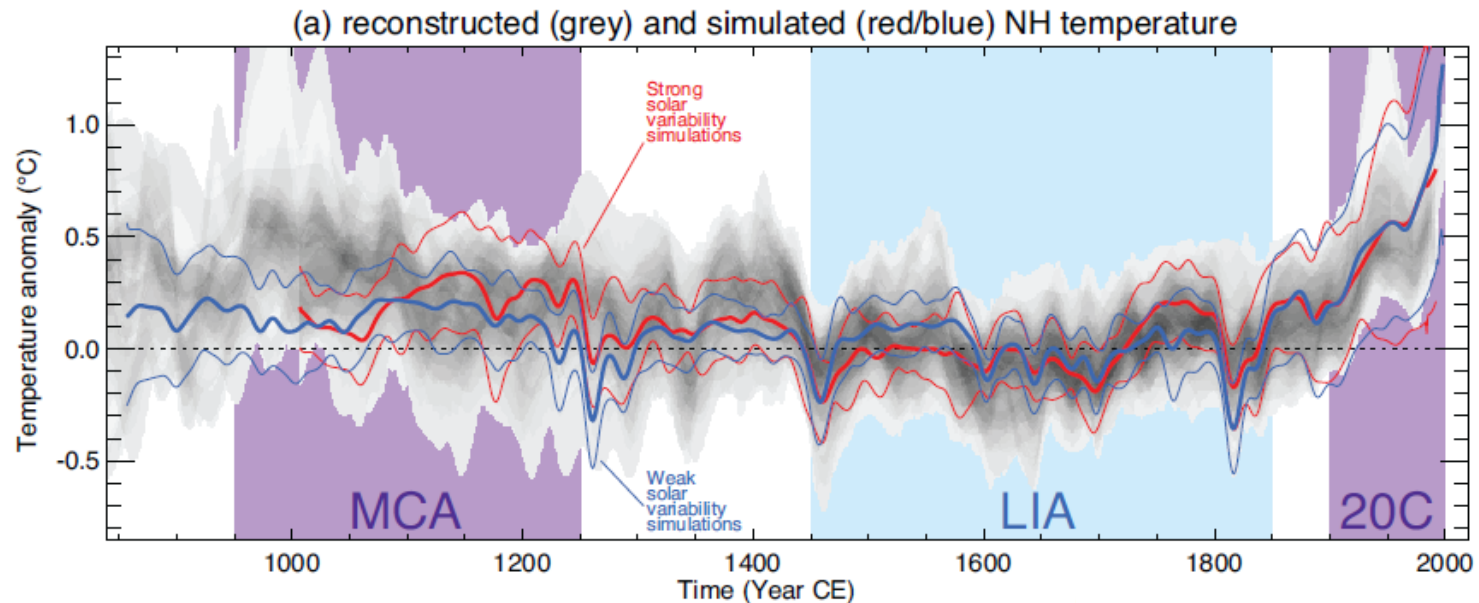
Global sea level changes during past warm periods



There is very high confidence that the maximum global mean sea level during the last interglacial period (129,000 to 116,000 years ago) was, for several thousand years, at least 5 m higher than present and high confidence that it did not exceed 10 m above present (contribution from Greenland 1.4-4.3 m and Antarctica)



Observed recent climate change in the context of interglacial climate variability



There is high confidence that annual mean surface warming since the 20th century has reversed long-term cooling trends of the past 5000 years in mid-to-high latitudes of the Northern Hemisphere



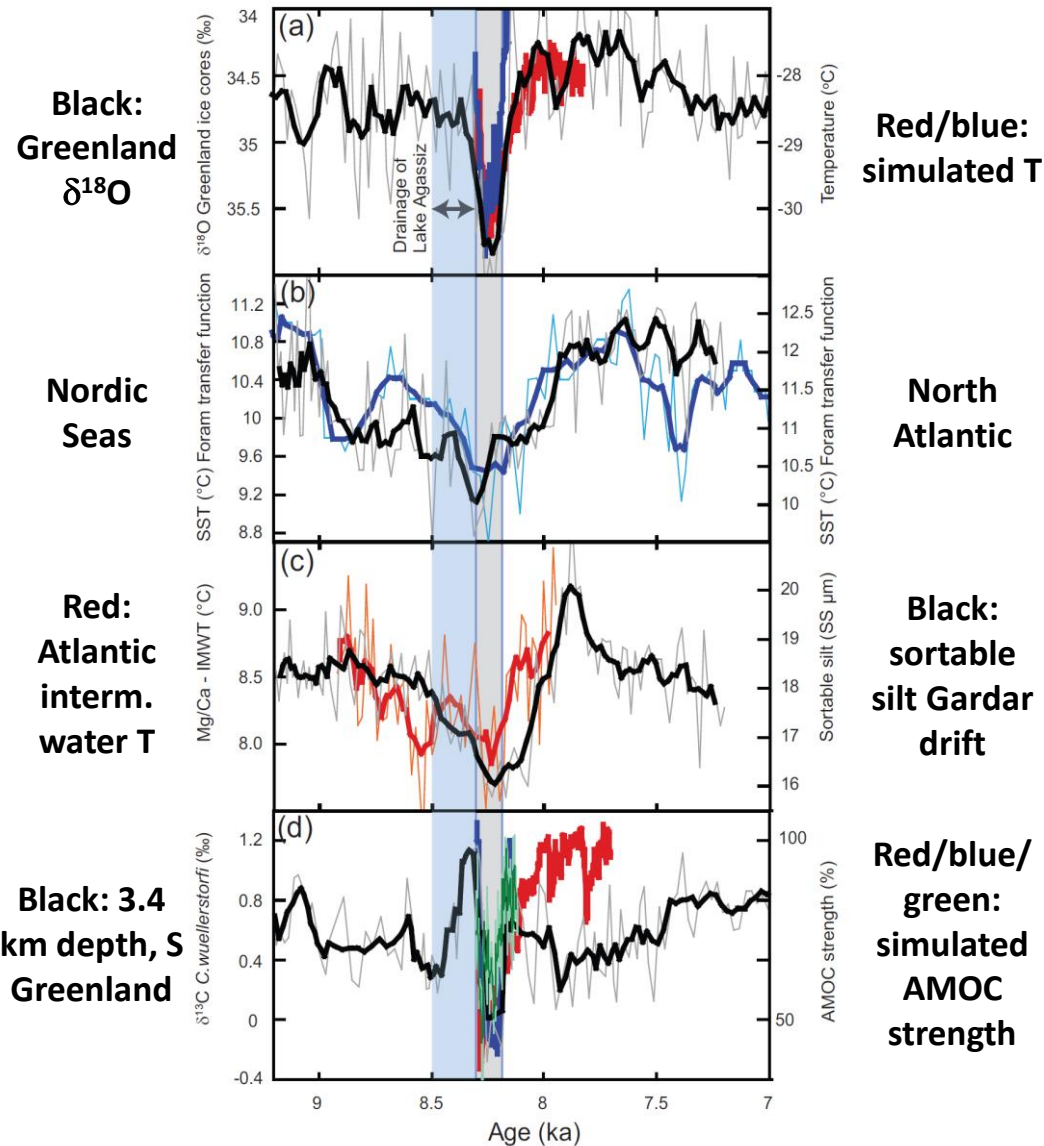
Observed recent climate change in the context of interglacial climate variability

There is high confidence for droughts during the last millennium of greater magnitude and longer duration than those observed since the beginning of the 20th century in many regions

With high confidence, floods larger than those recorded since 1900 occurred during the past five centuries in northern and central Europe, western Mediterranean region and eastern Asia



Abrupt climate change and irreversibility



With high confidence, the interglacial mode of the Atlantic meridional overturning circulation can recover from a short-term freshwater input into the subpolar North Atlantic



Observed recent climate change in the context of interglacial climate variability

It is virtually certain that orbital forcing will be unable to trigger widespread glaciation during the next 1000 years

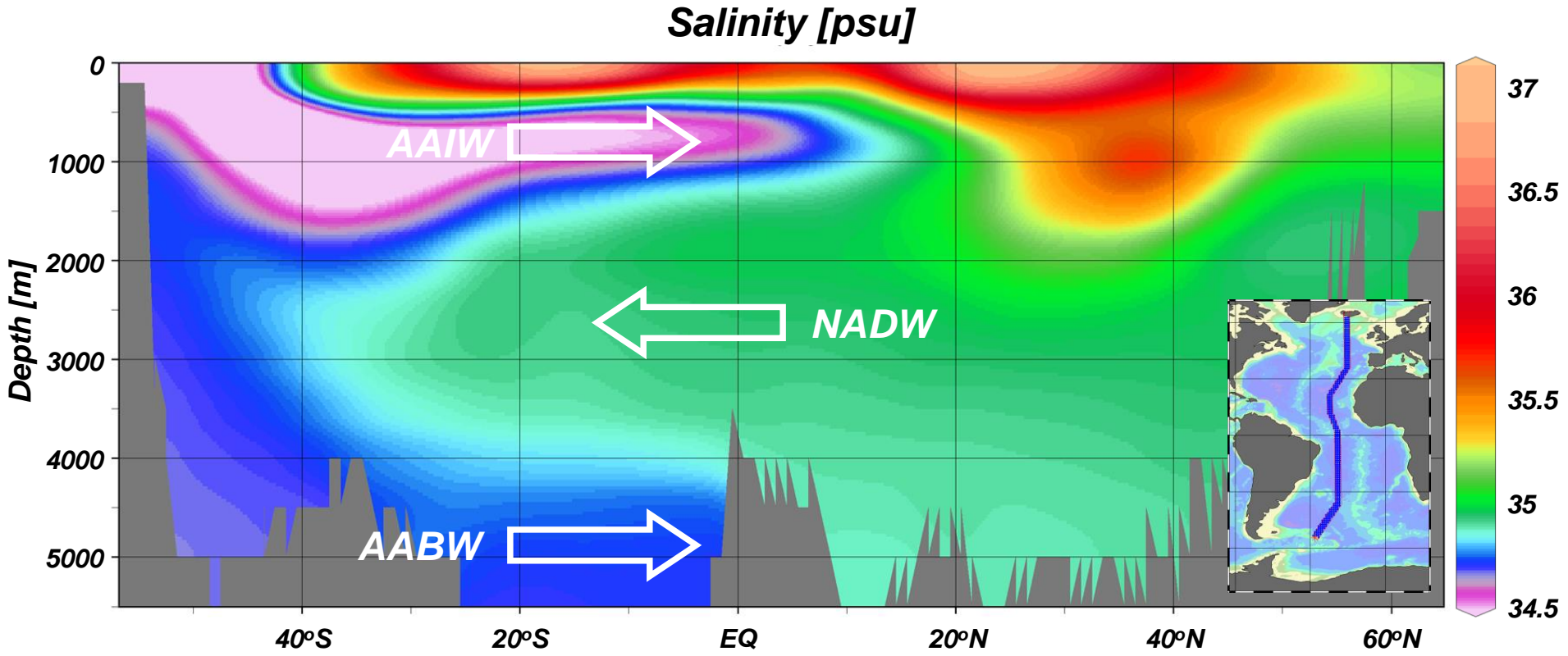
Paleoclimate records indicate that, for orbital configurations close to the present one, glacial inceptions only occurred for atmospheric CO₂ concentrations significantly lower than pre-industrial levels

Climate models simulate no glacial inception during the next 50,000 years if CO₂ concentrations remain above 300 ppm.



Abrupt climate change and irreversibility

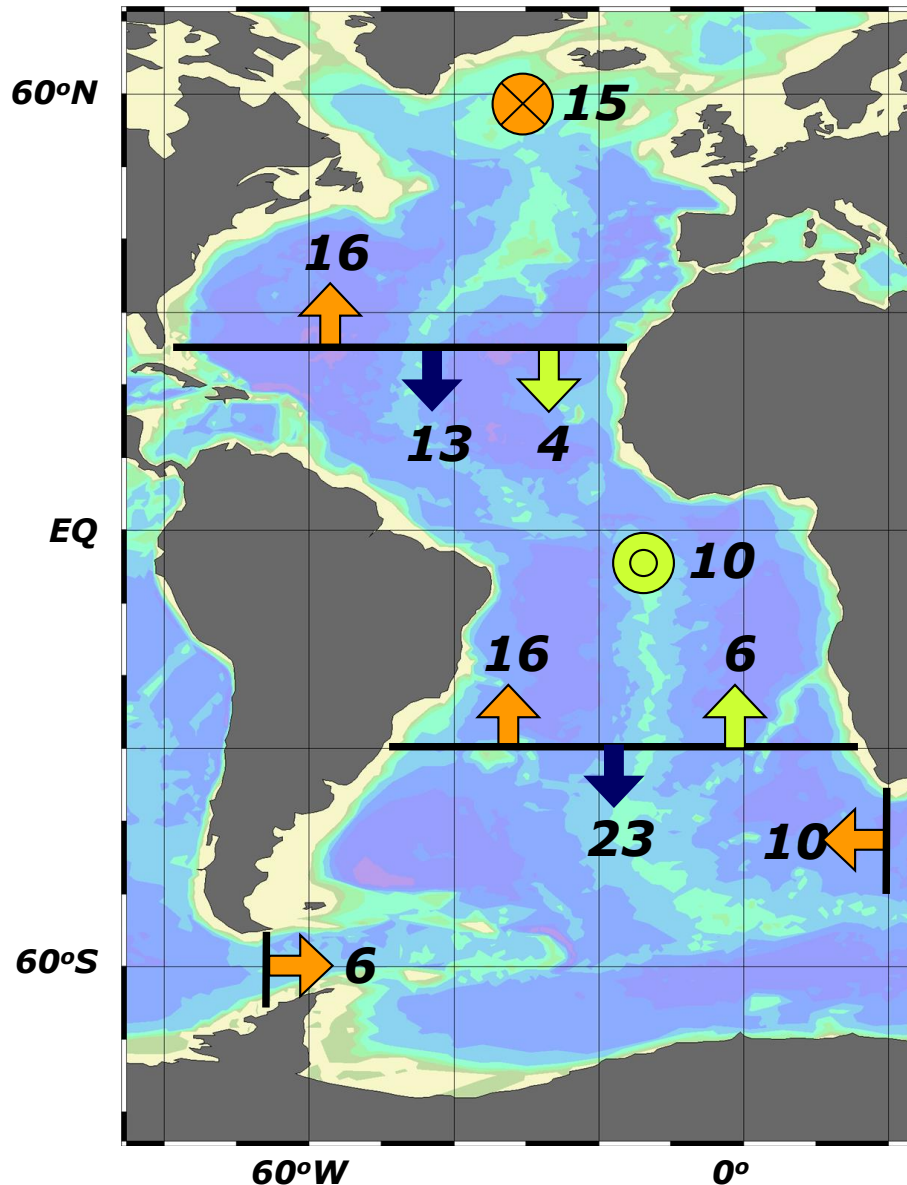
Confidence in the link between changes in North Atlantic climate and low-latitude precipitation patterns has increased since AR4








Atlantic meridional overturning circulation (AMOC)



Paleoclimate records: *the editor's cut*



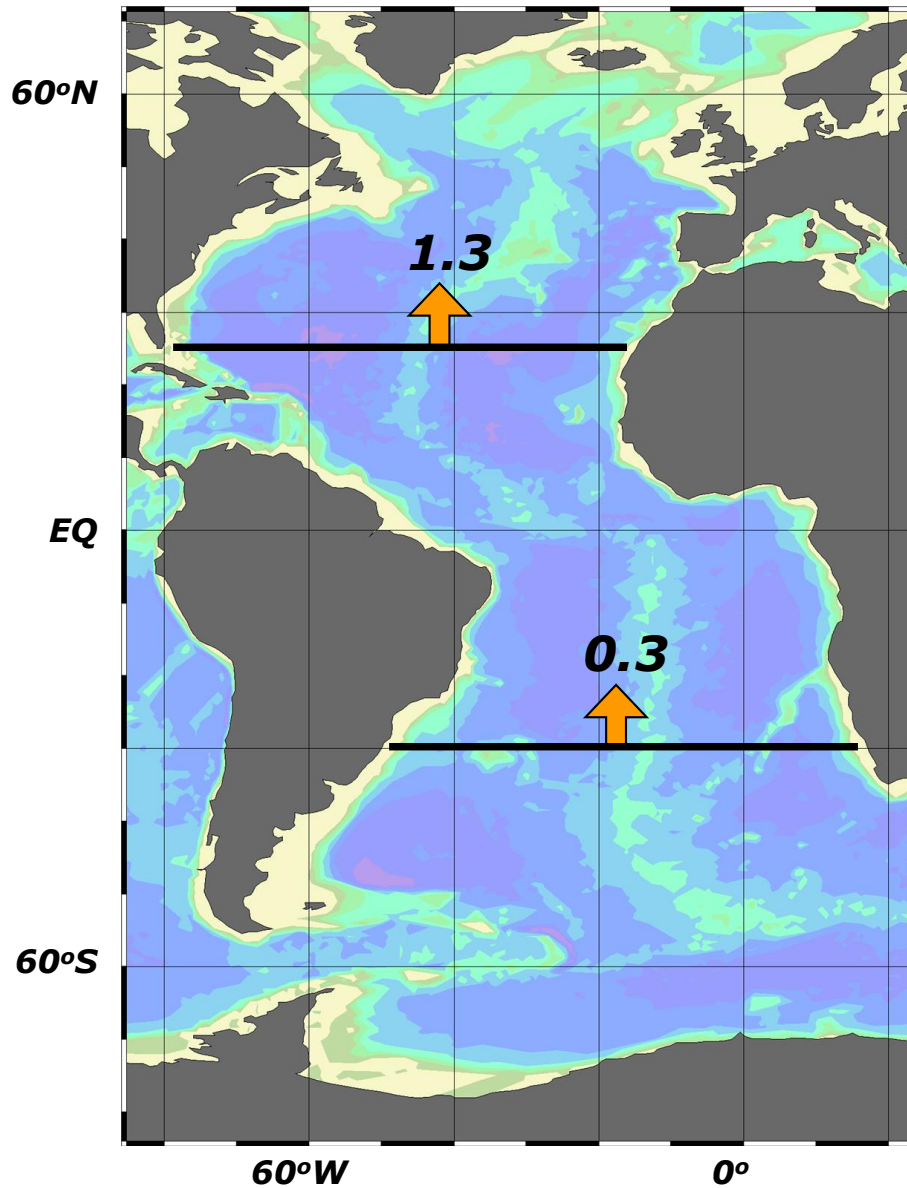
$$1Sv = 10^6 m^3 s^{-1}$$

-  *Shallow*
-  *Deep*
-  *Bottom*
-  *Upwelling*
-  *Downwelling*

The inflow of Indian Ocean high salinity waters preconditions the Atlantic for NADW formation



Paleoclimate records: *the editor's cut*



$1PW = 10^{15} W$

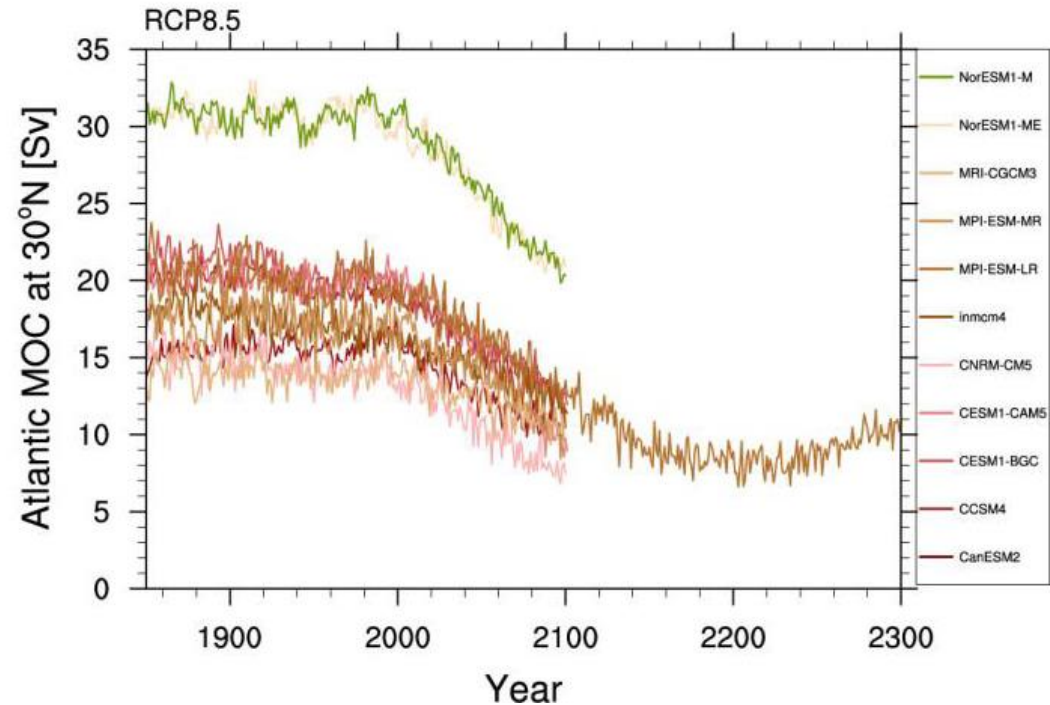
*Itaipu generates
0,000014 PW
(14 GW)*

*Large influence
over global
climate*



IPCC AR5 numerical models suggest a decrease in strength of the AMOC

Between 20–30% (RCP4.5) e 36–44% (RCP8.5) in 2100



Changes in AMOC may impact global climate



nature
climate change

LETTERS

PUBLISHED ONLINE: 29 JANUARY 2012 | DOI: 10.1038/NCLIMATE1353

Enhanced warming over the global subtropical western boundary currents

Lixin Wu^{1*}, Wenju Cai², Liping Zhang¹, Hisashi Nakamura³, Axel Timmermann⁴, Terry Joyce⁵, Michael J. McPhaden⁶, Michael Alexander⁷, Bo Qiu⁴, Martin Visbeck⁸, Ping Chang⁹ and Benjamin Giese⁹

SCIENCE

19 JUNE 2015

Observing the Atlantic Meridional Overturning Circulation yields a decade of inevitable surprises

M. A. Srokosz^{1*} and H. L. Bryden²

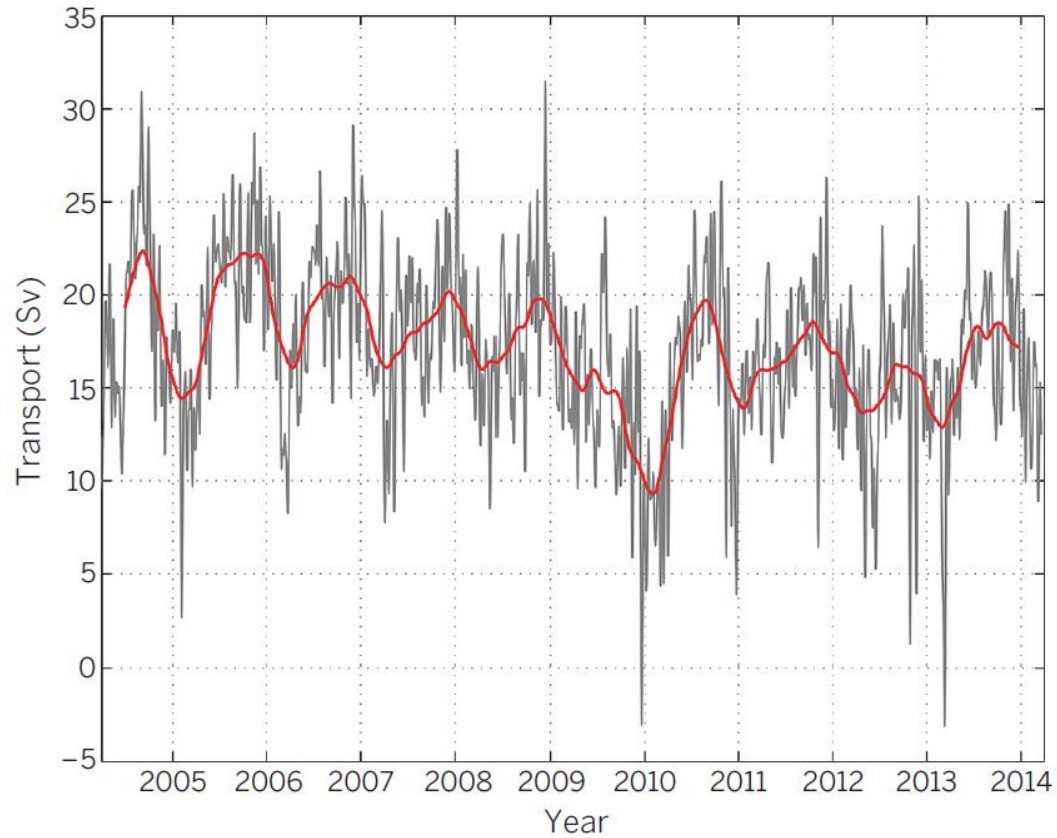
nature
climate change

ARTICLES

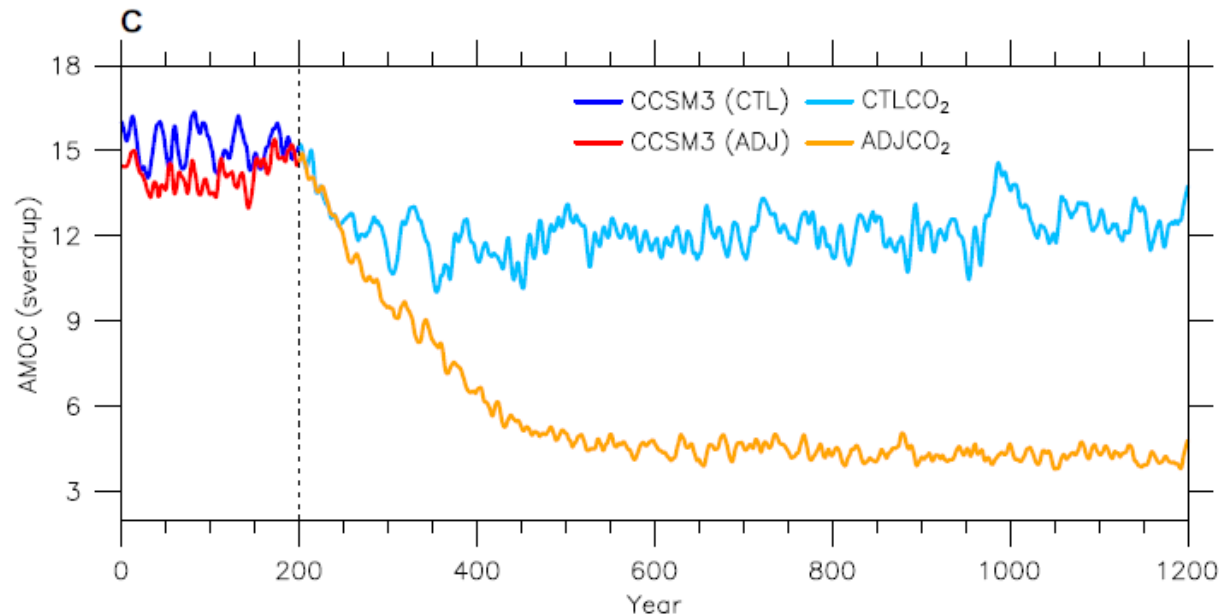
PUBLISHED ONLINE: 23 MARCH 2015 | DOI: 10.1038/NCLIMATE2554

Exceptional twentieth-century slowdown in Atlantic Ocean overturning circulation

Stefan Rahmstorf^{1*}, Jason E. Box², Georg Feulner¹, Michael E. Mann^{3,4}, Alexander Robinson^{1,5,6}, Scott Rutherford⁷ and Erik J. Schaffernicht¹



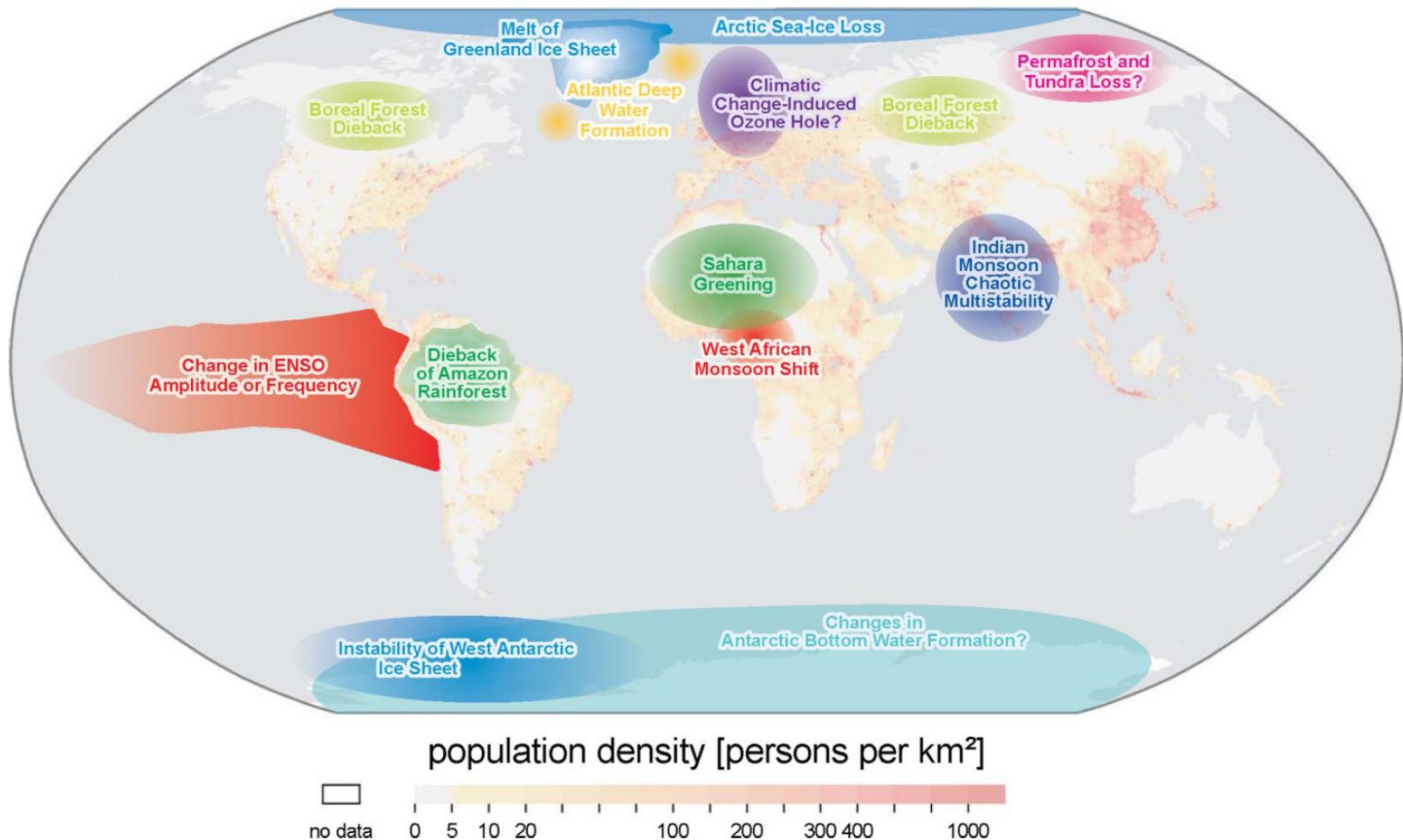
Instrumental records are too short to appropriately investigate changes in AMOC



Overlooked possibility of a collapsed Atlantic meridional overturning circulation in warming climate



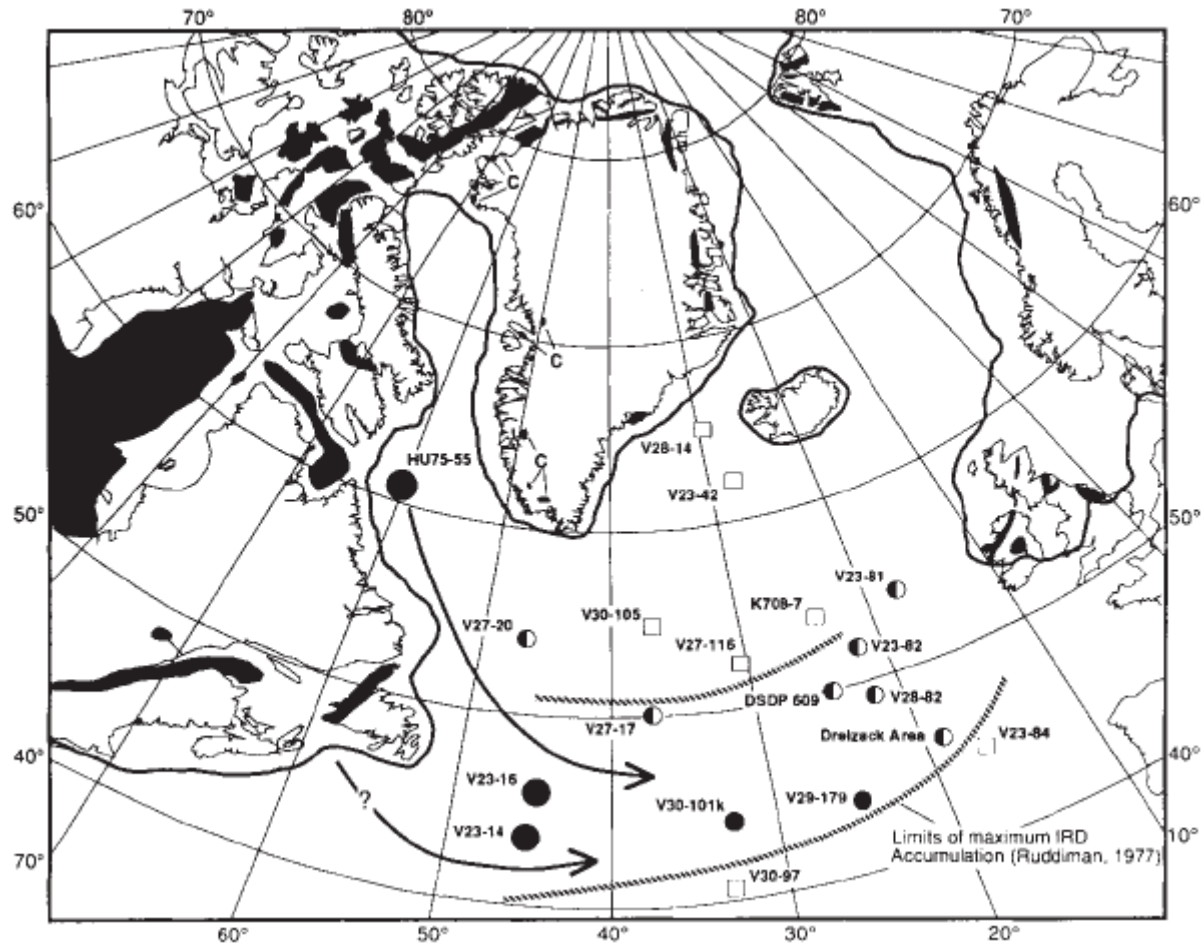
Paleoclimate records: *the editor's cut*



Formation of NADW: a tipping element in the climate system whose threshold could be reached this century

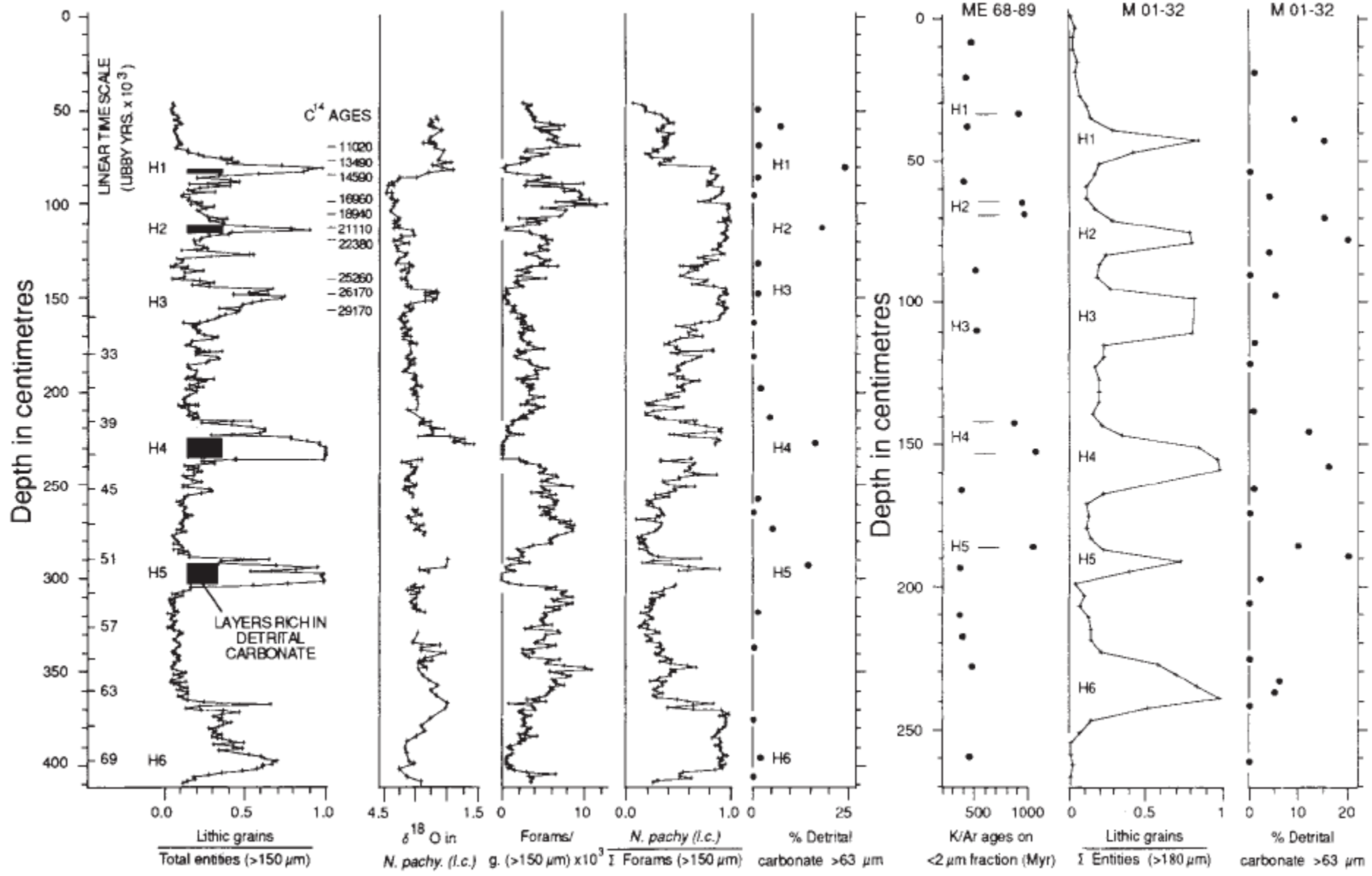


Paleoclimate records: *the editor's cut*



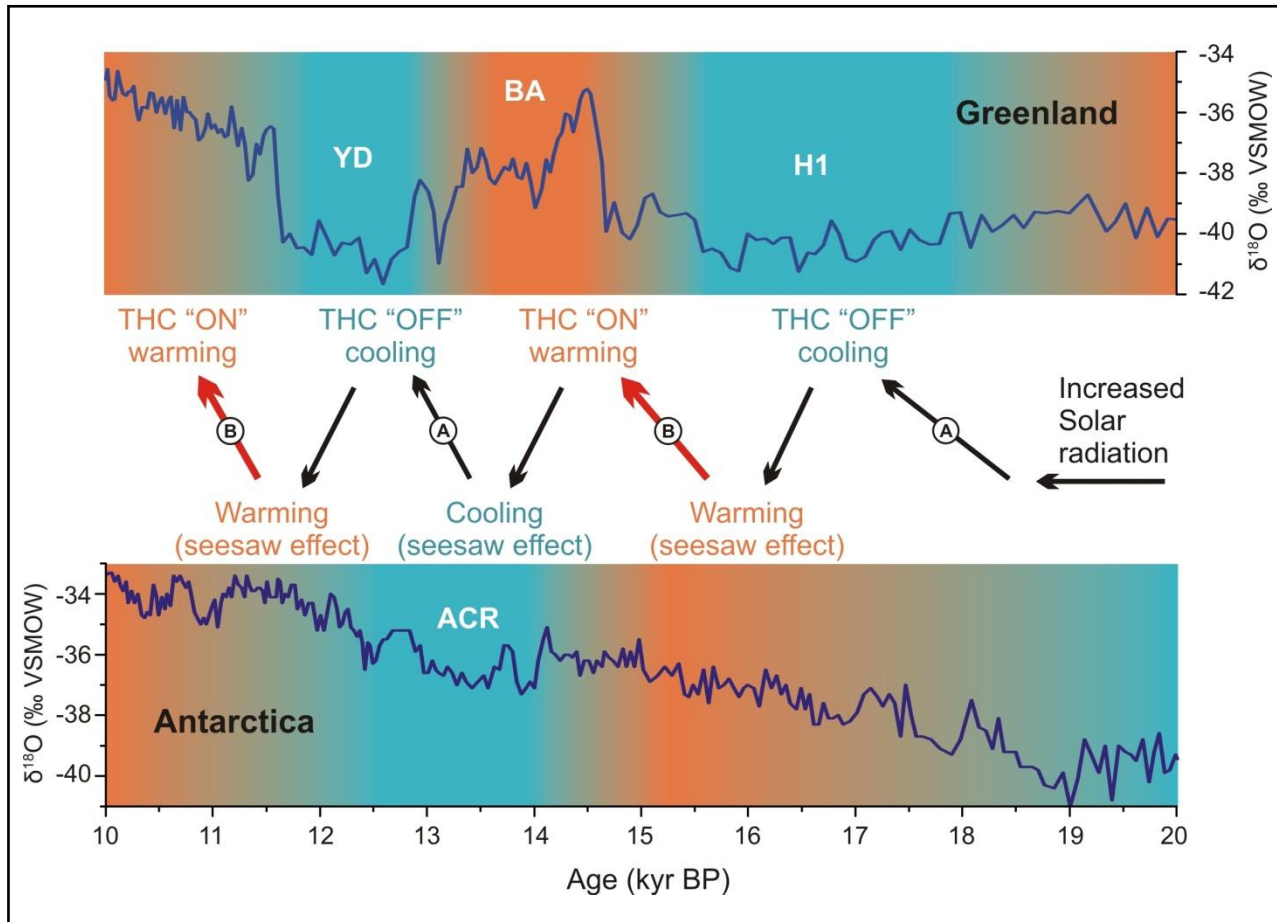


Paleoclimate records: *the editor's cut*



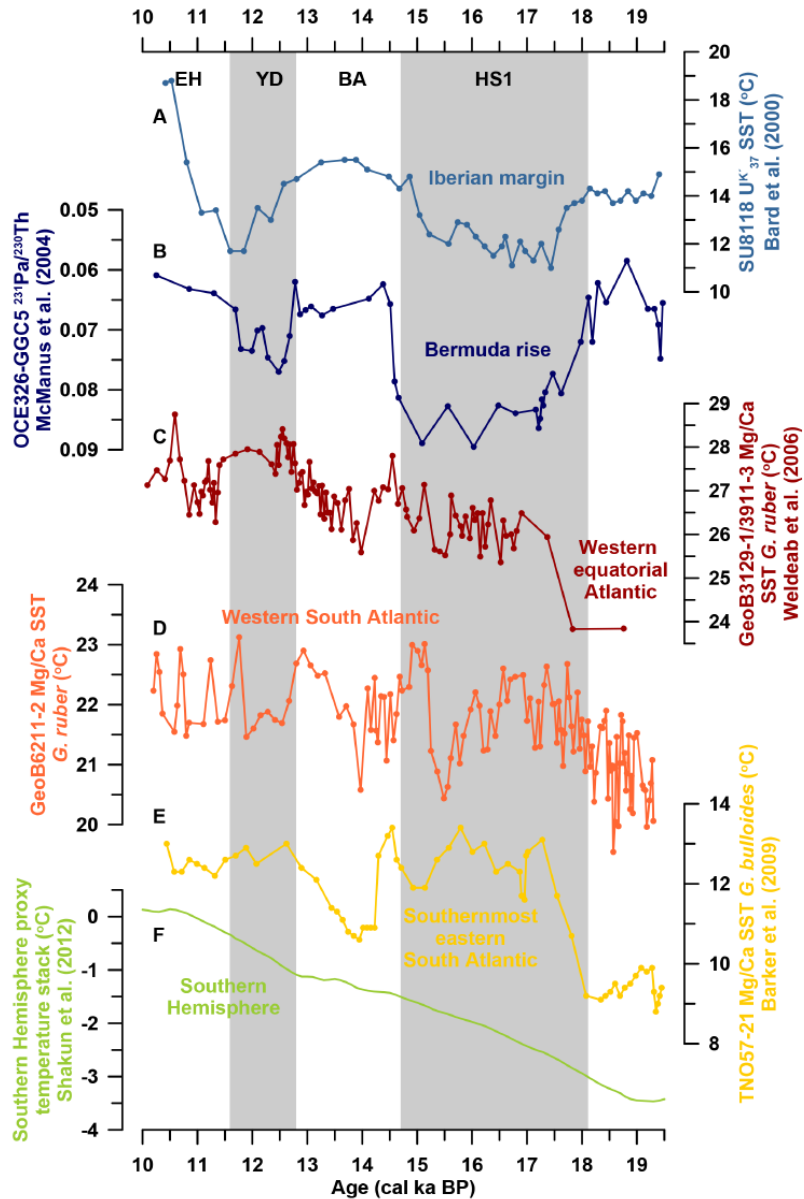


Paleoclimate records: *the editor's cut*





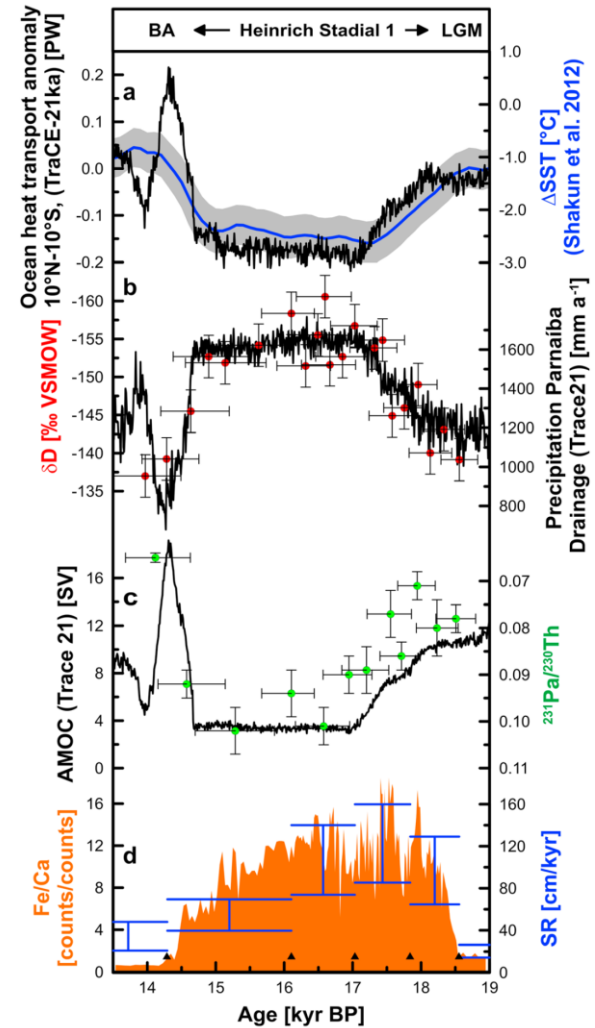
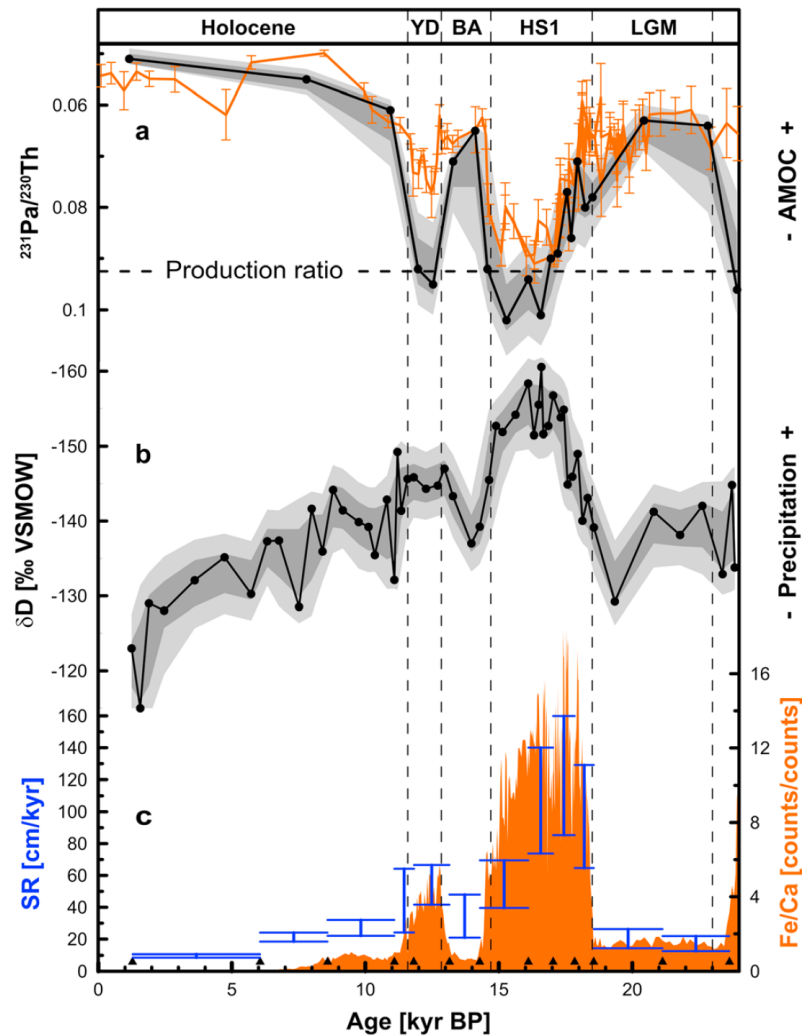
Paleoclimate records: *the editor's cut*



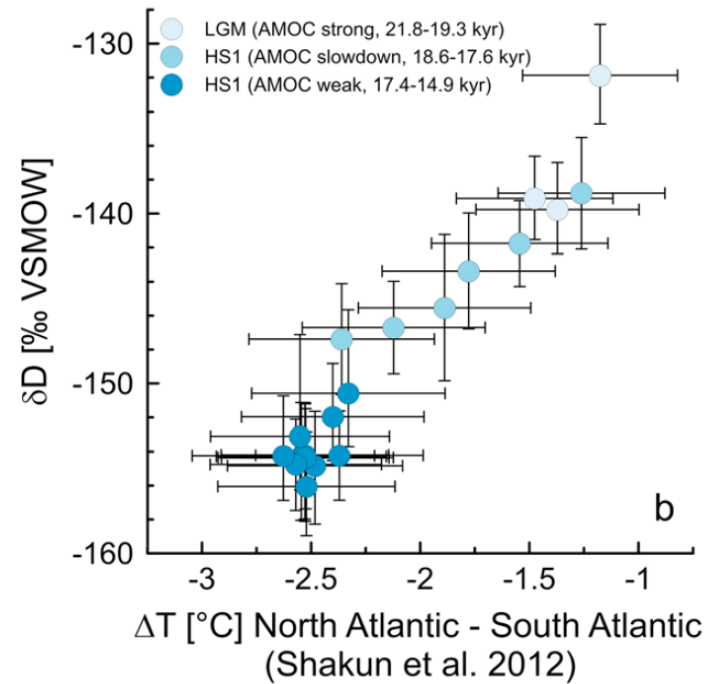
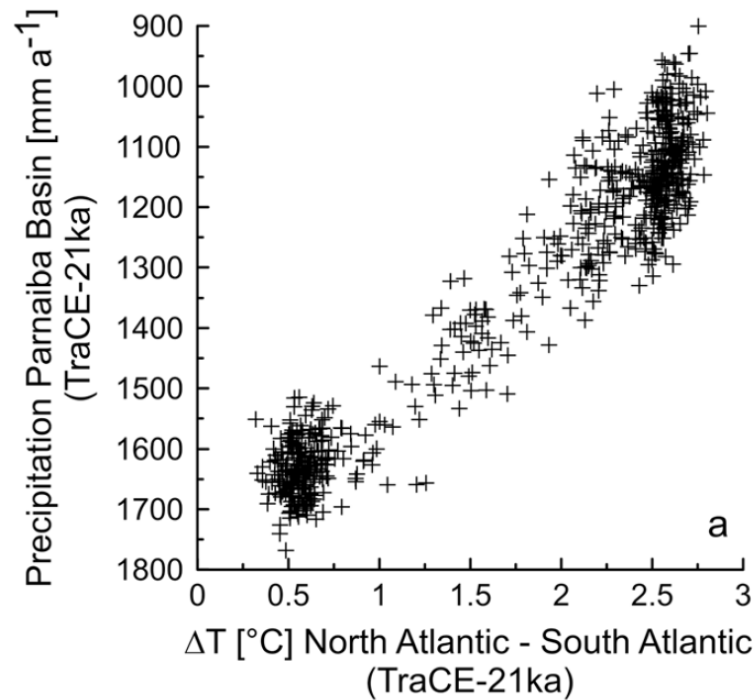
Opposite changes in sea surface temperatures between the North and South Atlantic during AMOC slowdown events



Paleoclimate records: *the editor's cut*



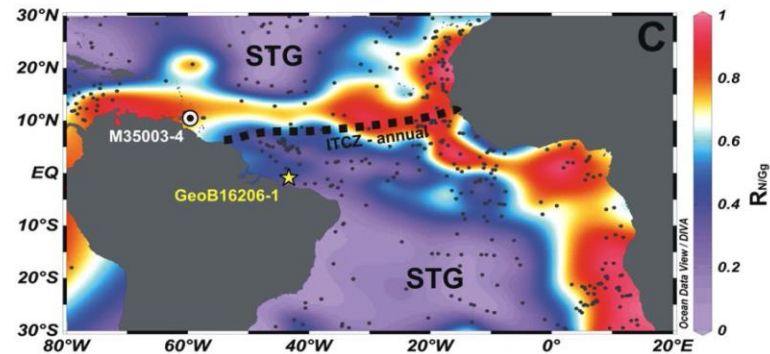
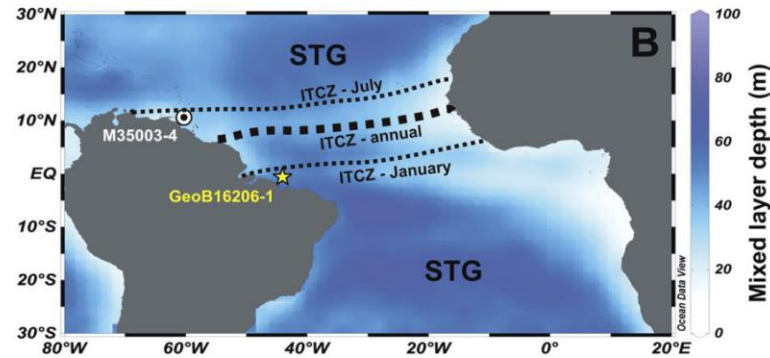
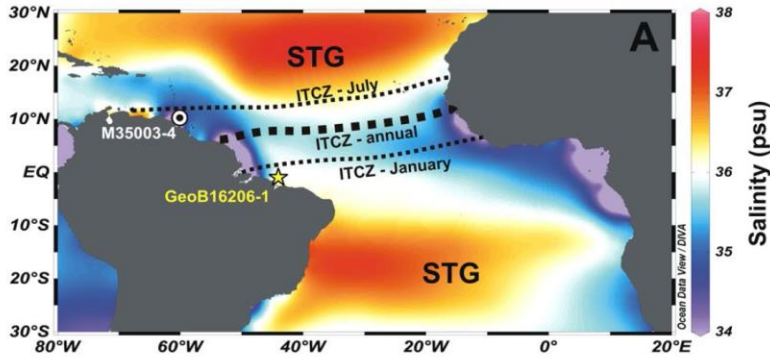
Precipitation over NE Brazil varied linearly and synchronously with AMOC strength



Intertropical Convergence Zone over the western Atlantic closely followed the change in the meridional sea surface temperature gradient



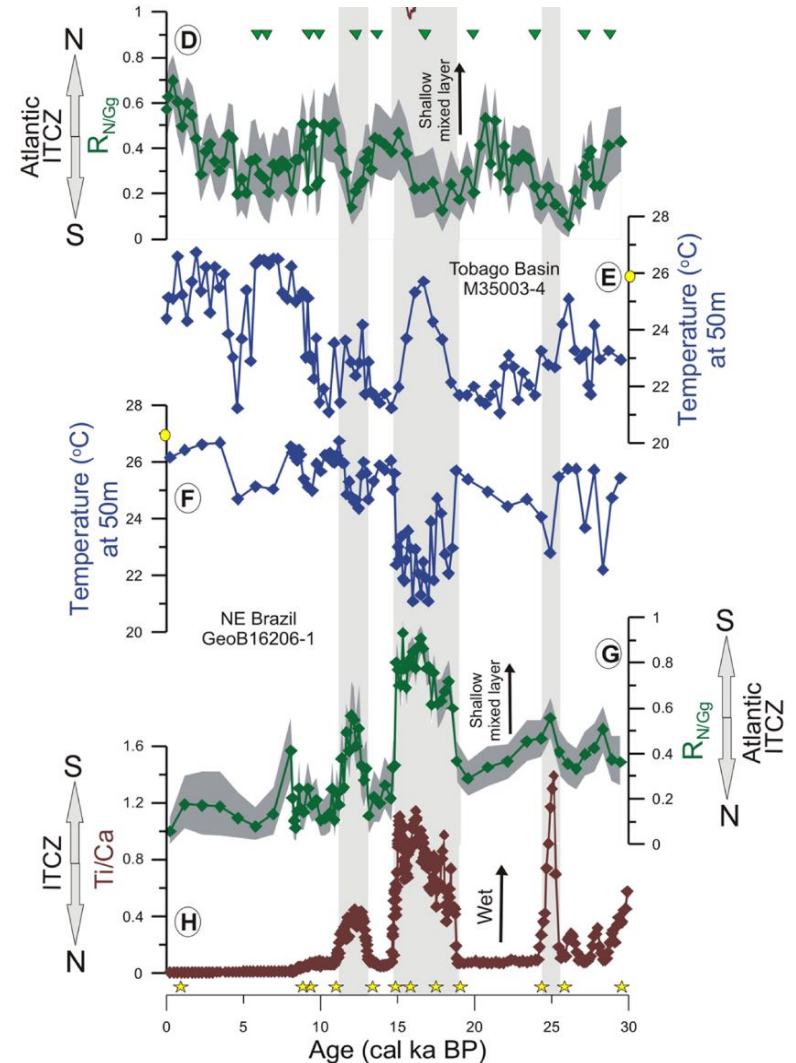
Paleoclimate records: *the editor's cut*

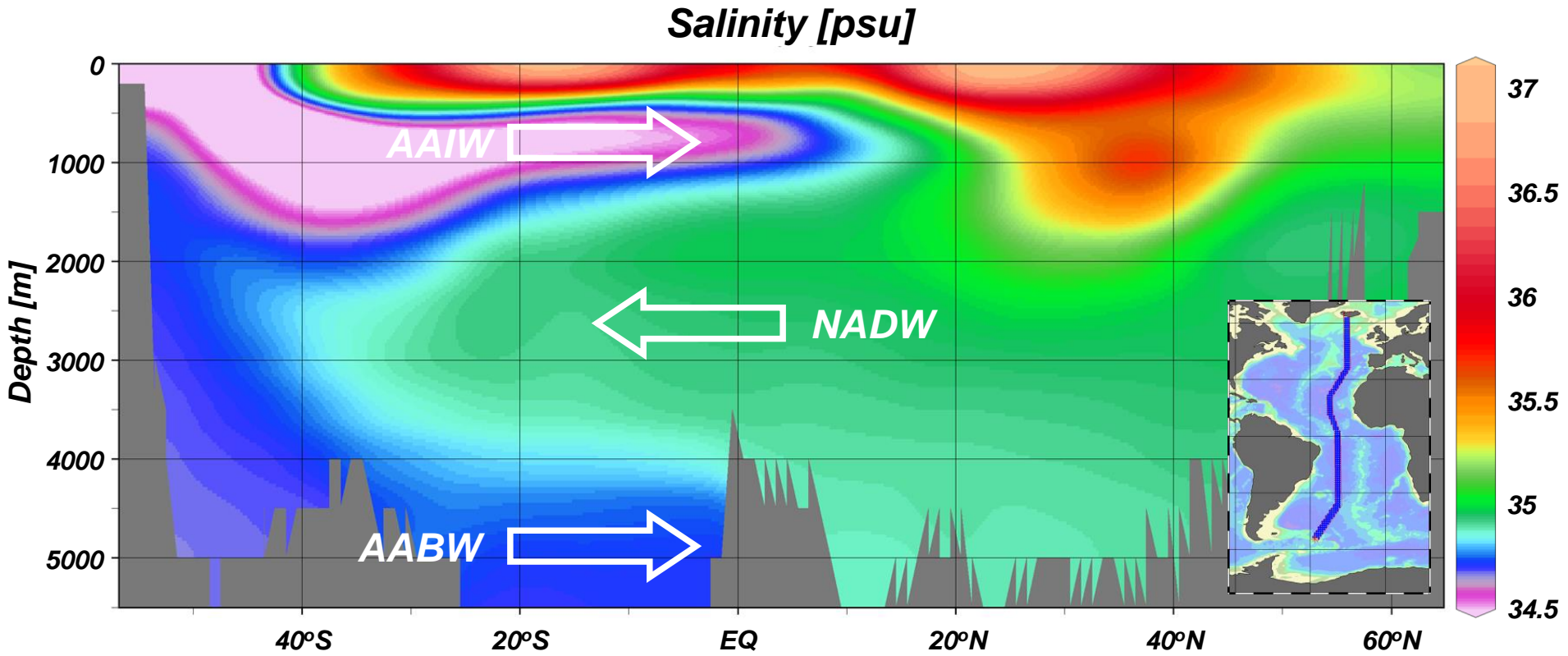


But, what happened with the shallowest mixed layer linked to the Intertropical Convergence Zone?

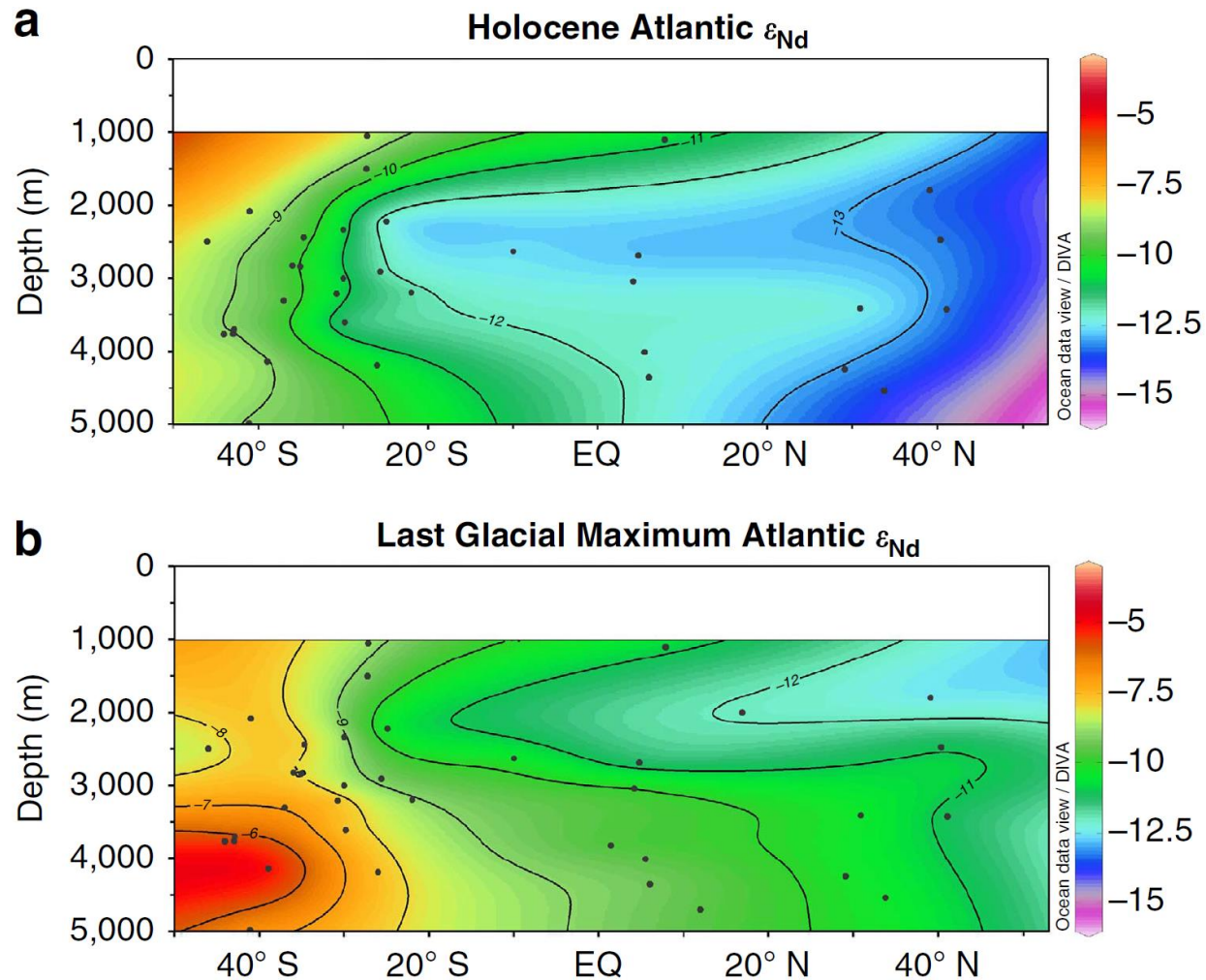


The shallowest mixed layer associated with the Intertropical Convergence Zone unambiguously shifted southwards during periods of a weakened AMOC





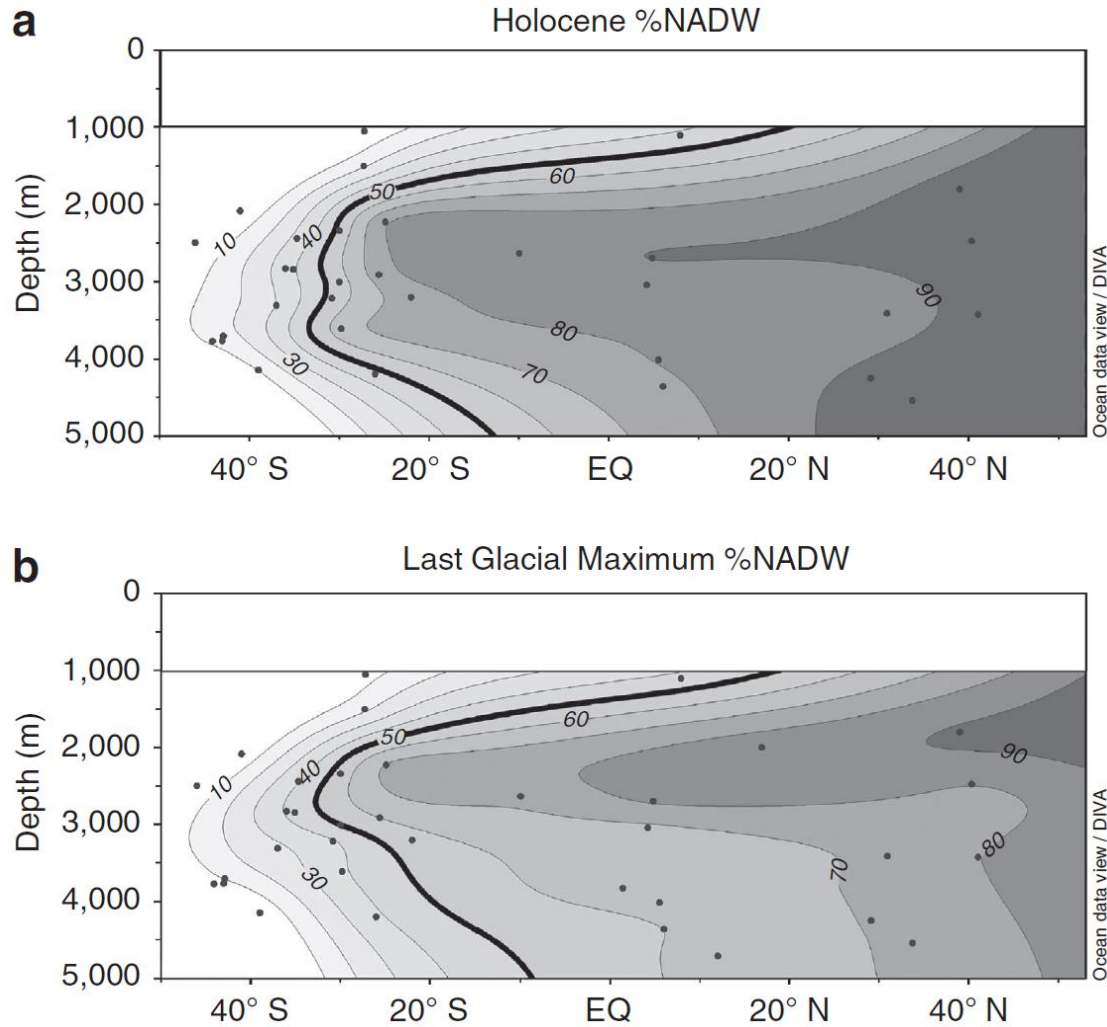
And, what about deep water geometry in the Atlantic during the Last Glacial Maximum?



Marked N-S gradient at deep waters suggest the sustained production of NADW and another northern-sourced water mass at intermediate water depth



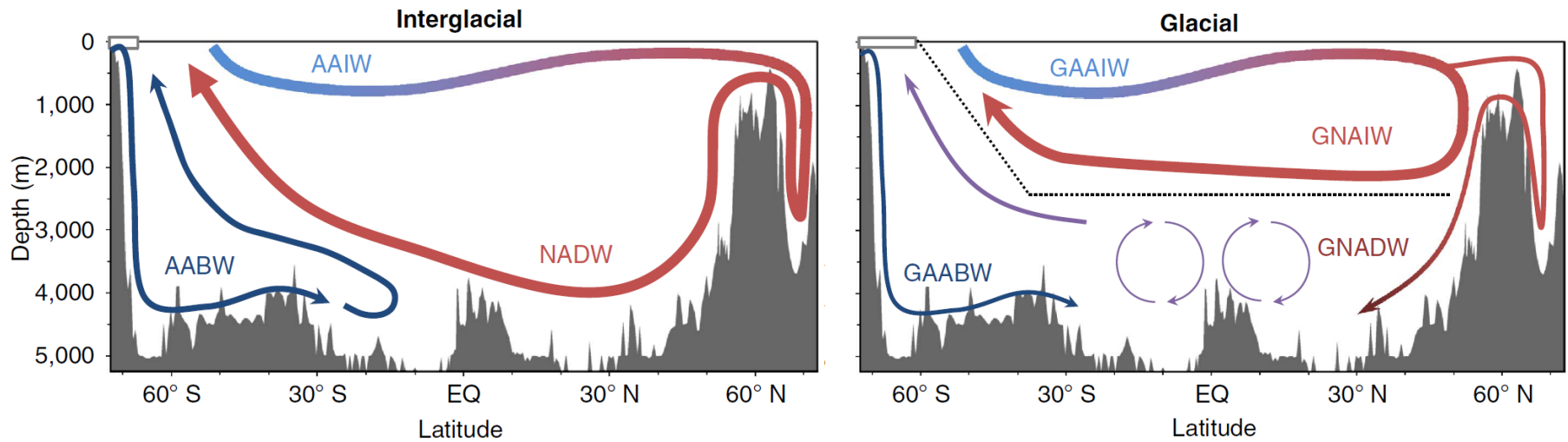
Paleoclimate records: *the editor's cut*



Marked N-S gradient at deep waters suggest the sustained production of NADW and another northern-sourced water mass at intermediate water depth



Paleoclimate records: *the editor's cut*



A greater amount of respired carbon must have been stored in the abyssal Atlantic during the Last Glacial Maximum by a sluggish deep overturning cell, comprised of well-mixed northern- and southern-sourced waters



Suggested reading



Lowe et al., 2007. Understanding Quaternary Climate Change. In: Elias (ed.), Encyclopedia of Quaternary Science, Elsevier, 28-39

Masson-Delmotte, et al. 2013. Information from Paleoclimate Archives. In: Stocker et al. (eds.), Climate Change 2013: The Physical Science Basis. Contribution of Working Group I to the Fifth Assessment Report of the Intergovernmental Panel on Climate Change, Cambridge University Press, 383-464

Ruddiman W.F. 2014. Earth's Climate: Past and Future. W.H. Freeman and Company, 445 p

SPSAS Climate Change

Many thanks for your attention!

Prof. Dr. Cristiano M. Chiessi
chiessi@usp.br

