

São Paulo School of Advanced Science on Ocean Interdisciplinary Research and Governance

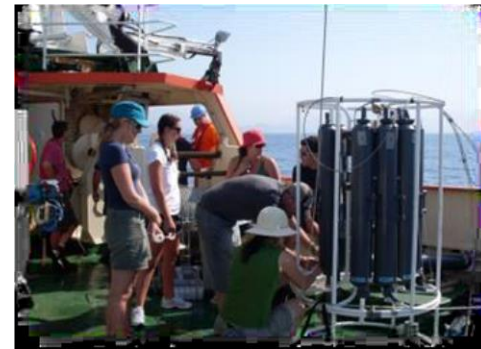
Climate Change and (Ocean) Variability

The memory of climate

Silvia I. Romero

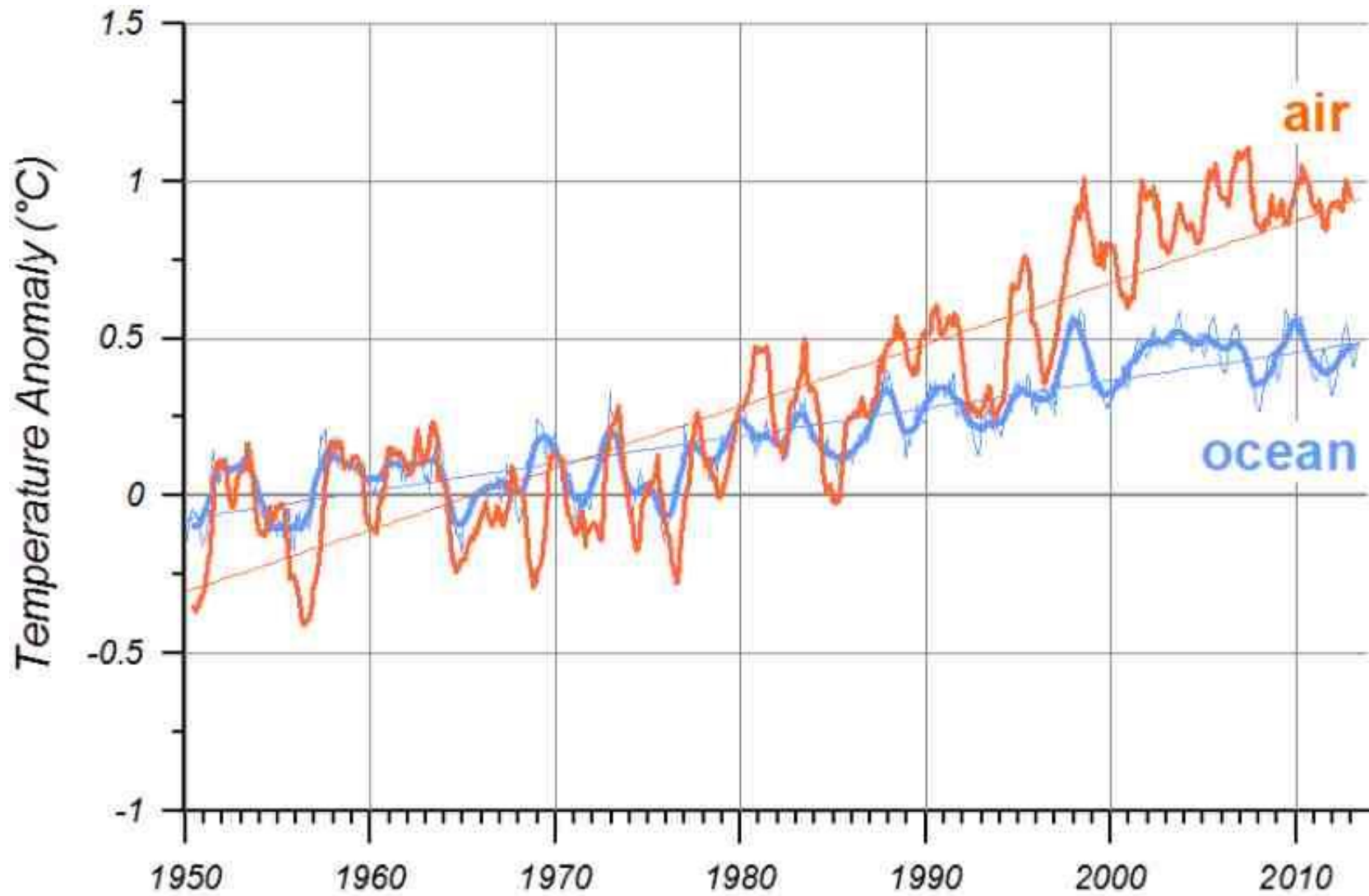
Departamento Oceanografía, Servicio de Hidrografía Naval,
Departamento Ciencias de la Atmosfera y los Océanos, Universidad de Buenos Aires,
Universidad de la Defensa,
Unité Mixte International CNRS/CONICET, Argentina

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- **Observations**
 - temperature (heat content),
 - sea level,
 - salinity and the hydrological cycle
- **Meridional overturning**
 - variability
 - impact on climate

Global warming – atmosphere and ocean



<http://vlb.ncdc.noaa.gov/cmb-faq/anomalies.php#mean>

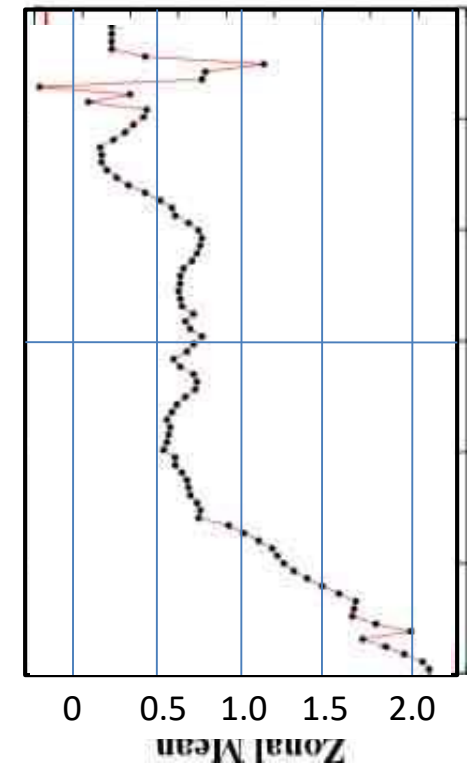
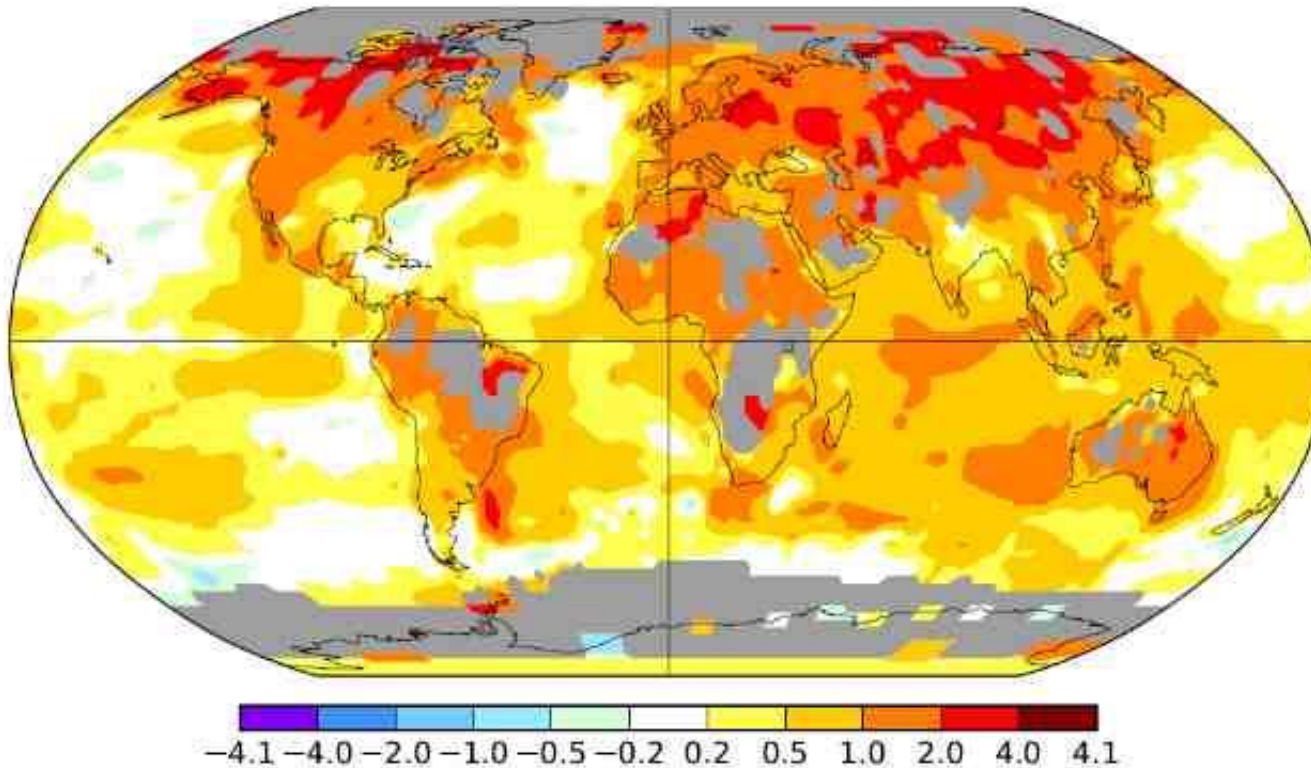
Sea surface temperature trend

1951-2015, °C

Annual J-D

L-OTI(°C) Change 1951-2015

0.73

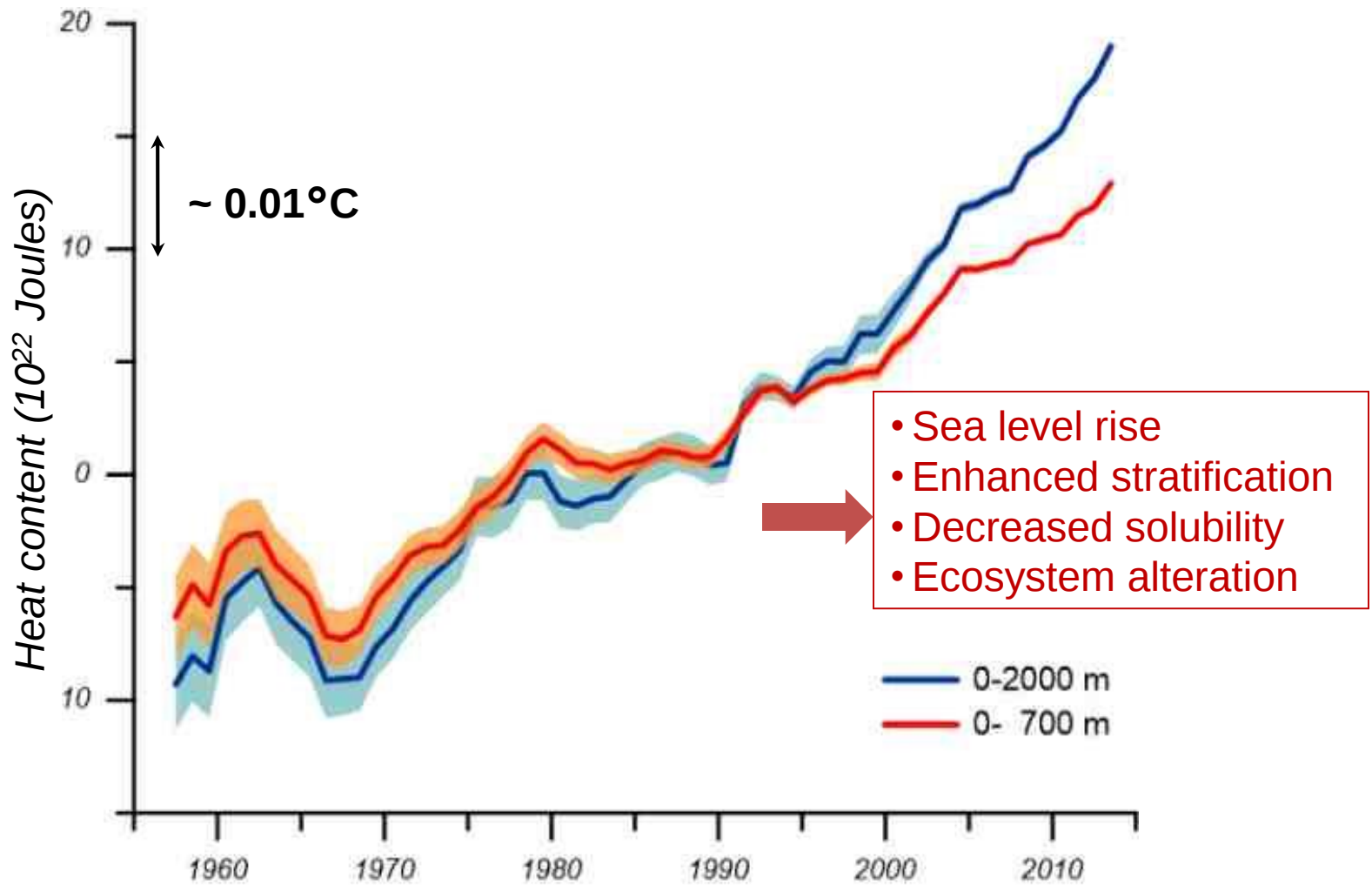


Data Sources

Land: Goddard Inst. Space Studies - *Ocean:* Hadley/Reynolds v.2

<http://data.giss.nasa.gov/gistemp/maps/>

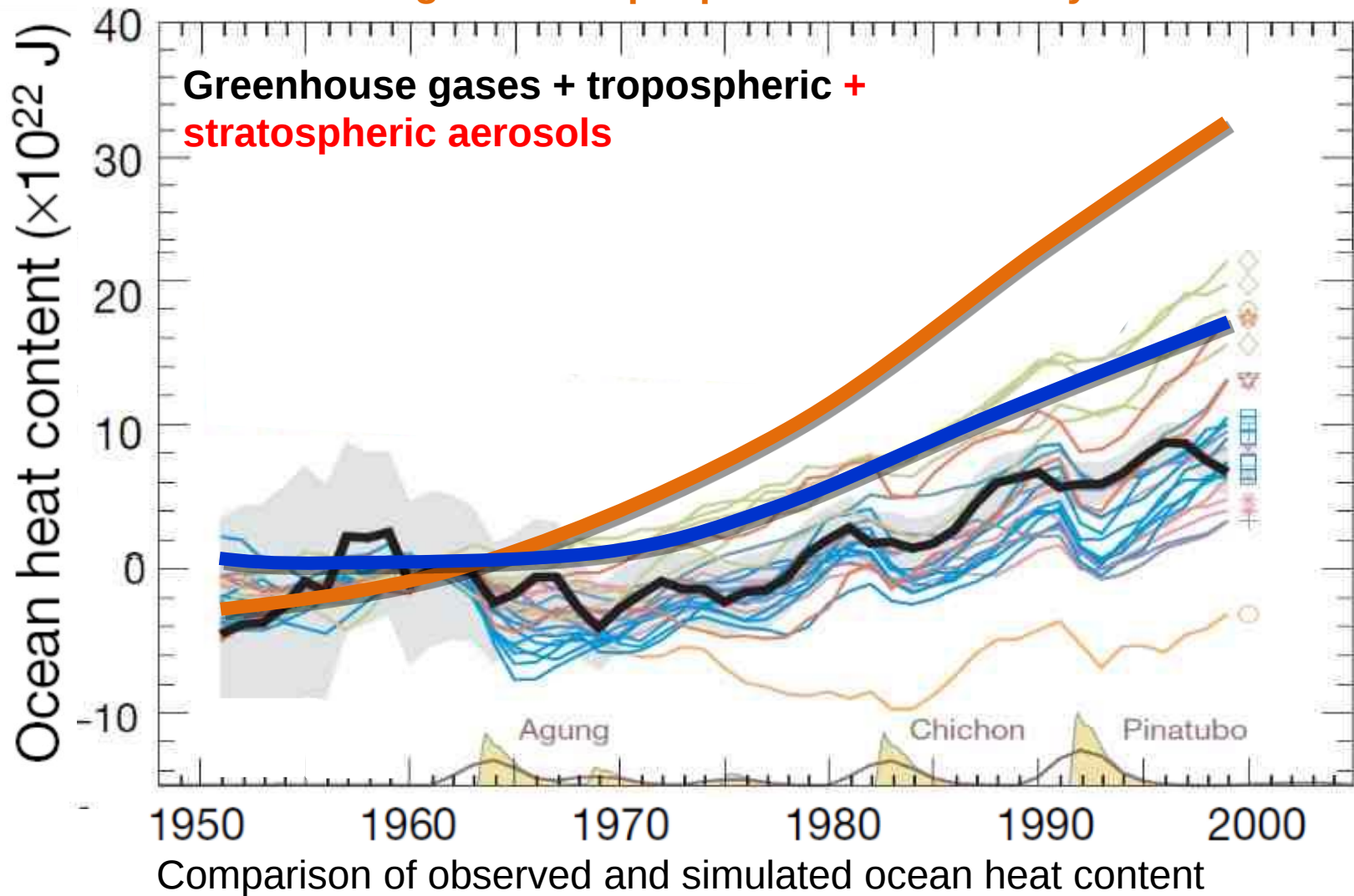
Changes in ocean heat content



Data from https://www.nodc.noaa.gov/OC5/3M_HEAT_CONTENT/

Attribution of warming trends in the upper ocean

Greenhouse gases + tropospheric aerosols only

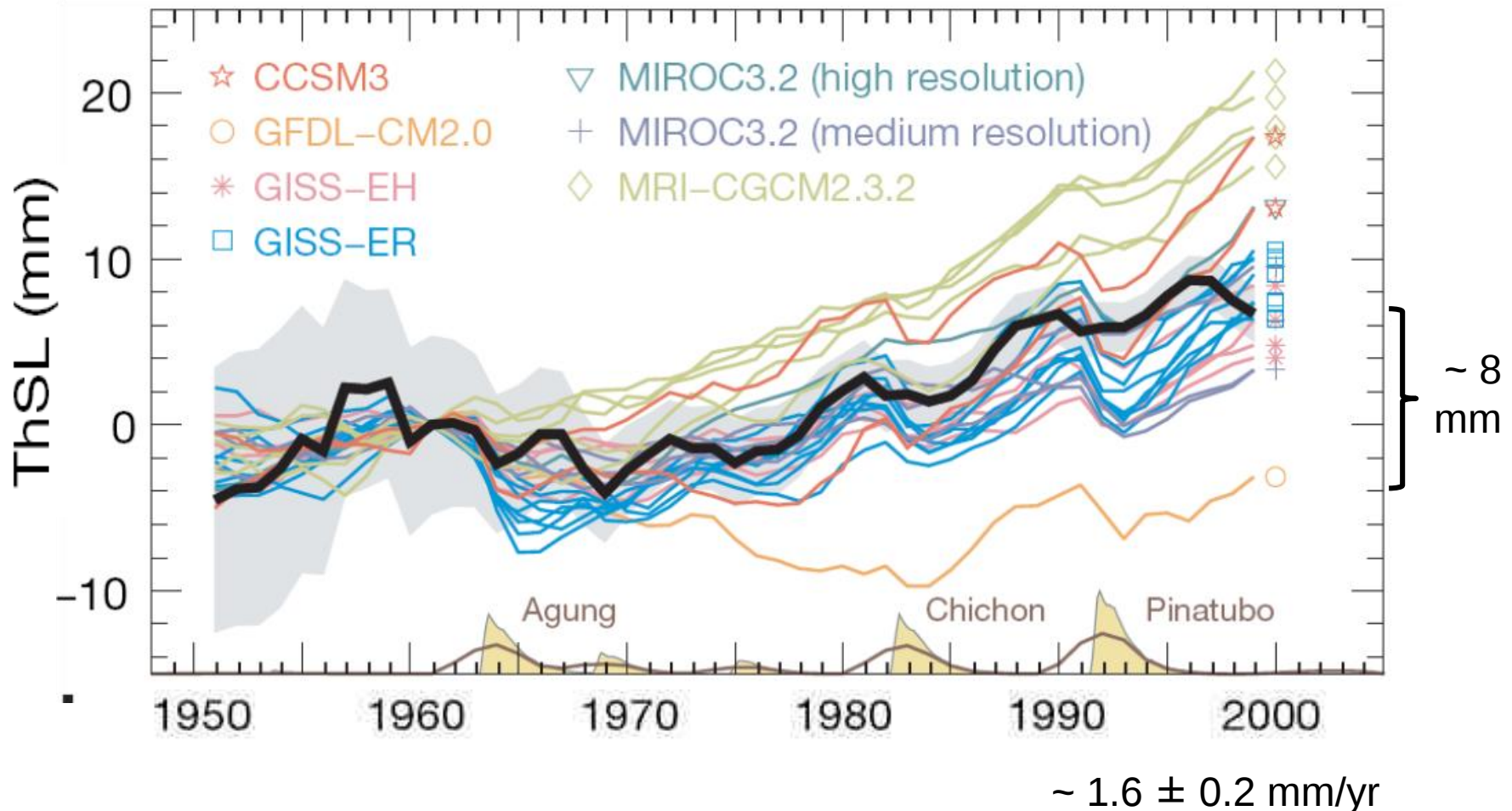


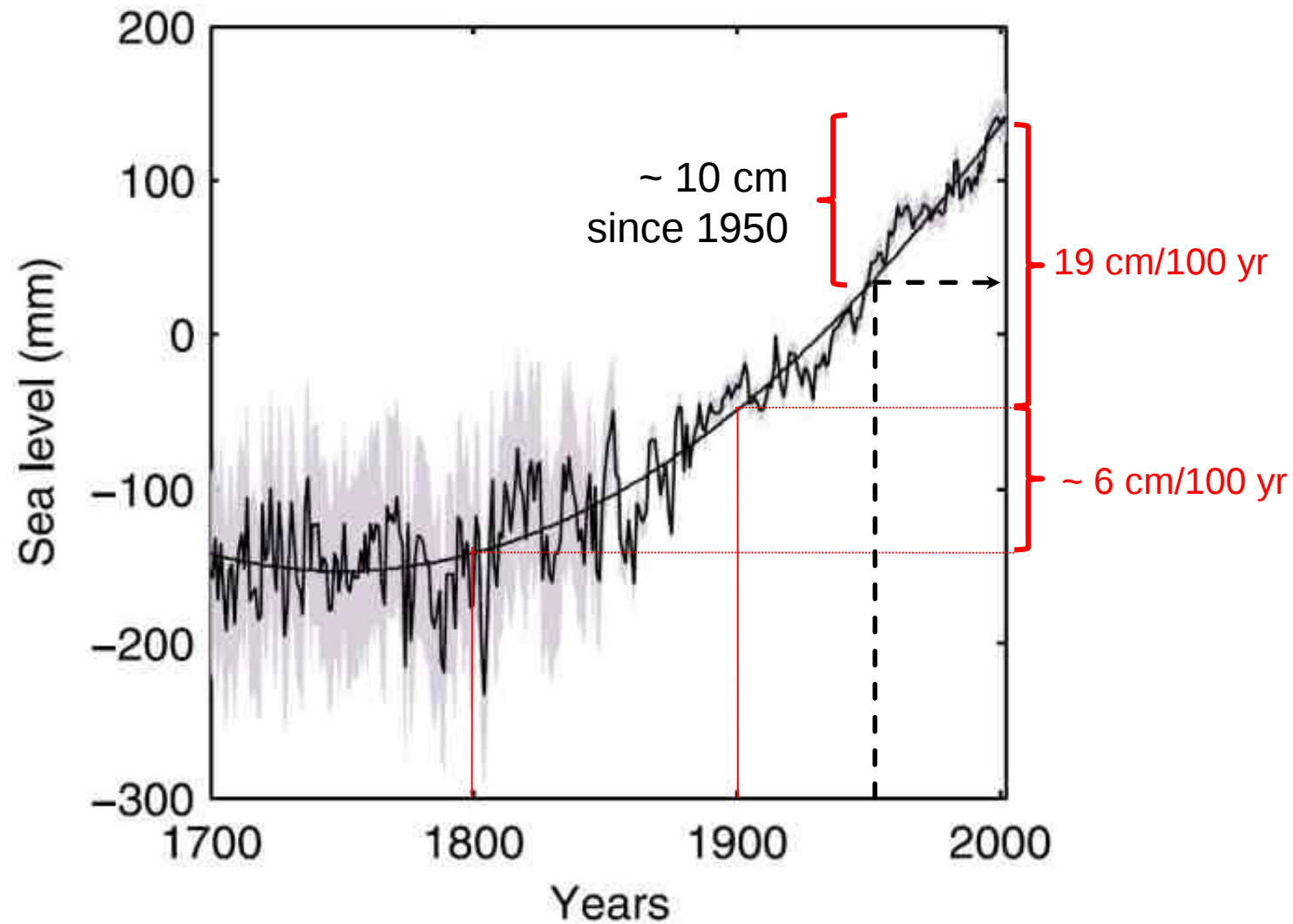
Sea level rise (from warming)

Recall thermal expansion $\alpha \Delta T = \Delta V/V$

warmer ocean $\Rightarrow$ larger volume

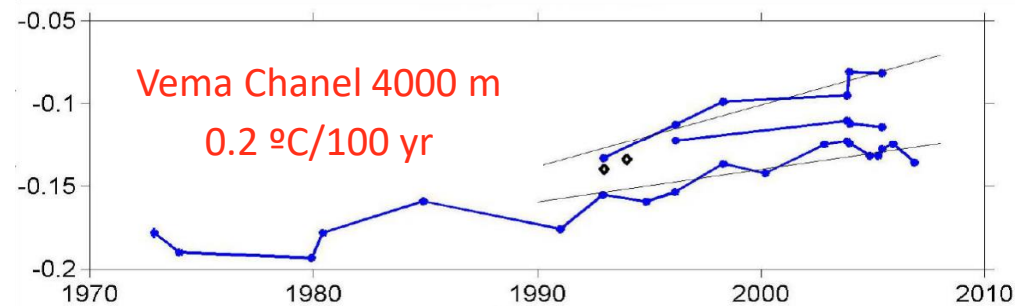
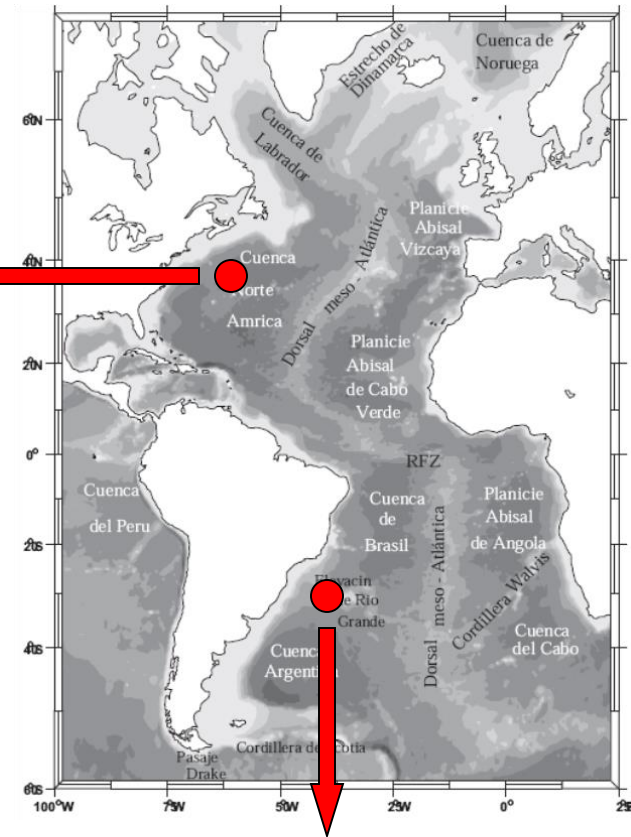
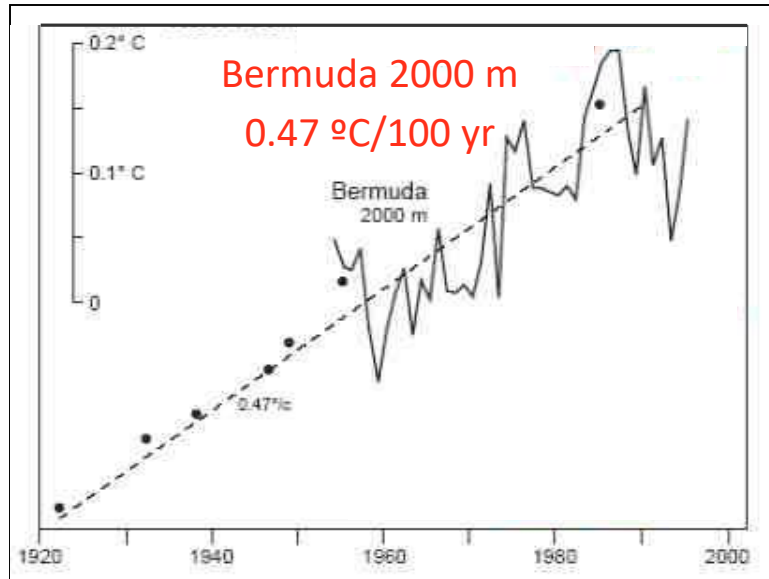
Greenhouse gases + tropospheric and stratospheric aerosols





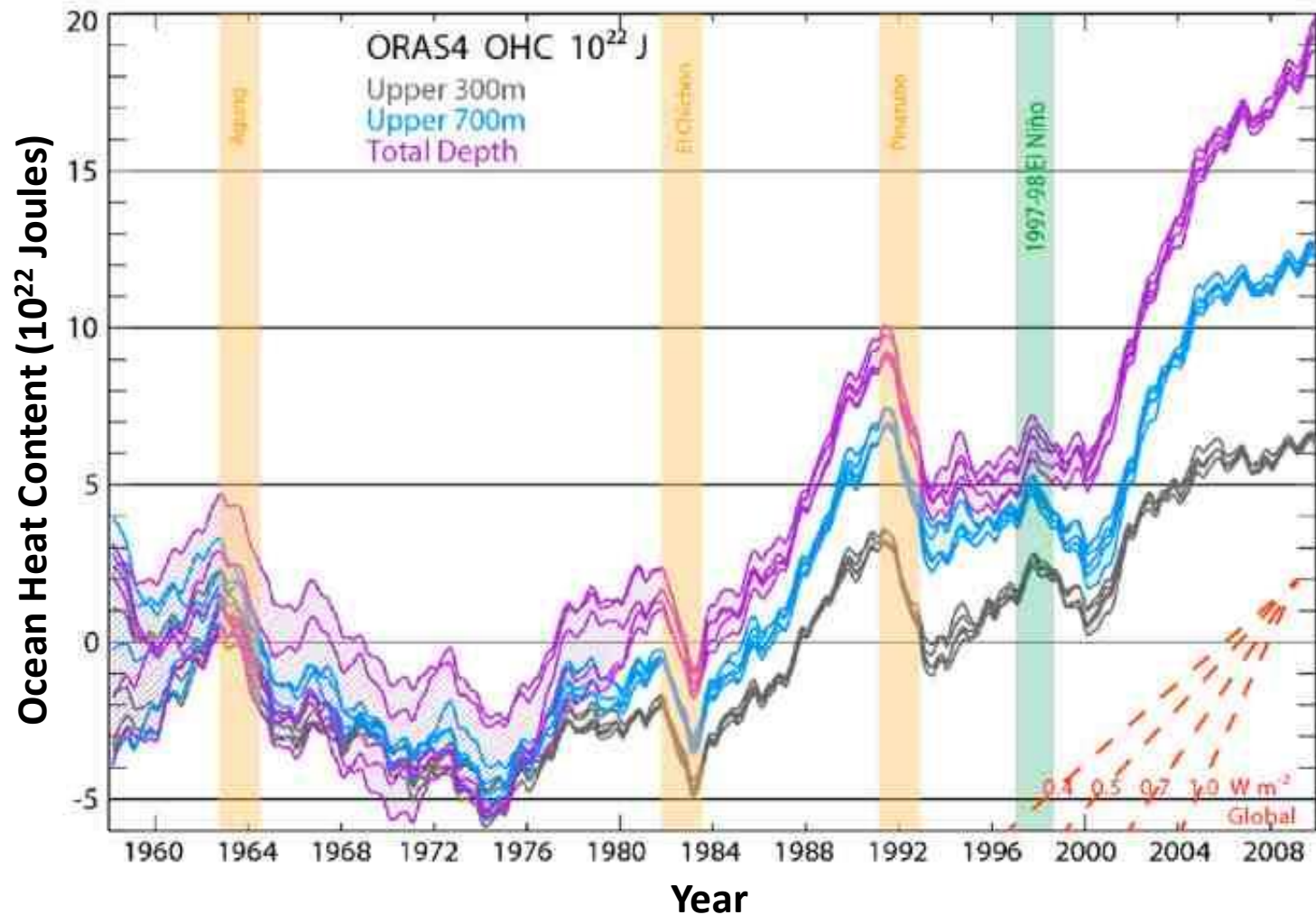
Recent rise started in late 19th century. Trend is non linear ($\sim 0.01 \text{ mm/yr}^2$)

Temperature trends in the deep ocean



Zenk & Morozov, *GRL*, 2007

Recent changes in upper ocean heat content

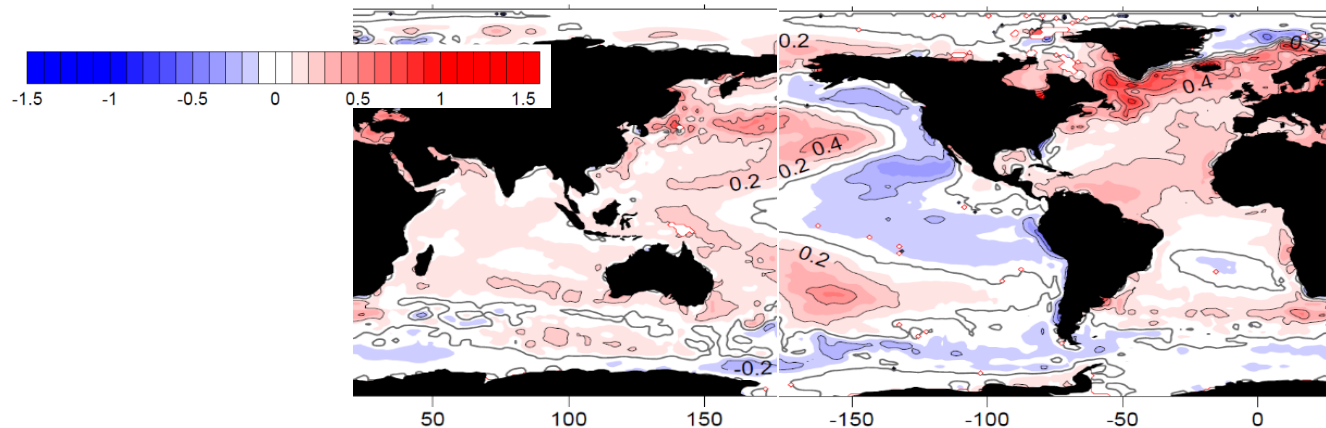
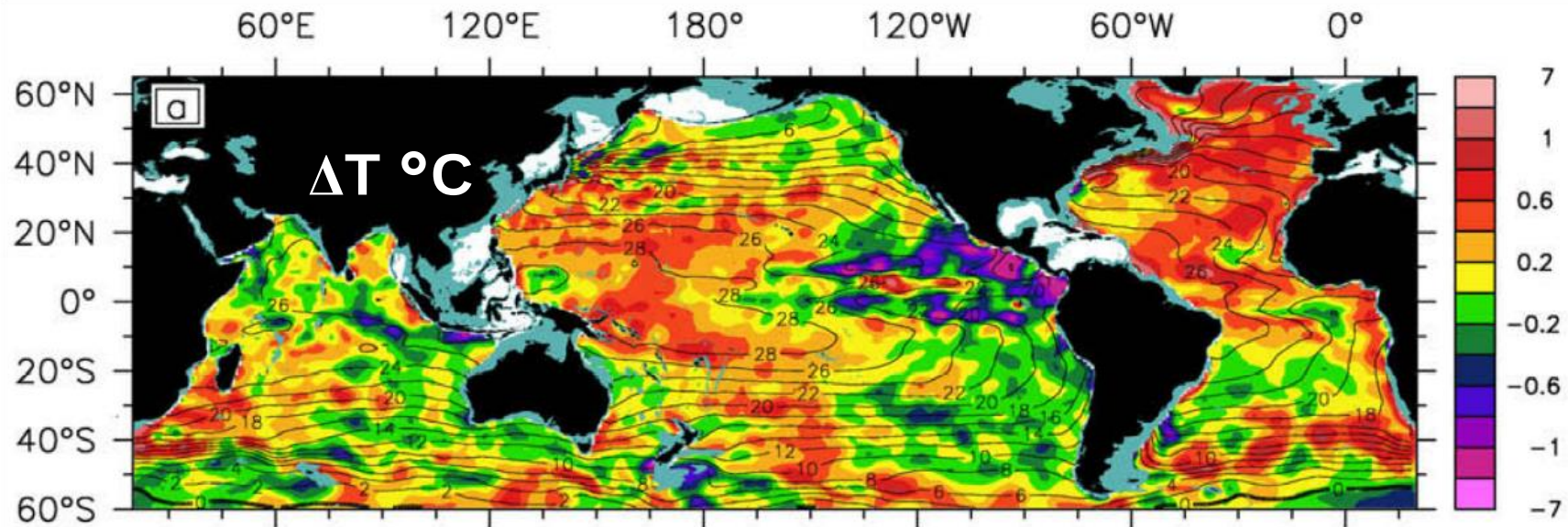


Heat content from an ocean reanalysis based on forced ocean model and T and S observations from various sources. 5 ensembles are shown for each layer.

Global temperature changes - I

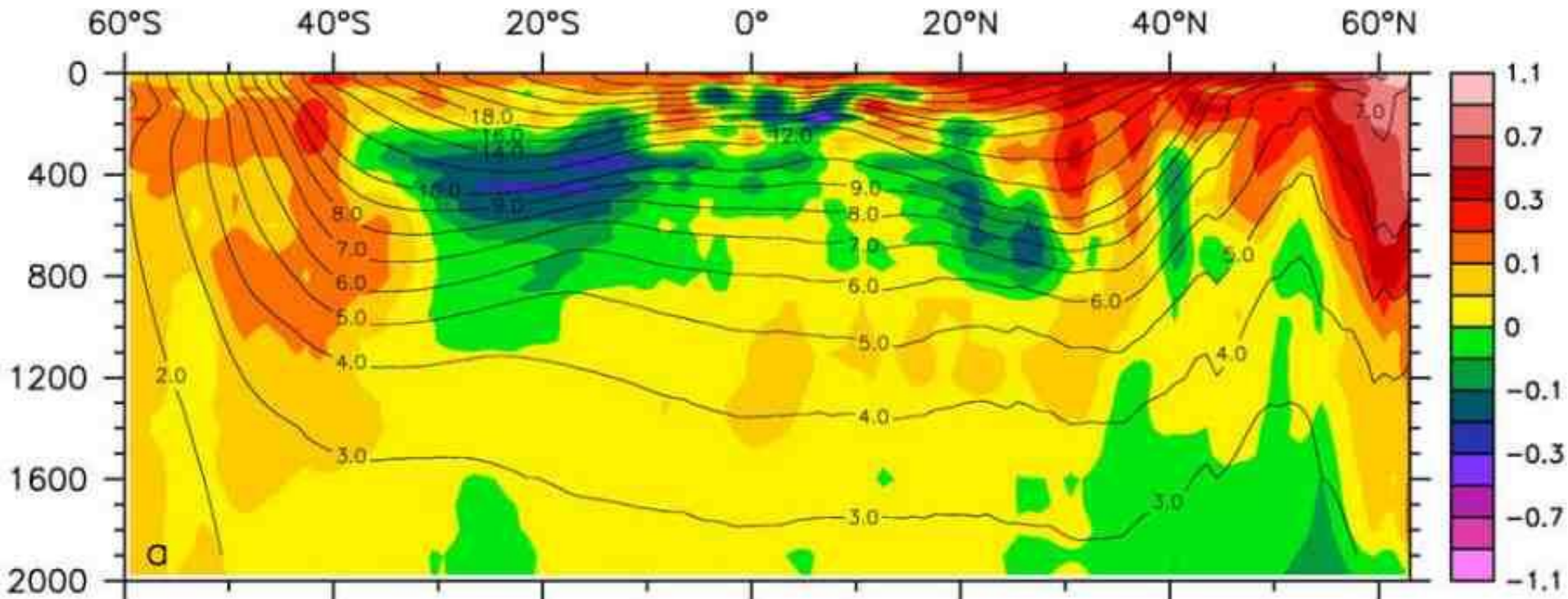
Argo 5yr (2004-2008) mean- WOA01
0-100 dbar

Roemmich & Gilson, *Progr. Oceanogr.*, 2009



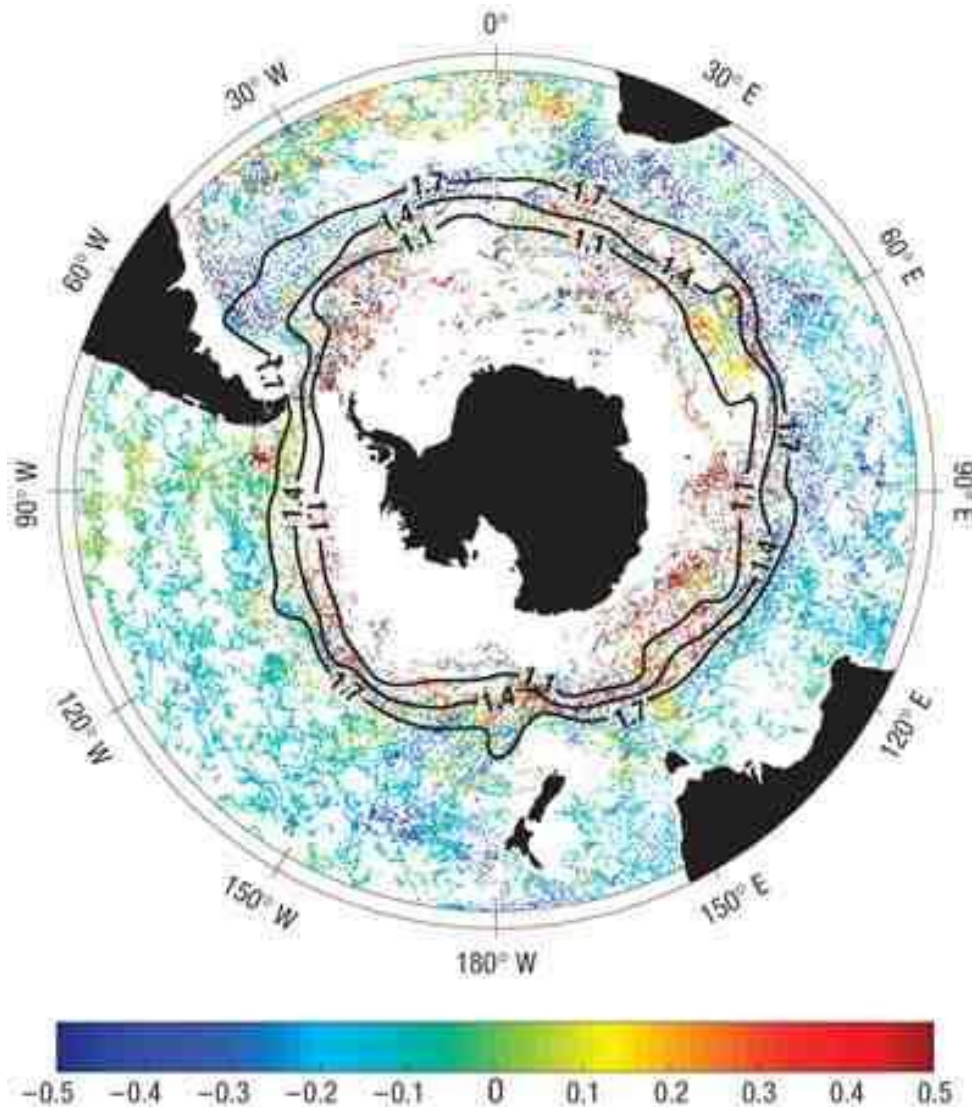
Reynolds SST
trend $^{\circ}\text{C}/\text{decade}$
(1981-2012)

Global temperature changes - II



Zonally averaged temperature versus pressure for the Argo 5-year mean (contours) and Argo-minus-WOA01 difference (color shading).

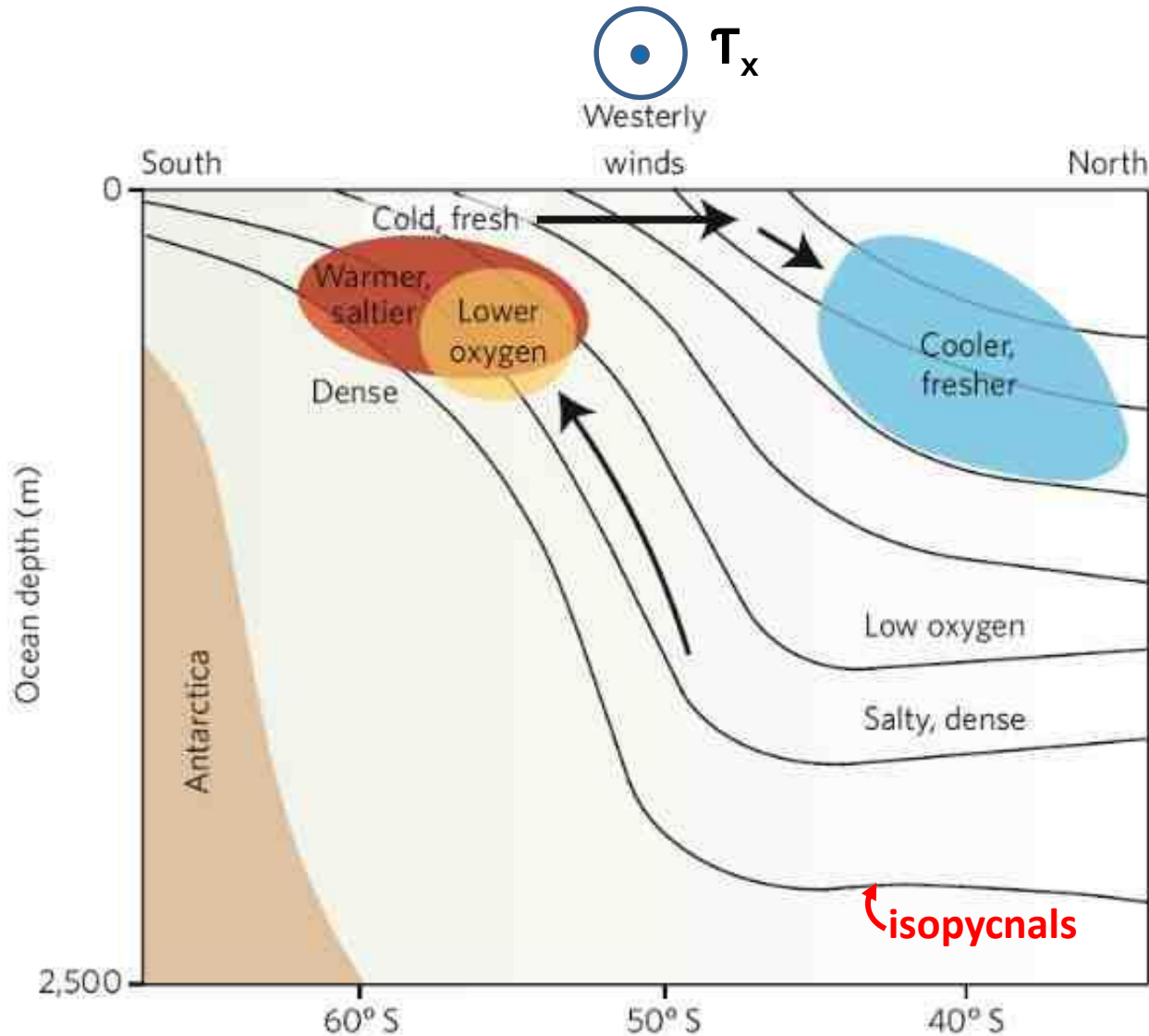
Temperature changes Southern Hemisphere



Recent temperature changes in the 26.9-27.7 kg/m³ neutral density range from a comparison of 52400 Argo profiles and climatology. Lines represent geostrophic streamlines of the Antarctic Circumpolar Current approximately following the Subantarctic and Antarctic (Polar) front (1.7 y 1.4 dyn meters, respectively).

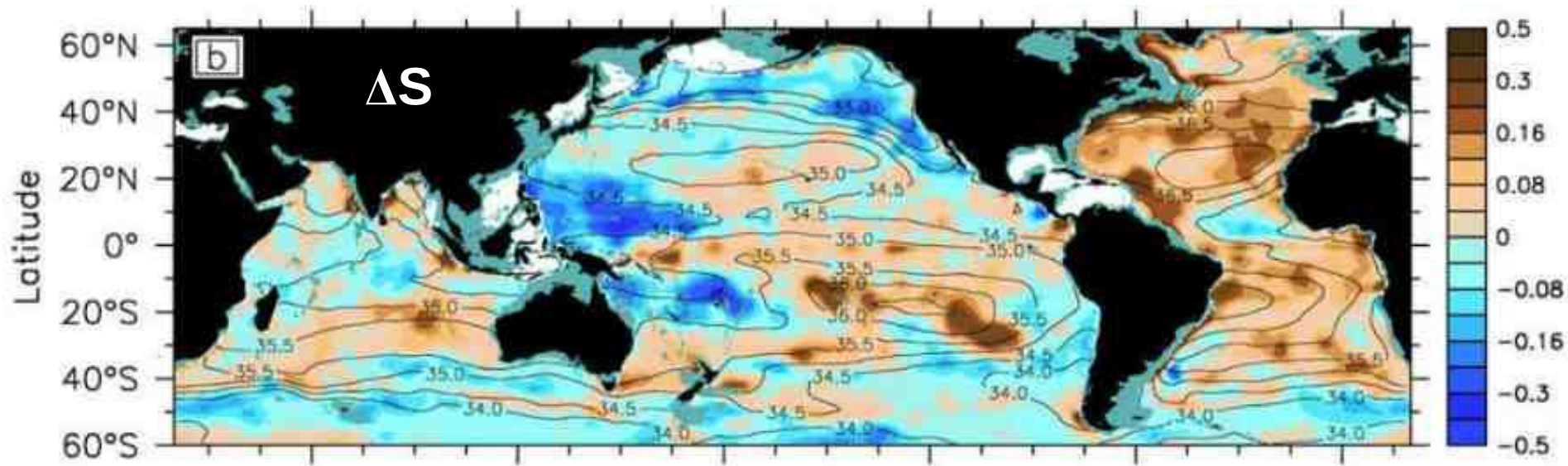
[Boning et al., *Nature Geosc.*, 2008](#)

Increased SH westerlies and expected variability



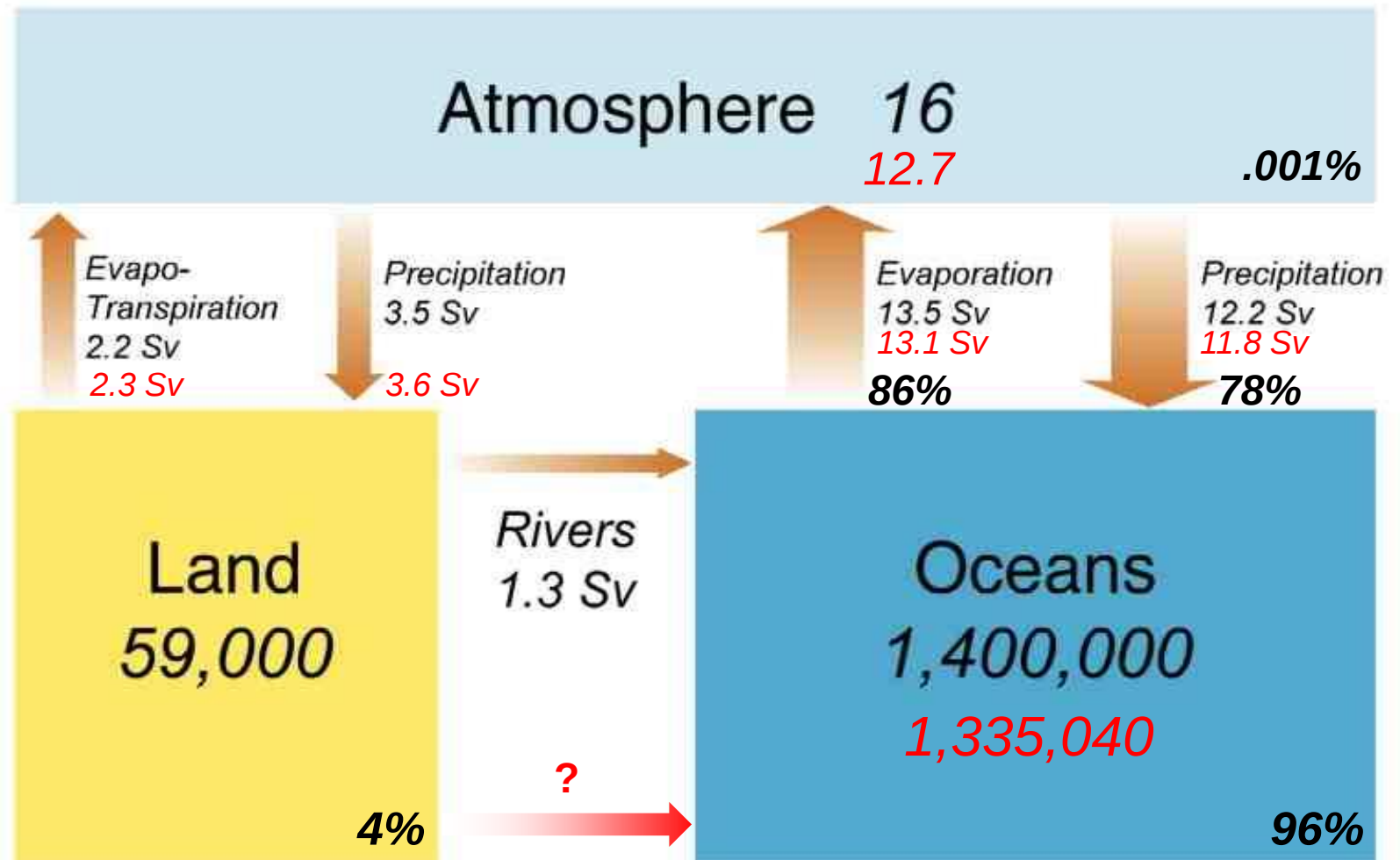
Schematic density section across the ACC and the expected temperature, salinity and dissolved oxygen changes due to an increase in SH westerlies. (Aoki et al., *GRL*, 2005; Togwiller et al., *Nature*, 2009).

Argo 5yr (2004-2008) mean- WOA01
0-100 dbar



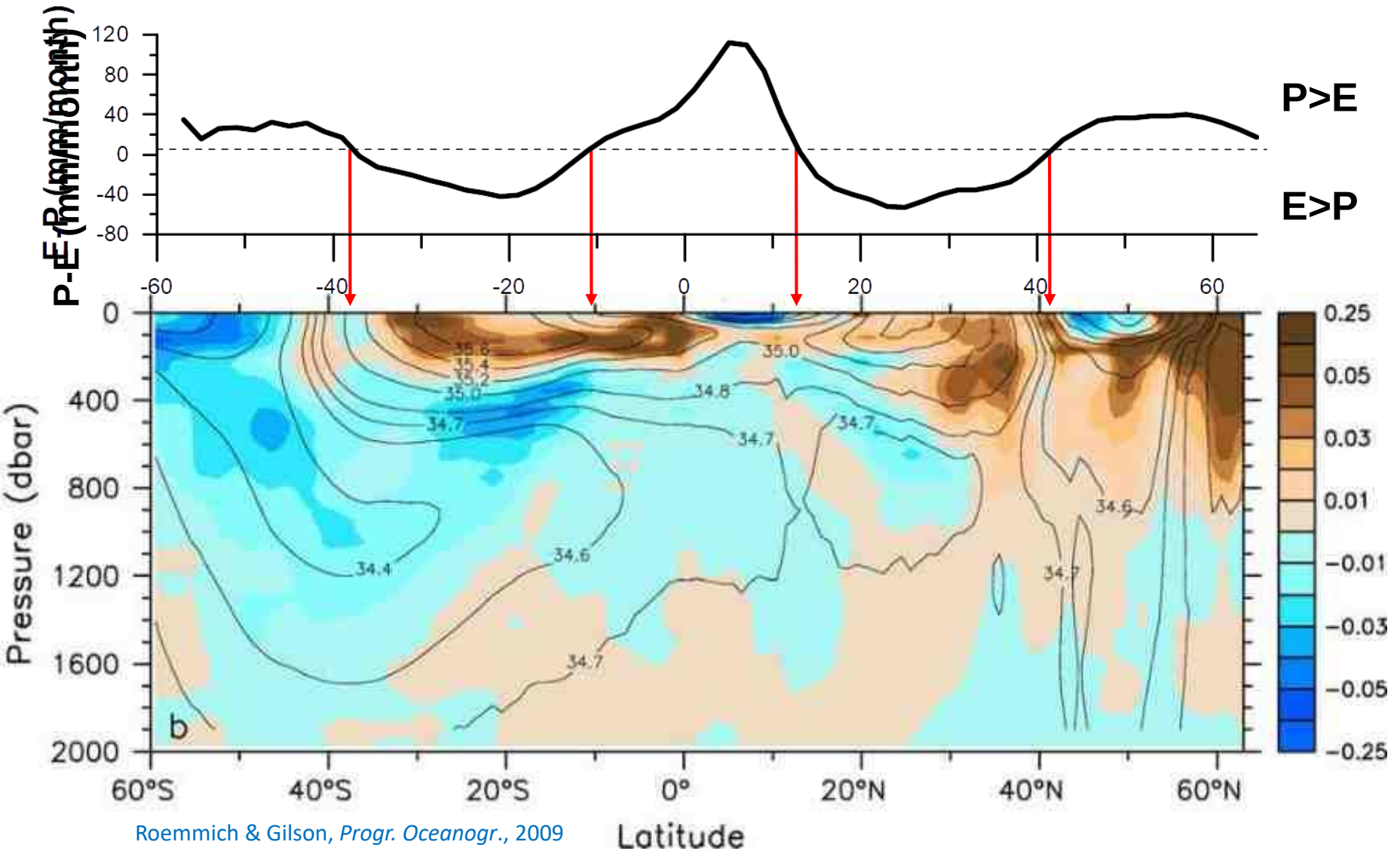
Map of 0–100 dbar vertically averaged salinity from the Argo 5-year mean (contours) and Argo-minus-WOA01 difference (color shading)





Reservoirs in 10³ km³, Fluxes in 10⁶ m³/s (= Sv)

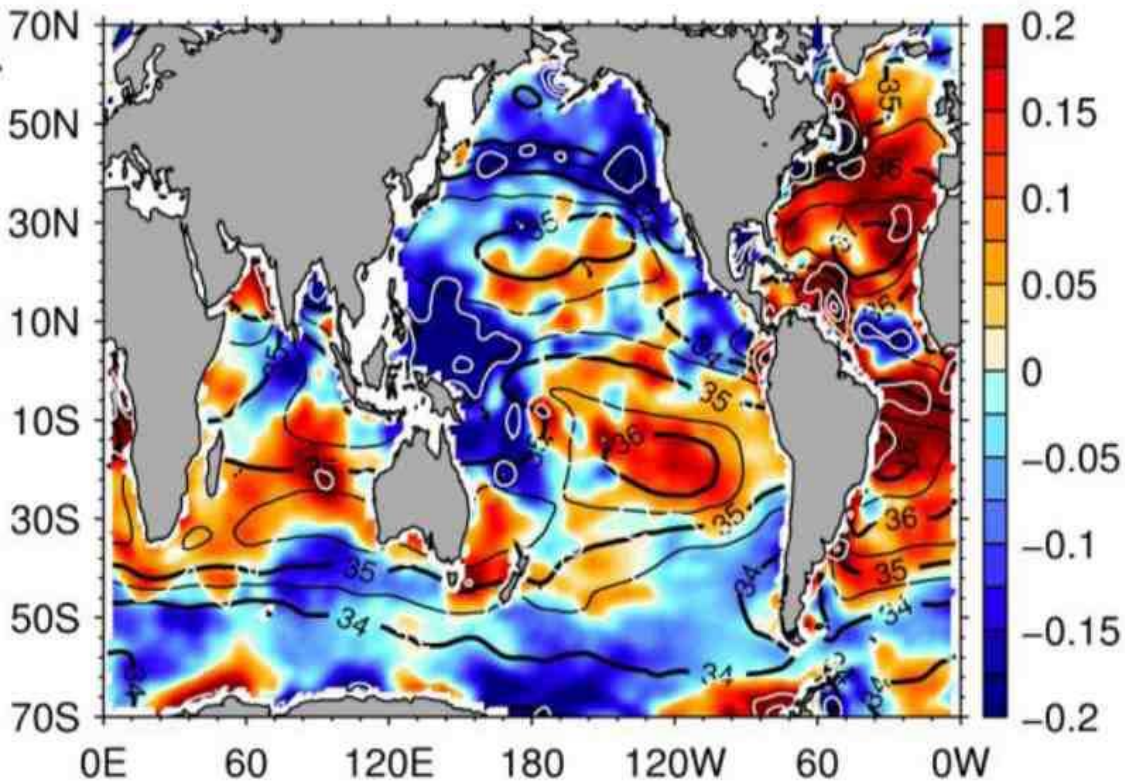
Global salinity changes - II



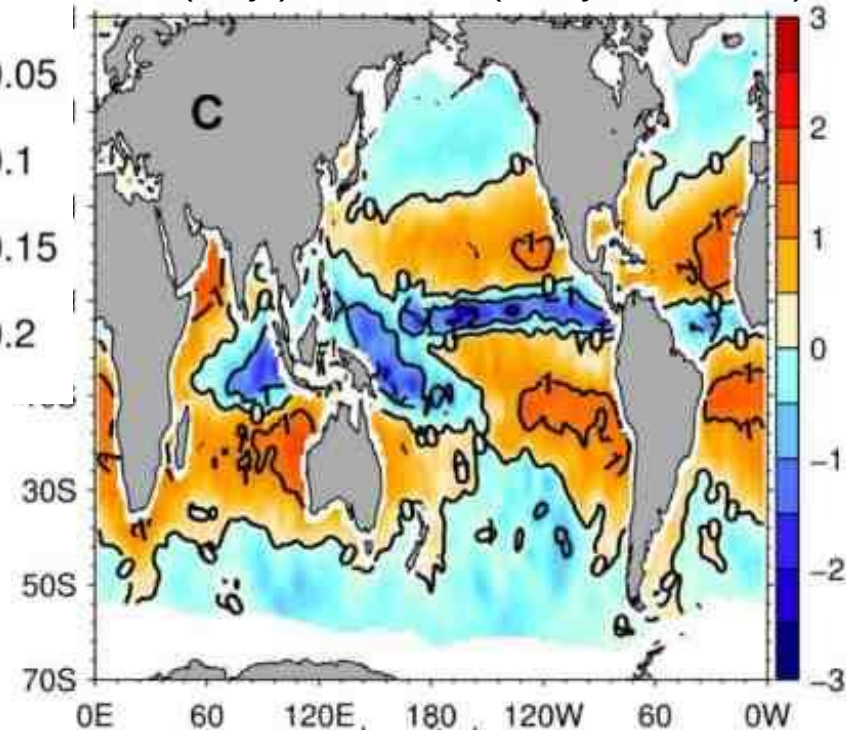
Zonally averaged salinity versus pressure for the Argo 5-year mean (contours) and Argo-minus-WOA01 difference (color shading).

Global salinity changes - III

50-year (1950-2000) linear surface salinity trend (pss/50 yr)

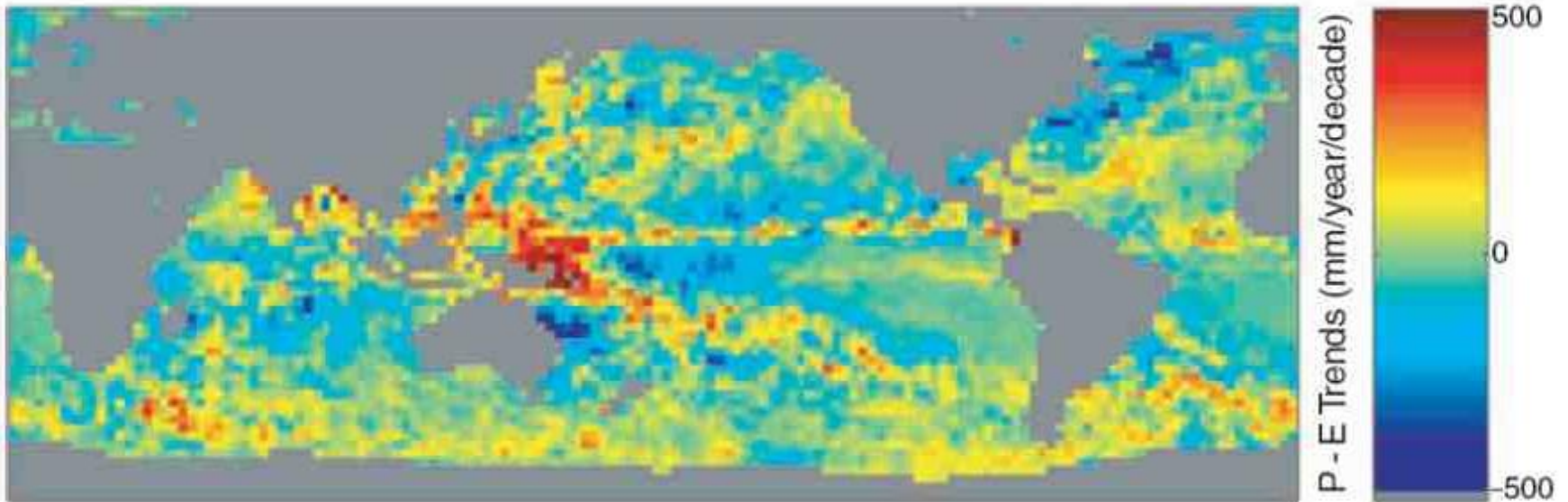


E-P (m^3/yr) 1980-1993 (Josey et al., 1998)



Trends in satellite-derived P – E

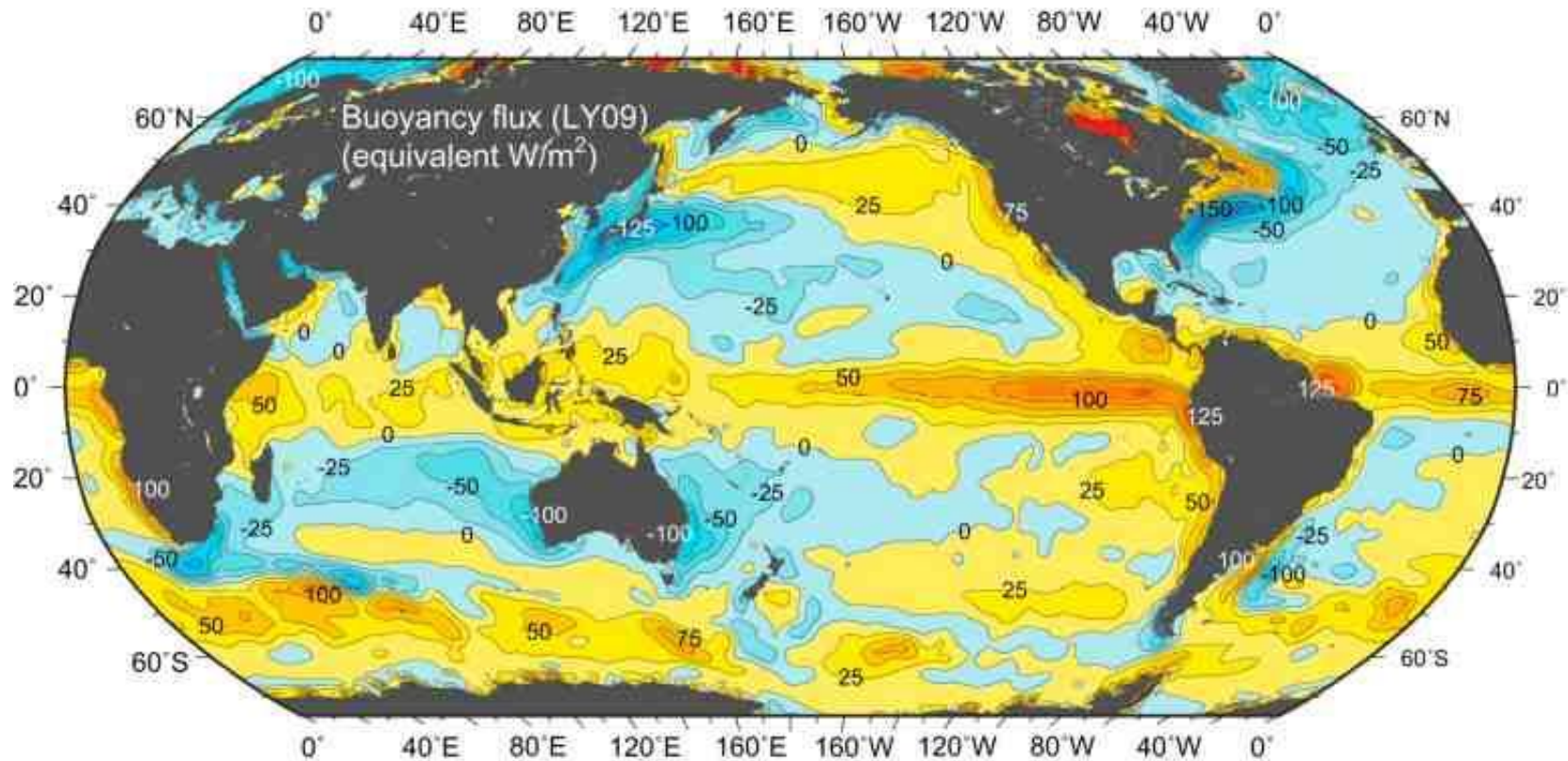
July 1987 - August 2006



Wentz et al., *Science* 2007

Surface thermohaline modification

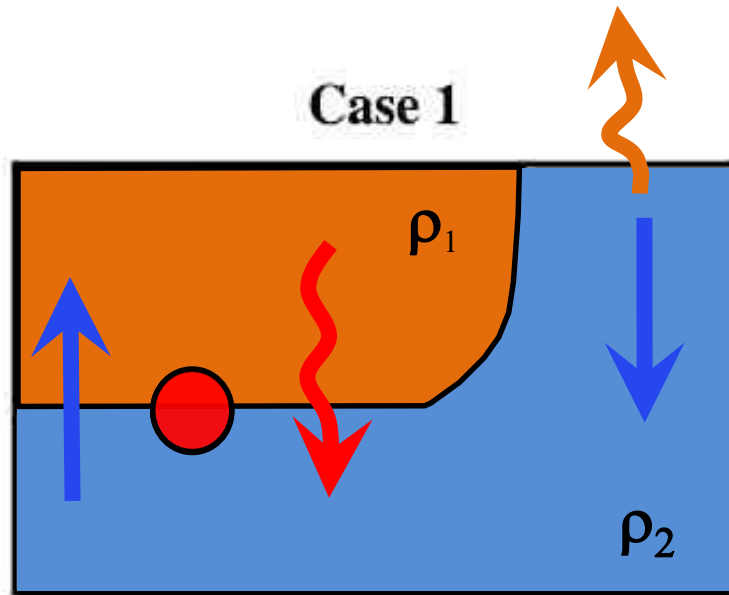
Surface Buoyancy Flux



Annual mean air-sea buoyancy flux converted to equivalent heat fluxes (W/m^2), based on Large and Yeager (2009) air-sea fluxes. **Positive values indicate that the ocean becomes less dense.** Contour interval is 25 W/m^2 .

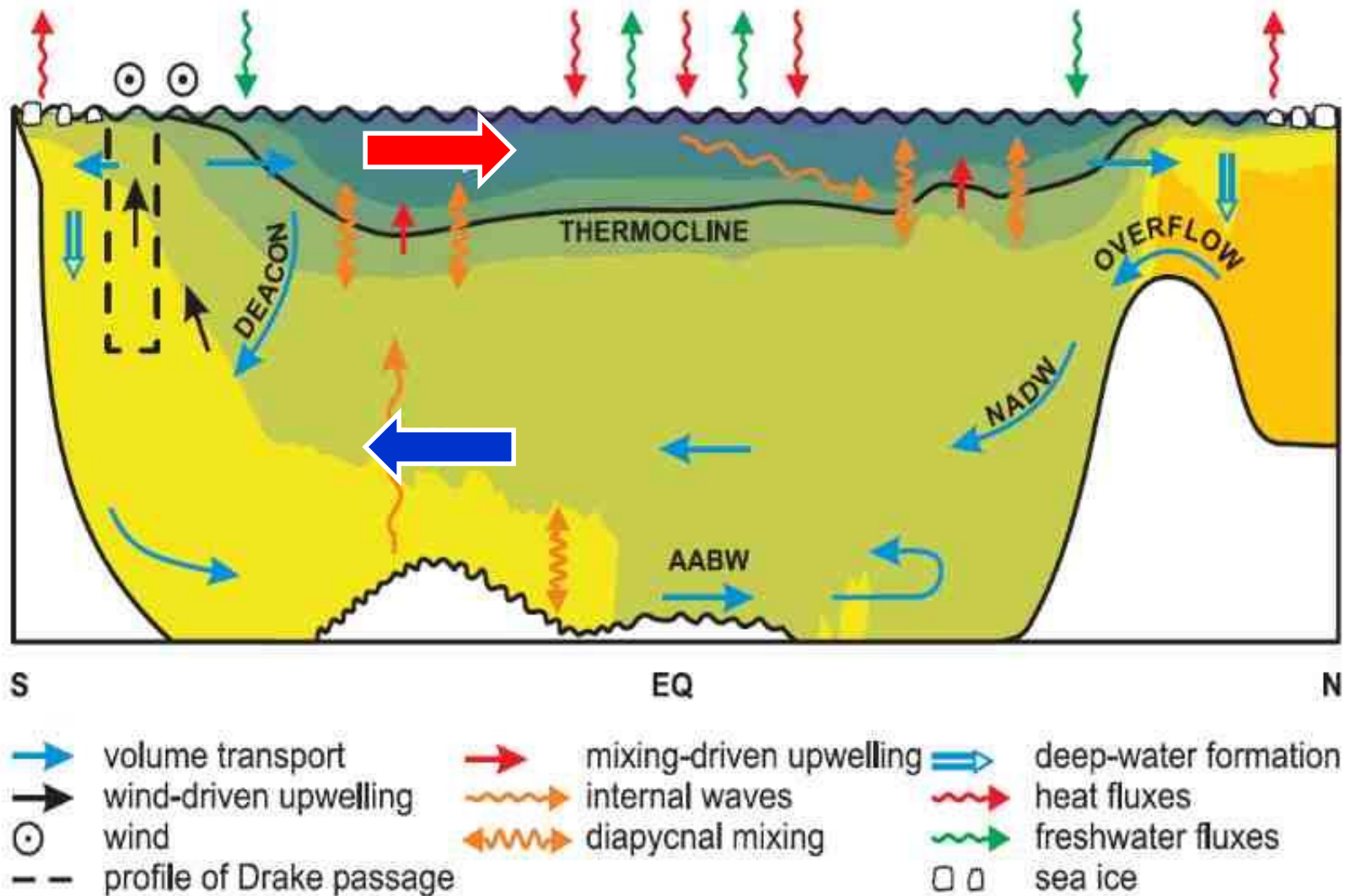
Extreme cases

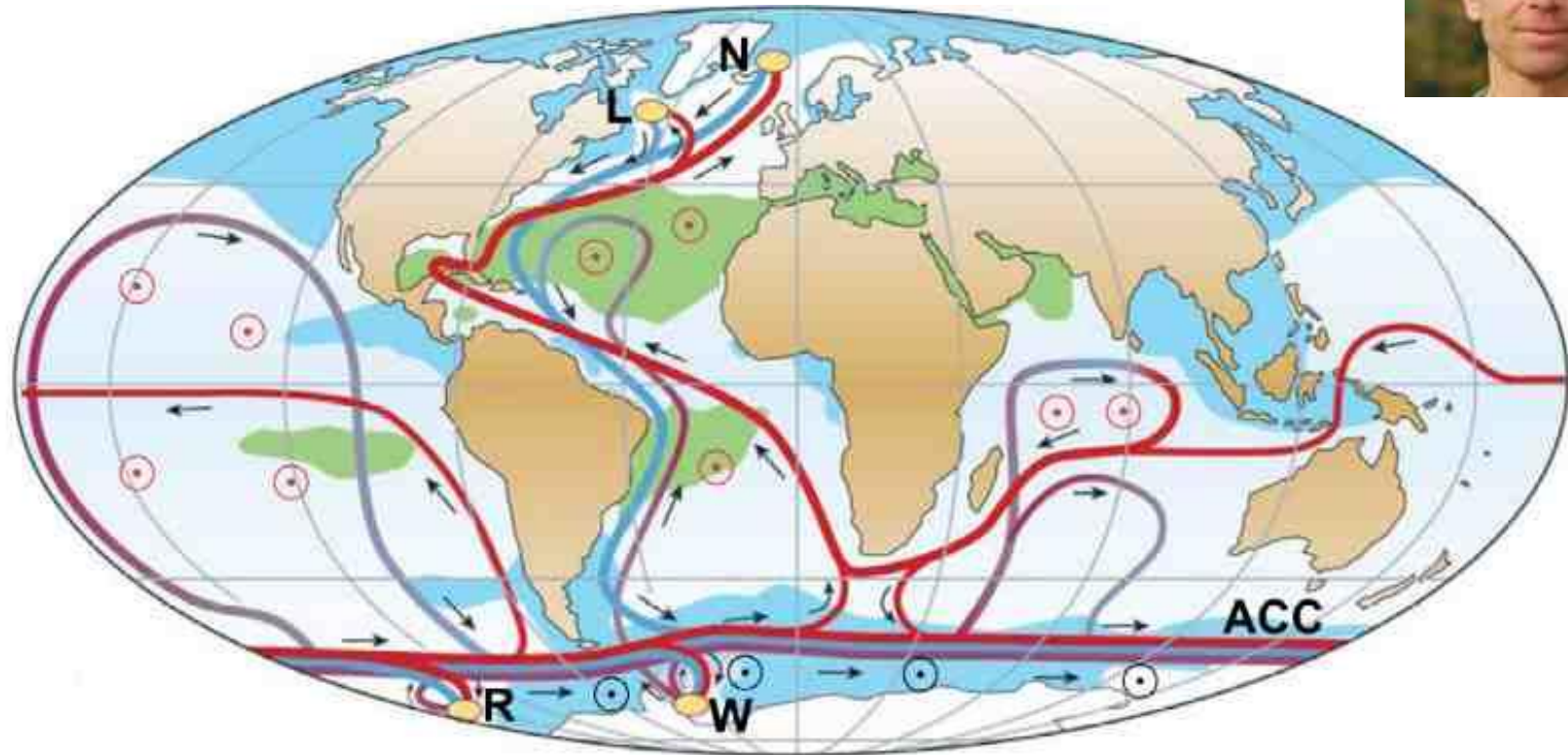
$$w \frac{\partial T}{\partial z} = \kappa \frac{\partial^2 T}{\partial z^2}$$



 dense to light water conversion

Overtaking, more realistic (?)





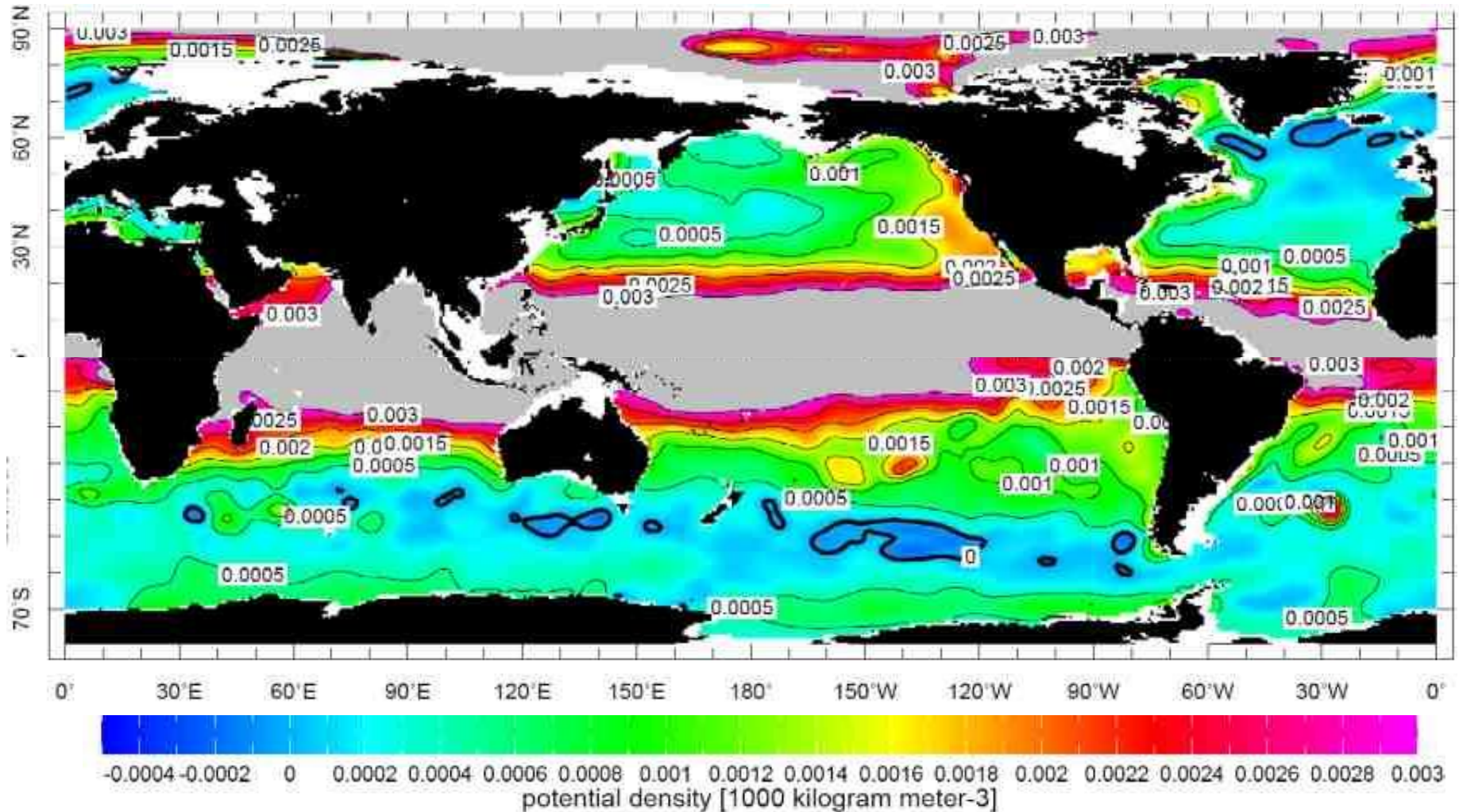
- Surface flow
- Deep flow
- Bottom flow
- Deep Water Formation

- ⊙ Wind-driven upwelling
- ⊙ Mixing-driven upwelling
- Salinity > 36 ‰
- Salinity < 34 ‰

- L** Labrador Sea
- N** Nordic Seas
- W** Weddell Sea
- R** Ross Sea

Vertical stratification – observations

$\Delta\sigma_\theta$ (kg/m³) 0-300 m winter (Feb-NH / Aug-SH)



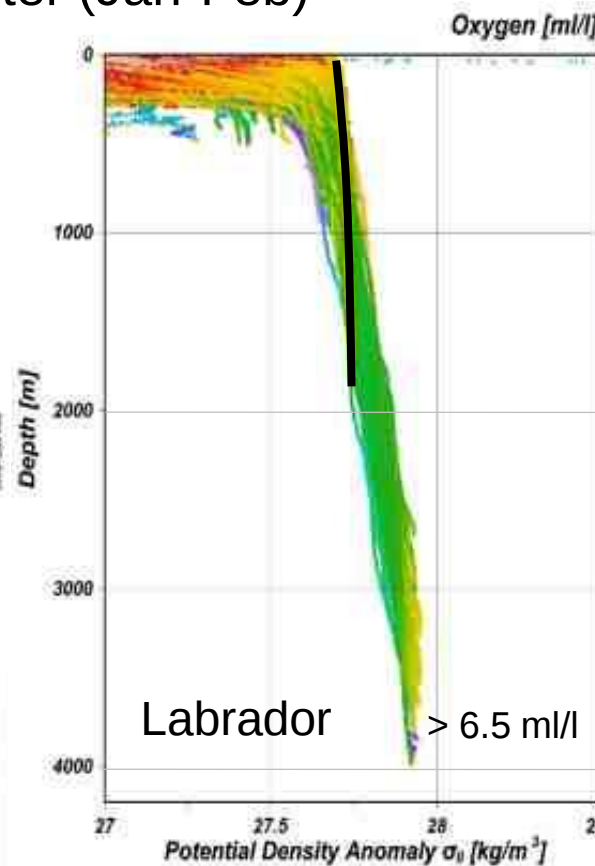
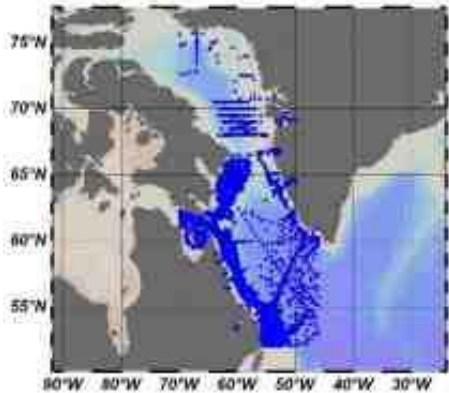
Weakly stratified

Strongly stratified

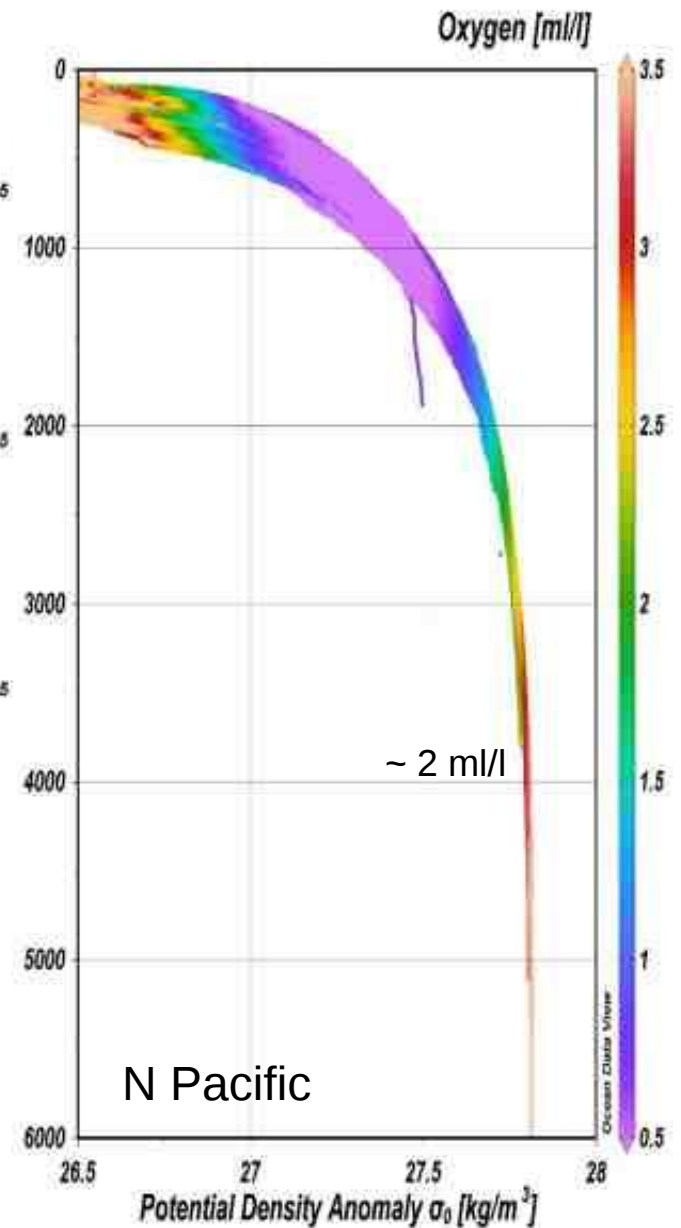
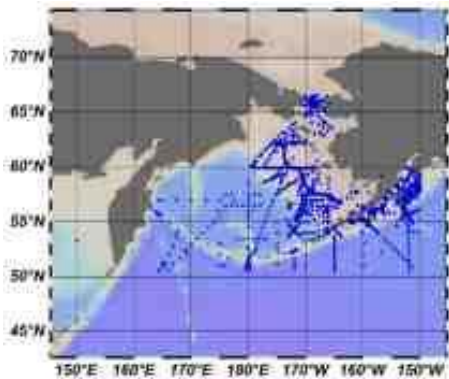
Vertical stratification – observations

σ_θ (kg/m³) winter (Jan-Feb) –

Labrador

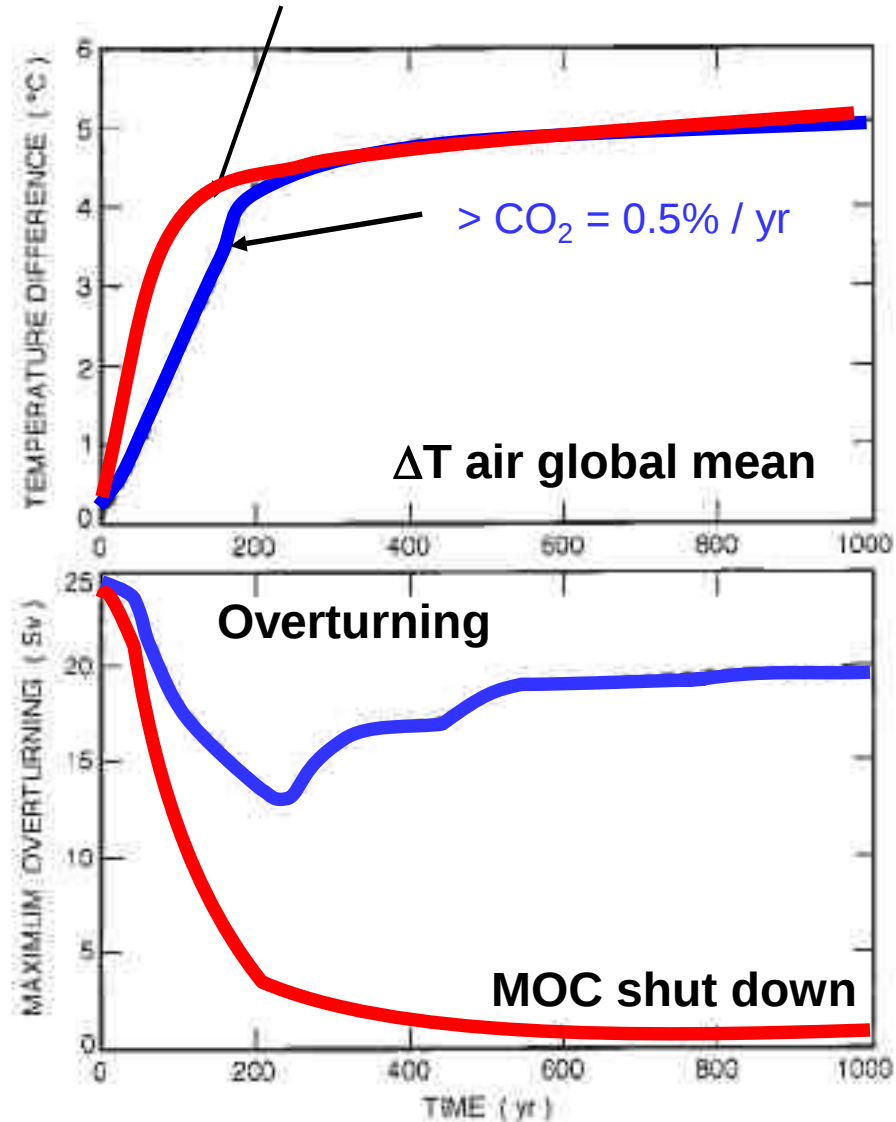


N Pacific



MOC sensitivity based on a simple 3 basin coupled model forced with different rates of CO₂ increase, both to 750ppm

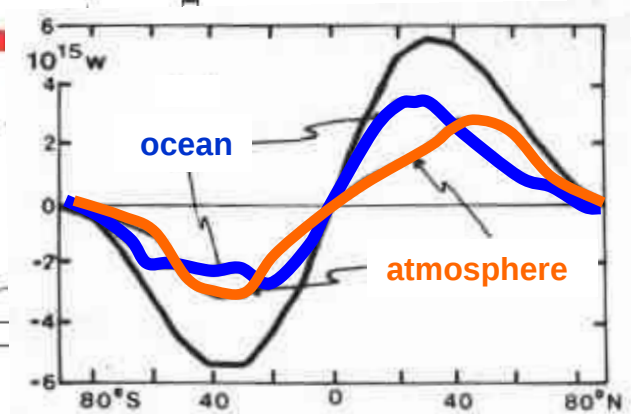
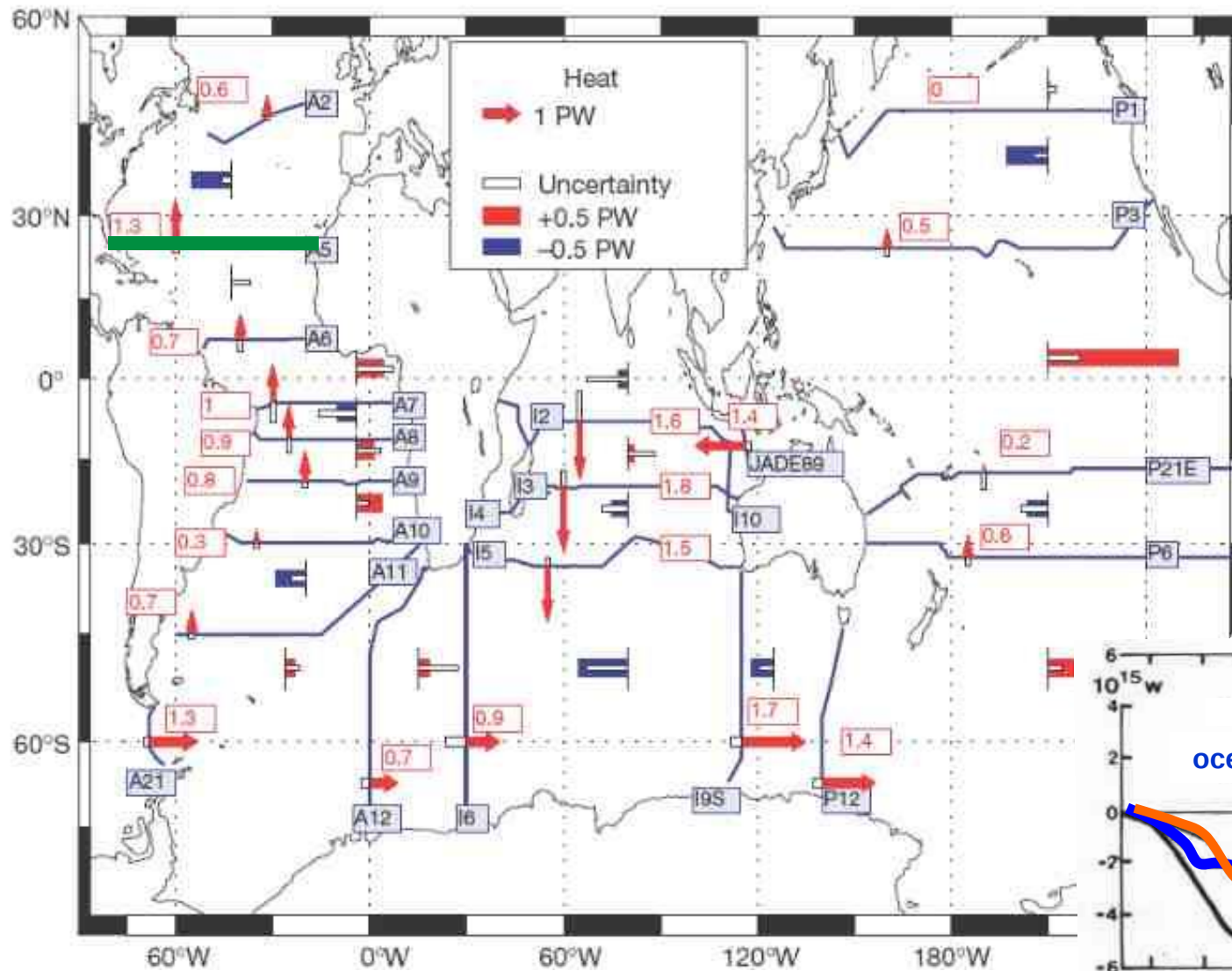
> CO₂ = 1% / yr



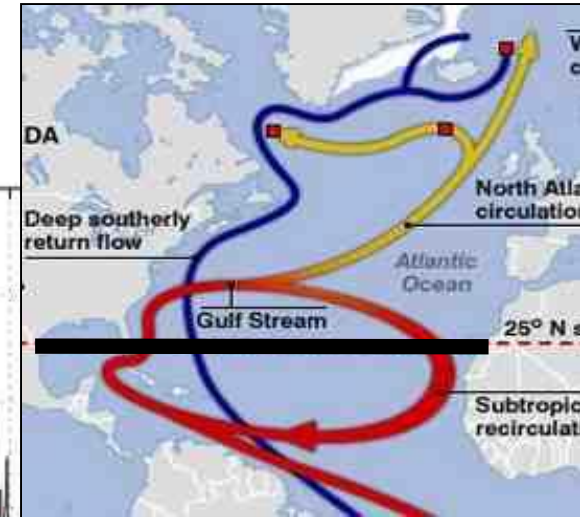
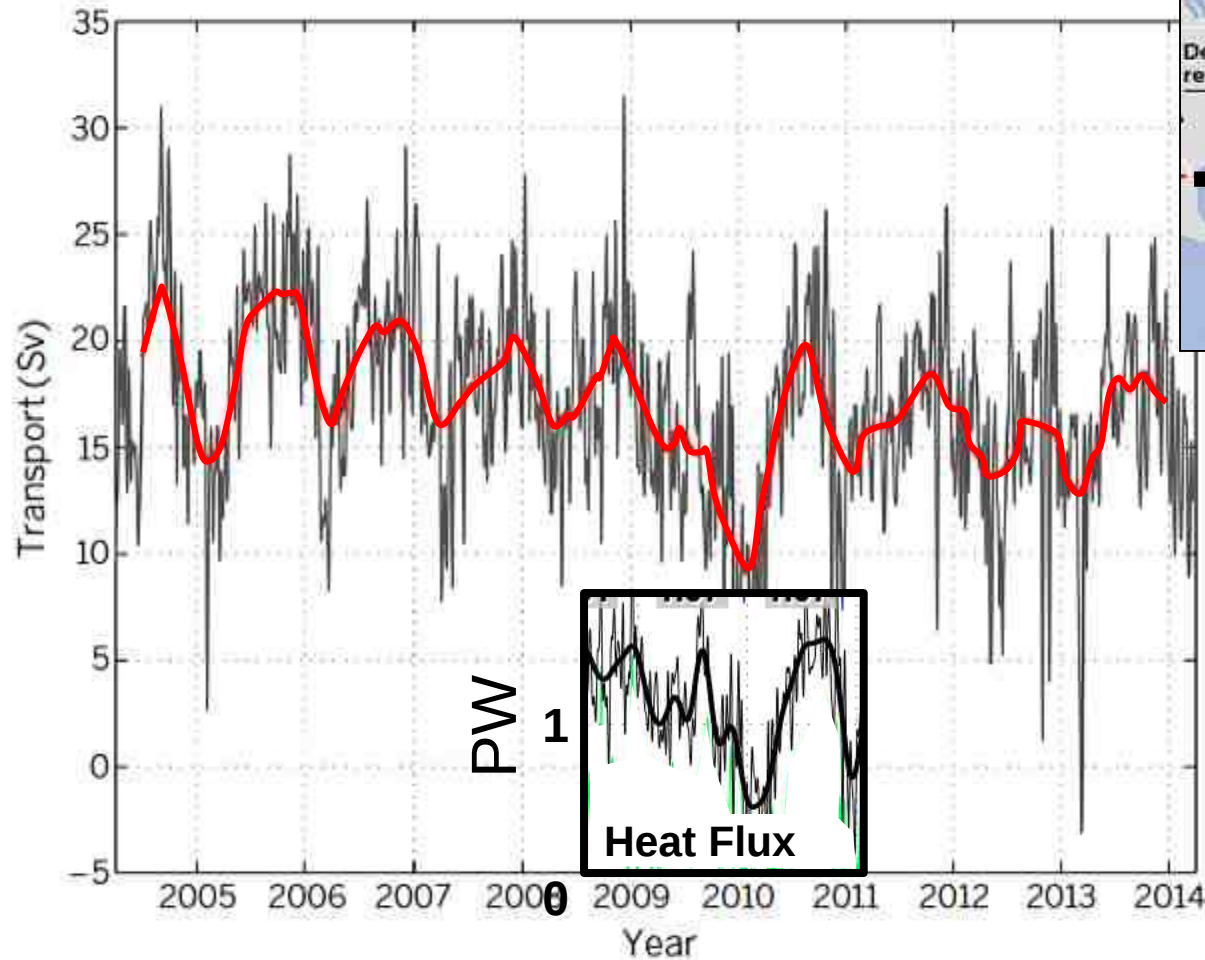
Global ocean meridional heat flux



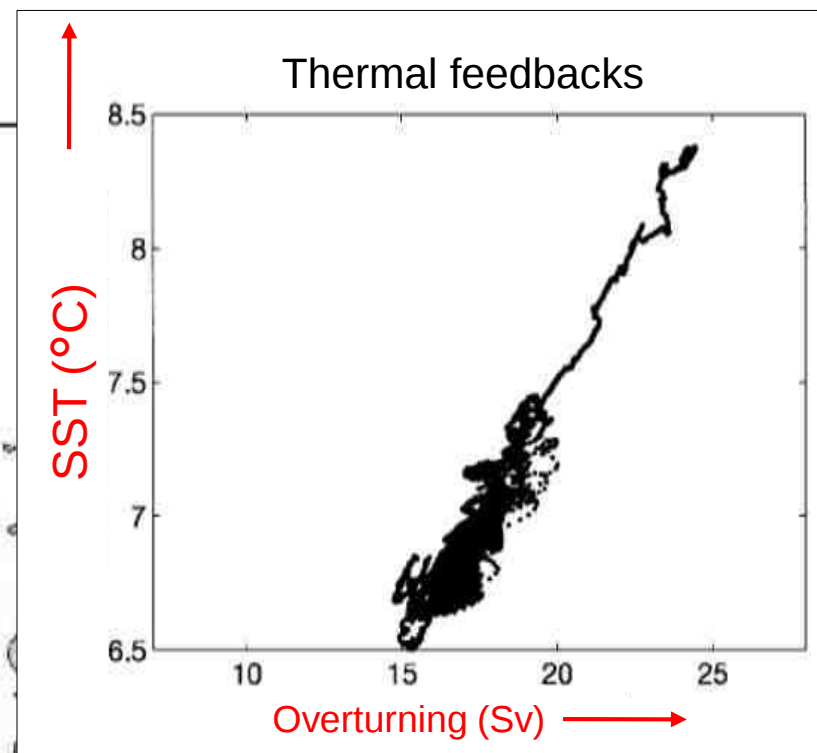
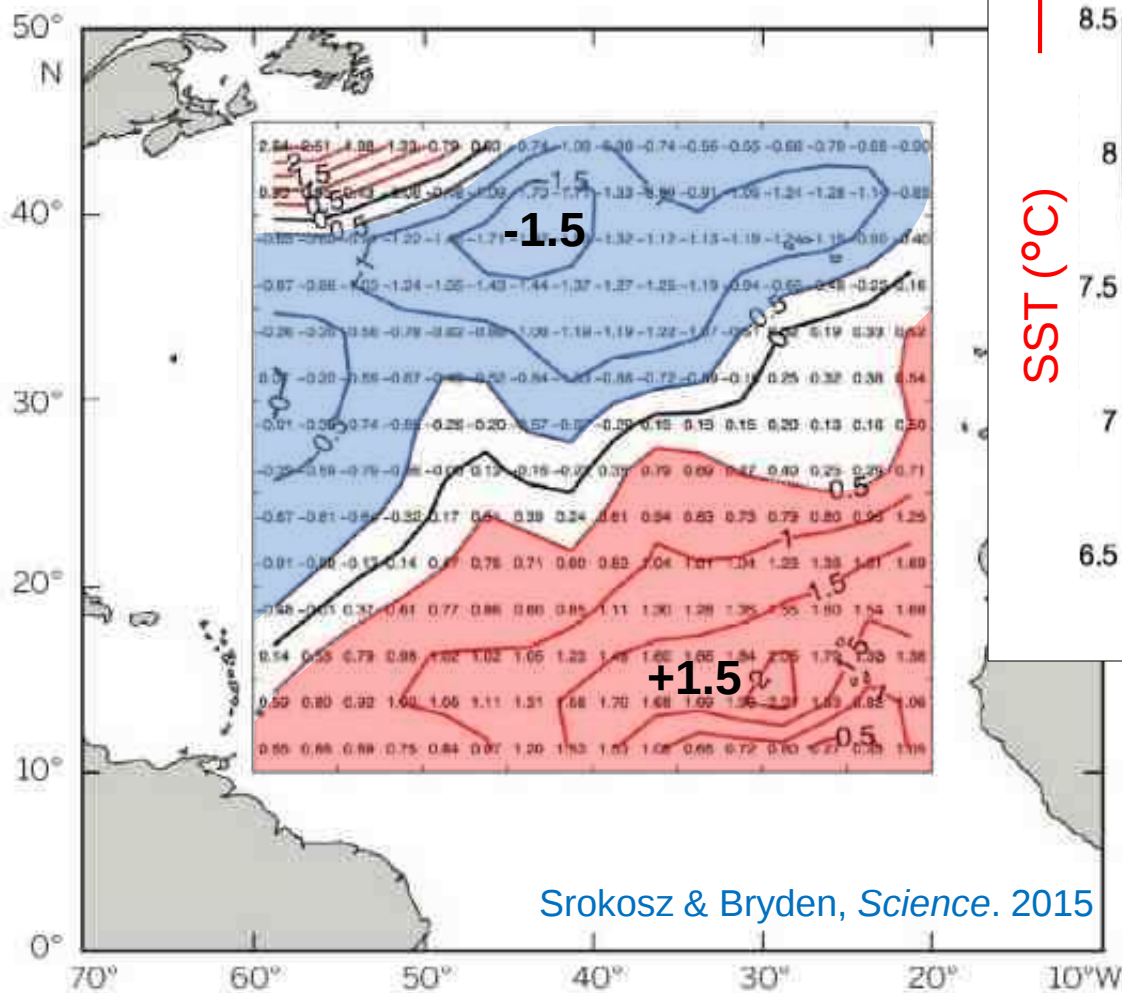
C. Wunsch



Atlantic meridional overturning at 26°N



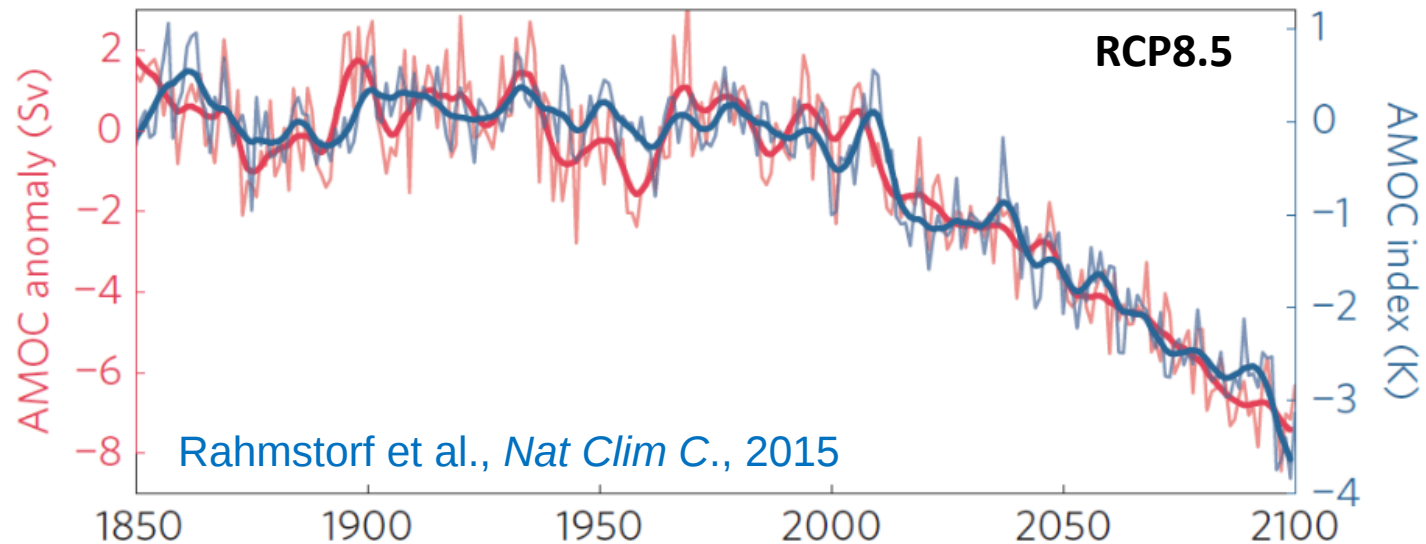
Srokosz & Bryden, *Science*. 2015



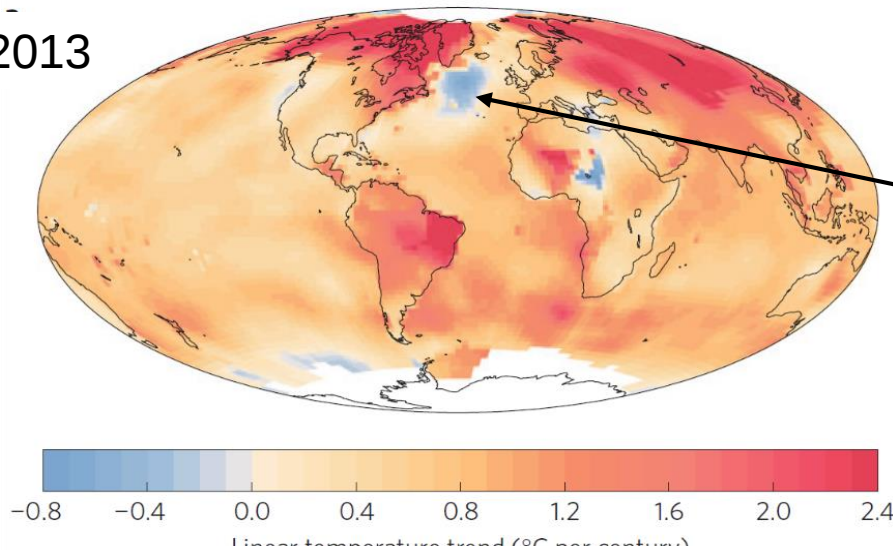
Rahmstorf & England, *JPO*. 1997

Temperature anomaly @ 50-m at the end of the 2009 - 2010 AMOC slowdown event

Subpolar surface temperature based AMOC proxy



1901-2013

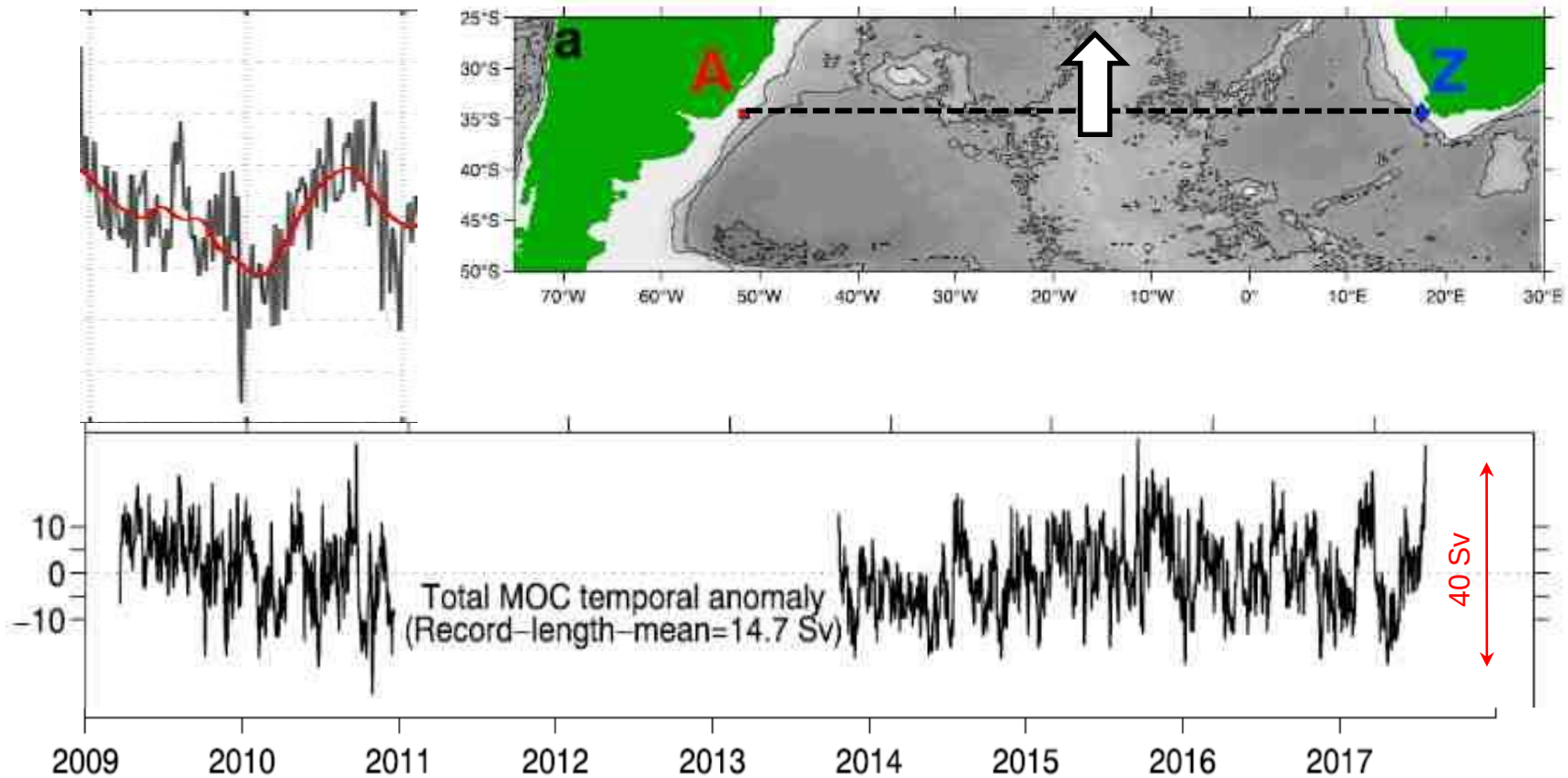


Cooling sub-polar N. Atlantic. Why isn't the AMOC recovering?

Observed SST trend
(°C/century)

Observed Atlantic overturning – 34°S

Total MOC variability at South Atlantic gateway



- The ocean contributes to about 30% of the meridional redistribution of heat on Earth
- Buoyancy fluxes, surface wind stress and vertical mixing sustain a vigorous meridional overturning circulation in the ocean, particularly in the Atlantic basin
- The MOC also contributes to the meridional heat flux, which is largest in the Atlantic and Indian Oceans
- Models and observations show that there are large changes in the intensity of the MOC and the associated heat transport
- Recent warming of the upper ocean may be associated with an acceleration of the hydrological cycle.
- There is strong evidence suggesting that these changes in the freshwater forcing at the surface have lead to changes in water mass properties and circulation, whose impact on the MOC is still poorly understood

¡GRACIAS!

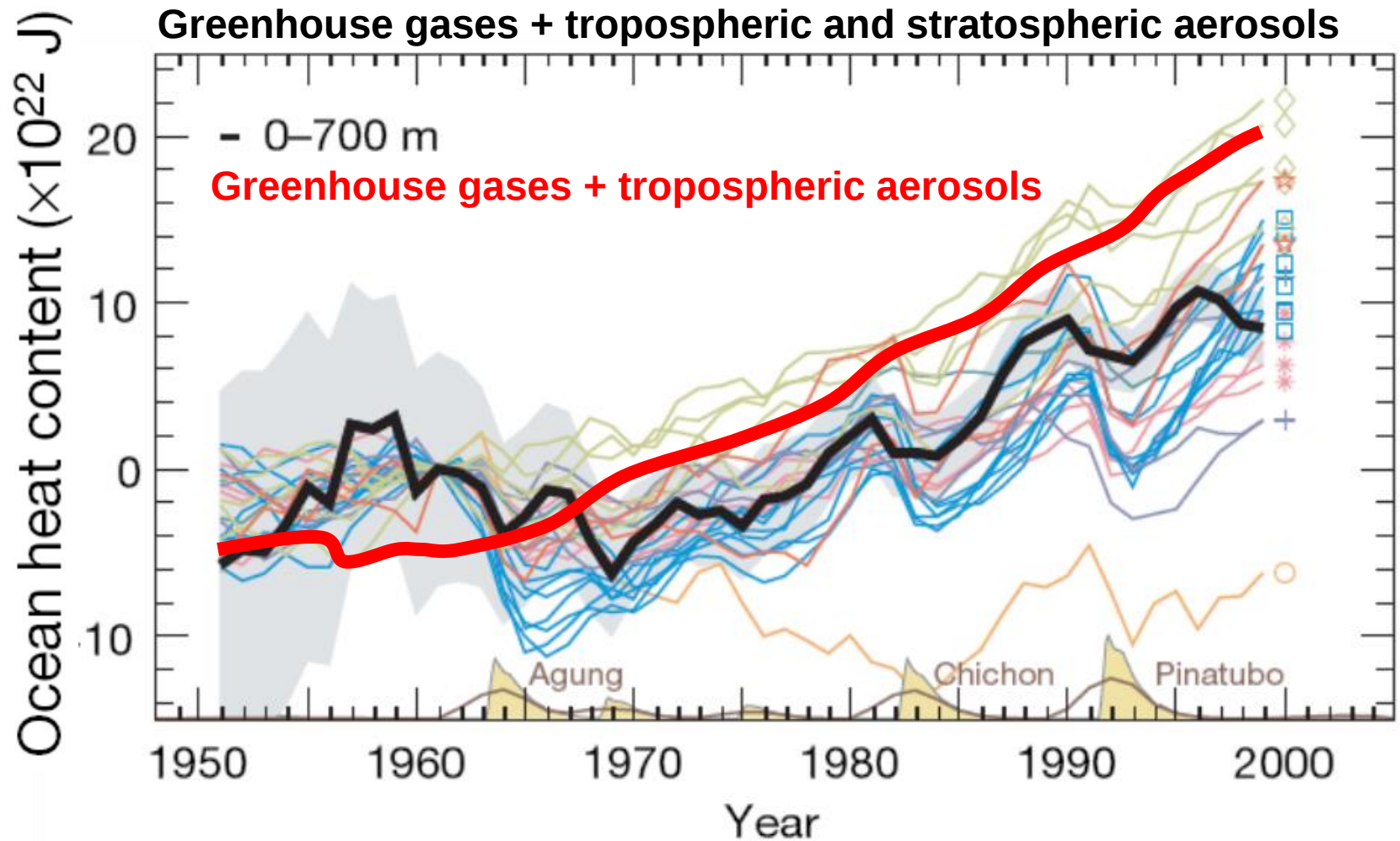


**Obrigada,
Thank you!**

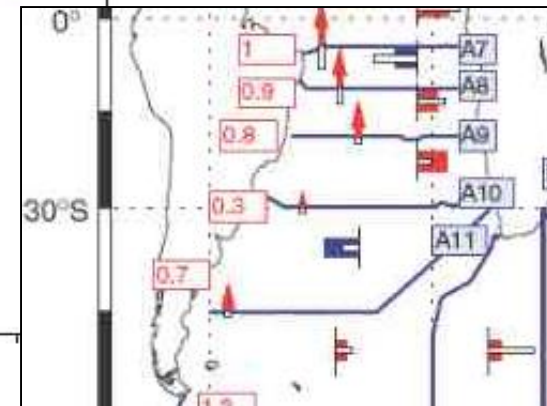
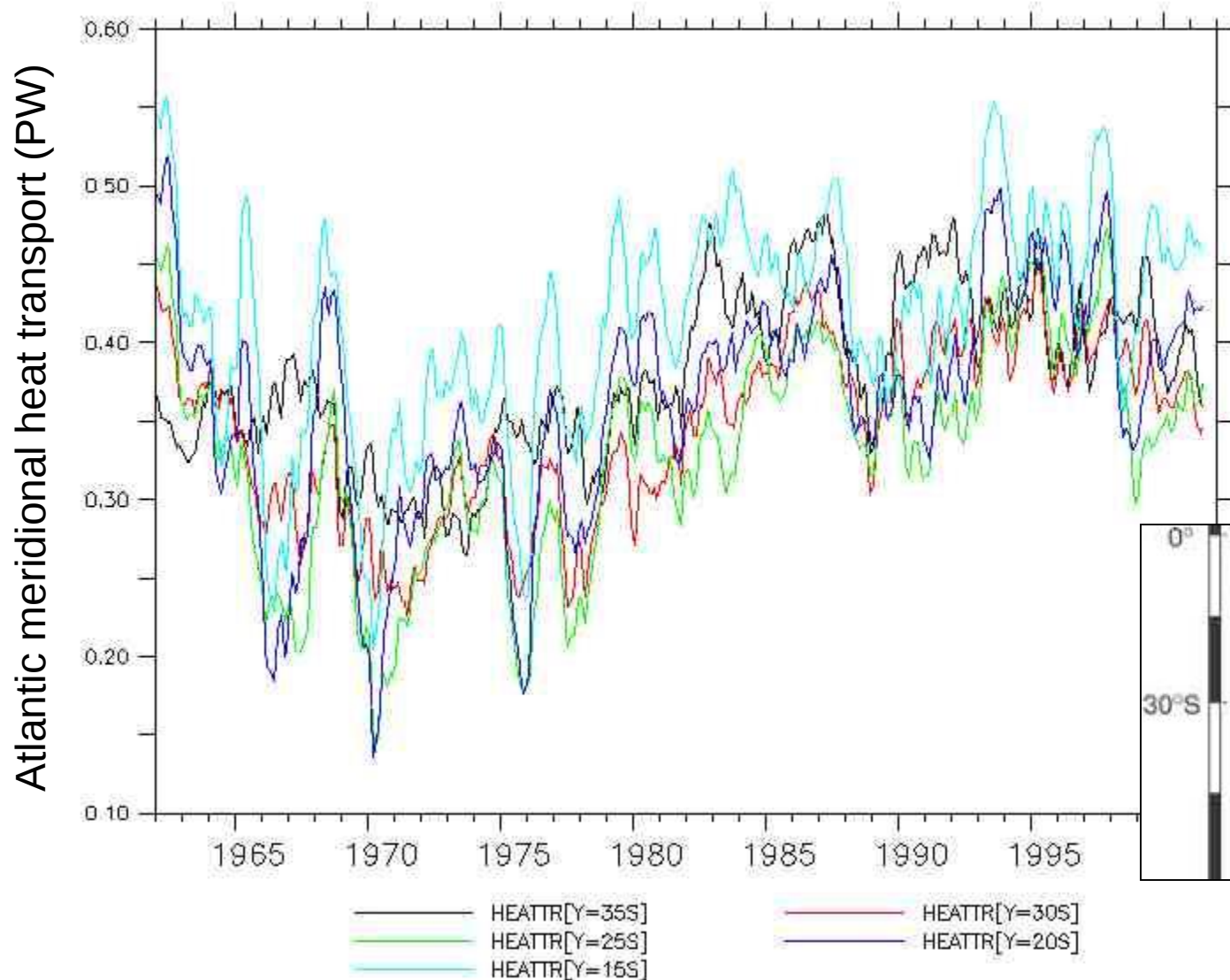


Supplementary Material

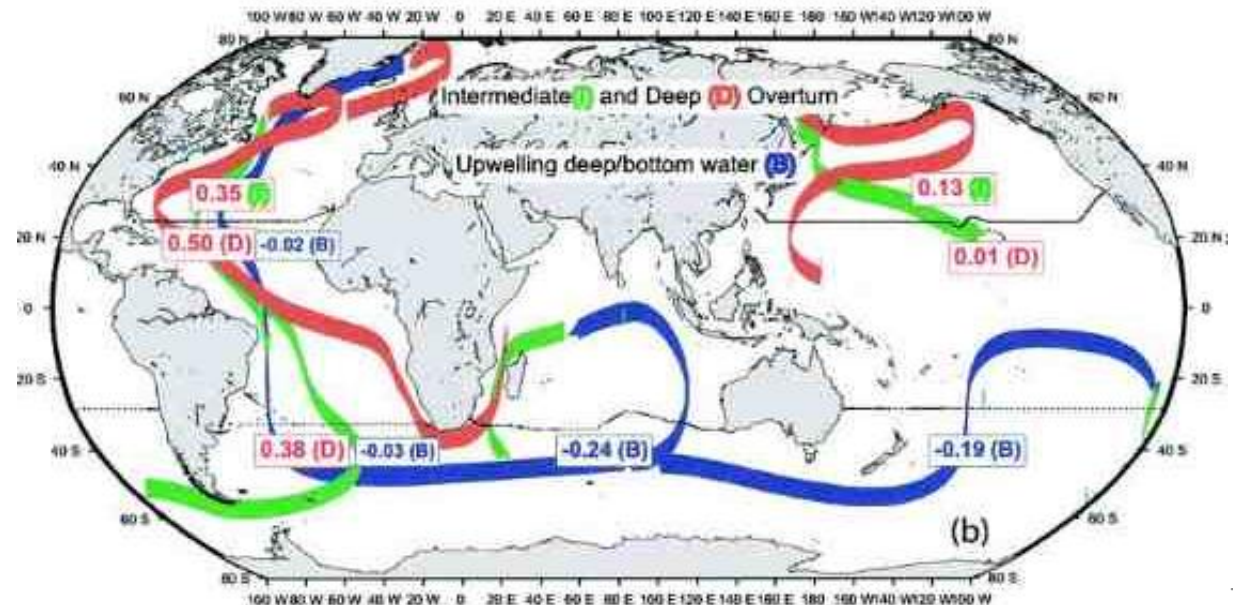
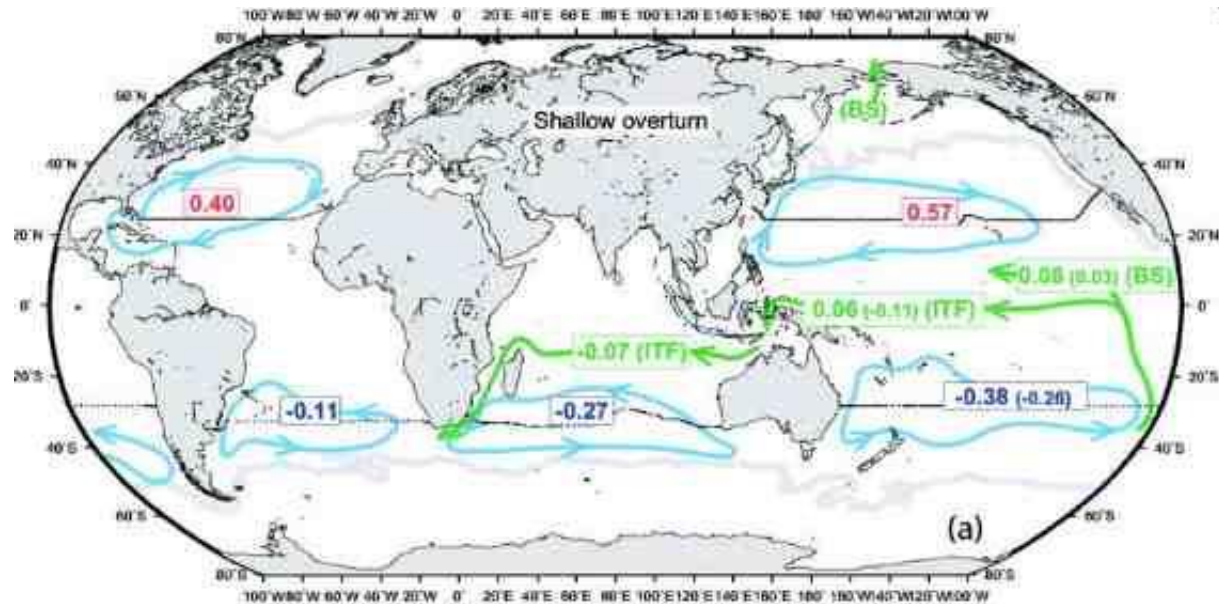
Warming trends in the upper ocean



South Atlantic meridional heat transport (gECCO)

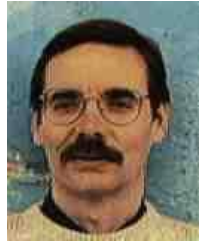
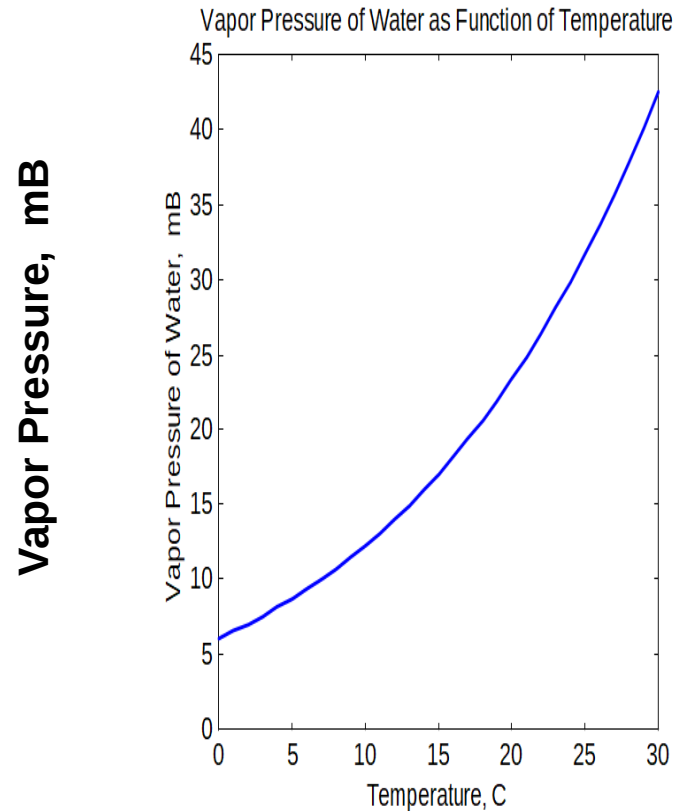


Ocean heat flux mechanisms



indirect effects of warmer atmosphere

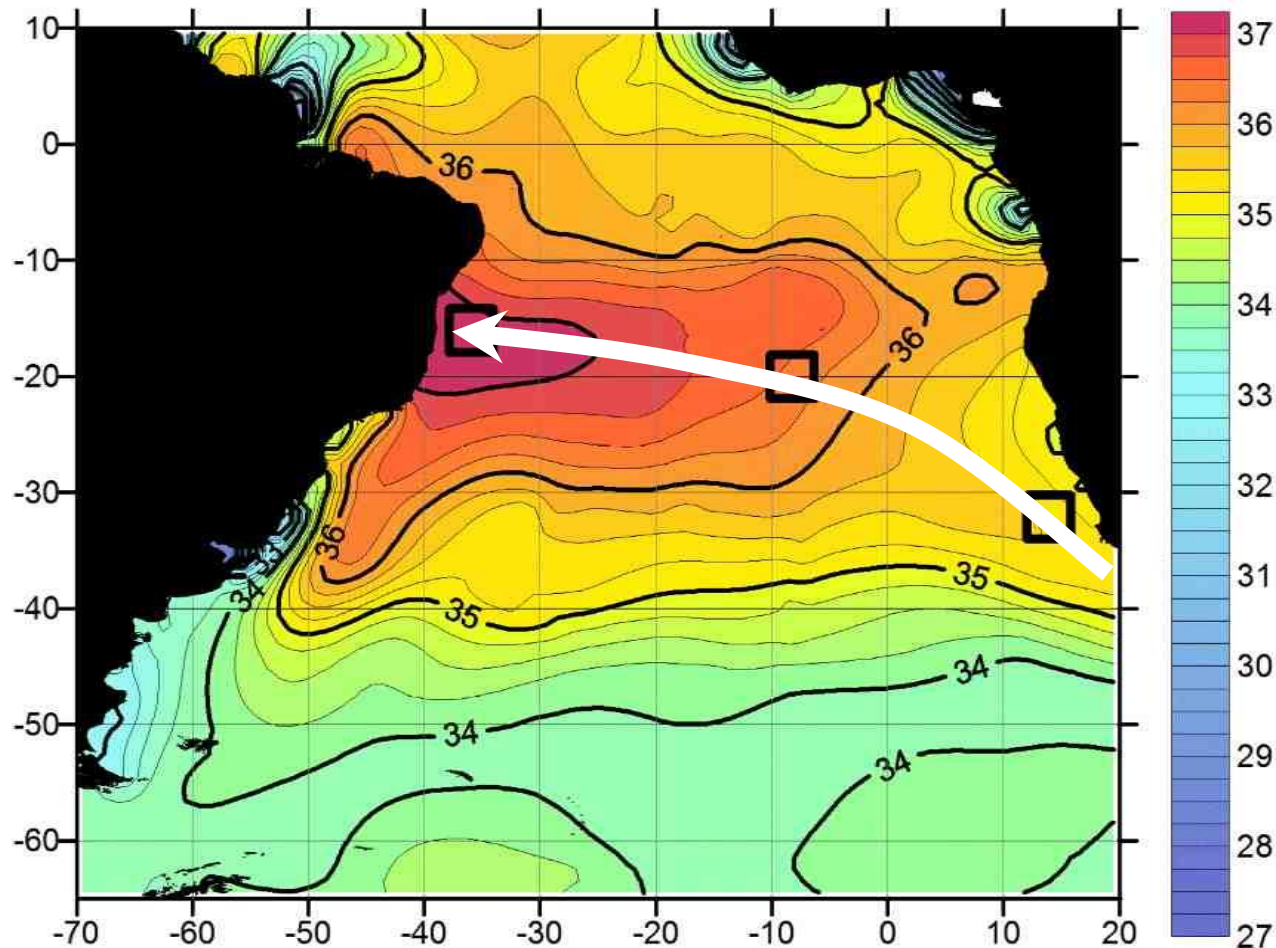
- A warmer atmosphere will carry more water vapor, because of the exponential increase of vapor pressure with temperature.
- An enhanced water cycle will change the distribution of salinity in the upper ocean.



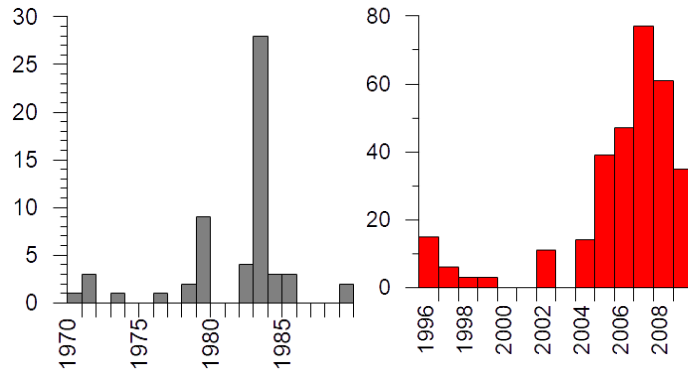
R. Schmitt

Local Upper Ocean Modifications in the South Atlantic

The South Atlantic is the entry point for the upper layer flow into the North Atlantic, required to compensate the export of deep water to the global ocean.

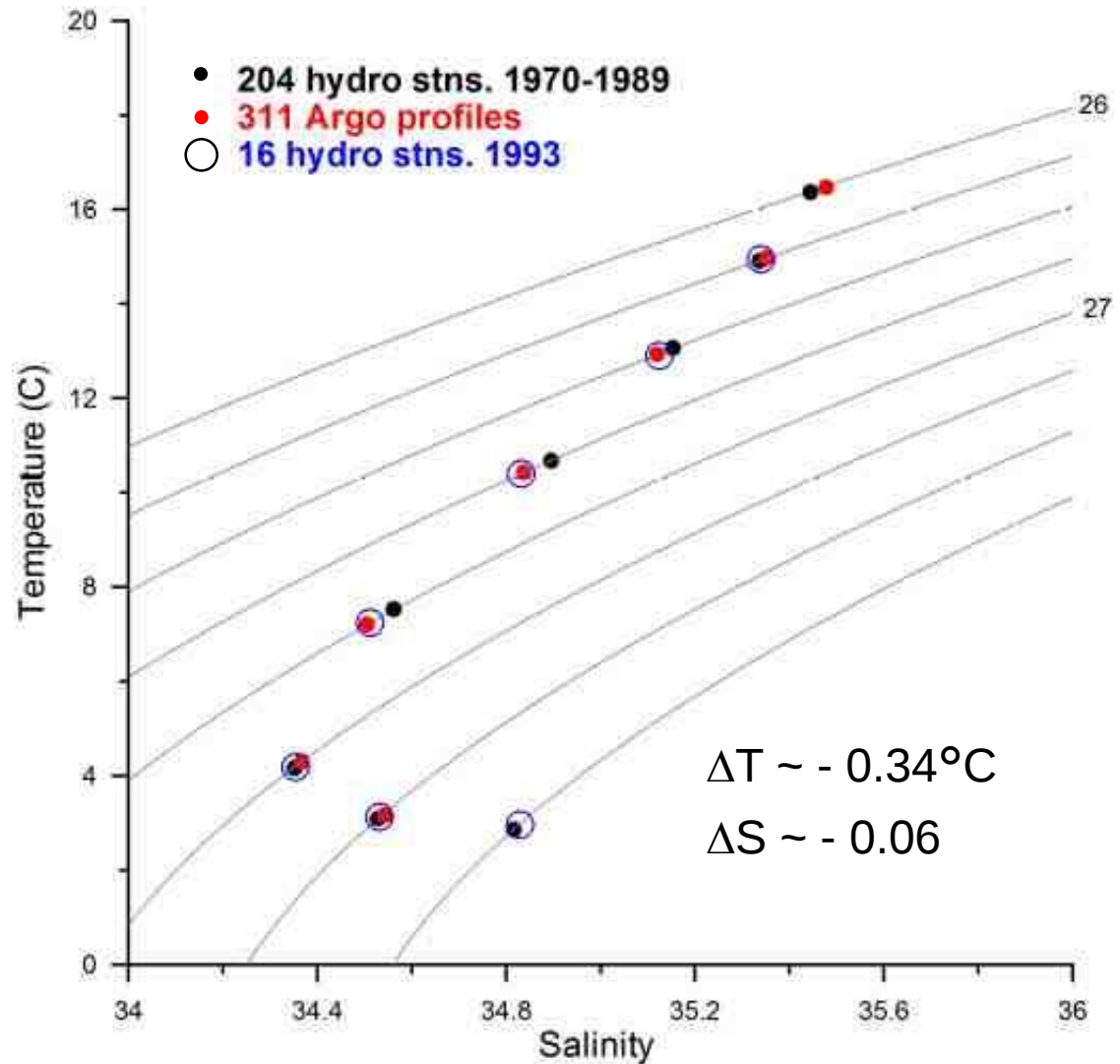


T - S changes on isopycnals - Benguela

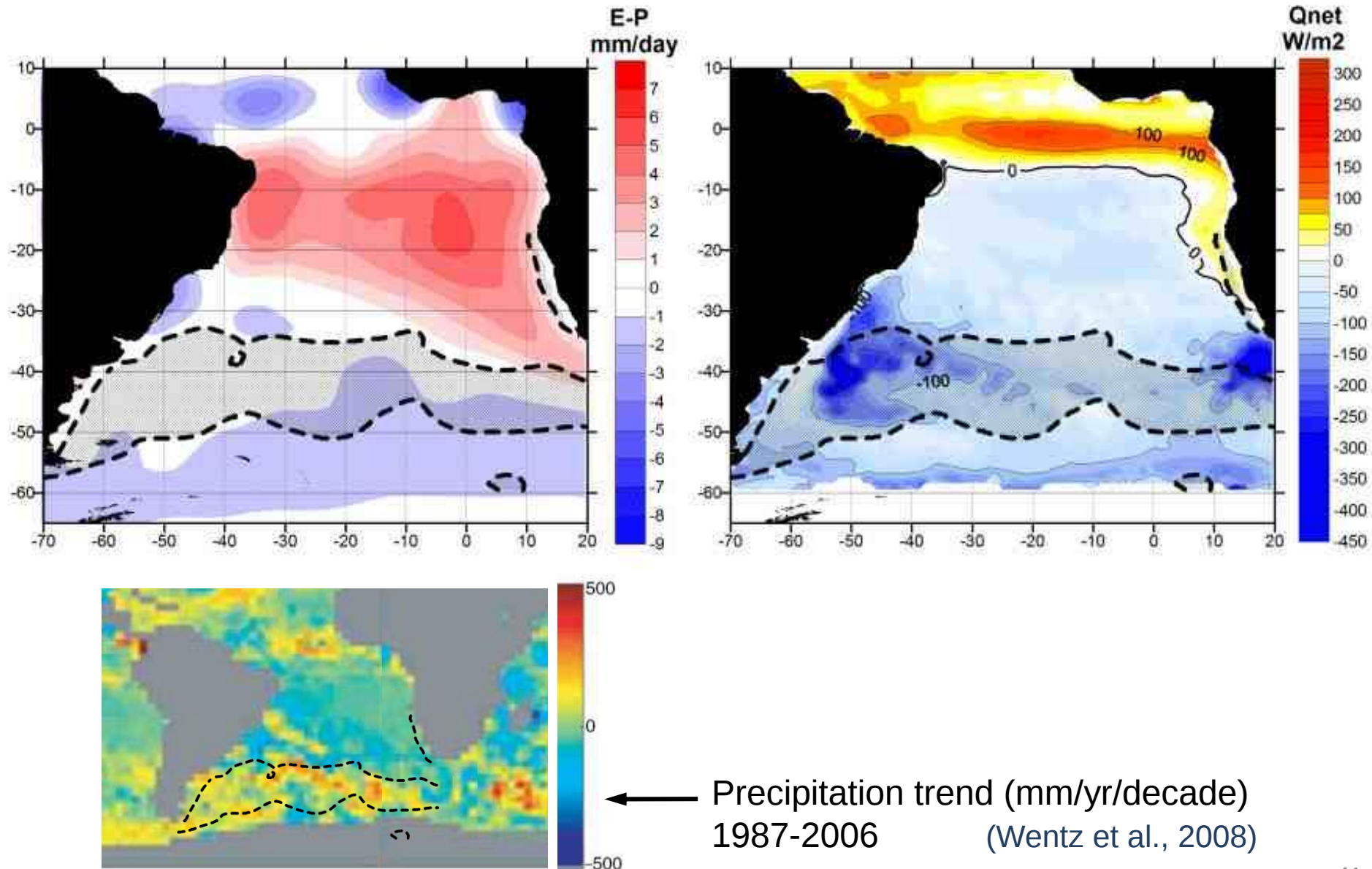


SEC: no significant changes are observed

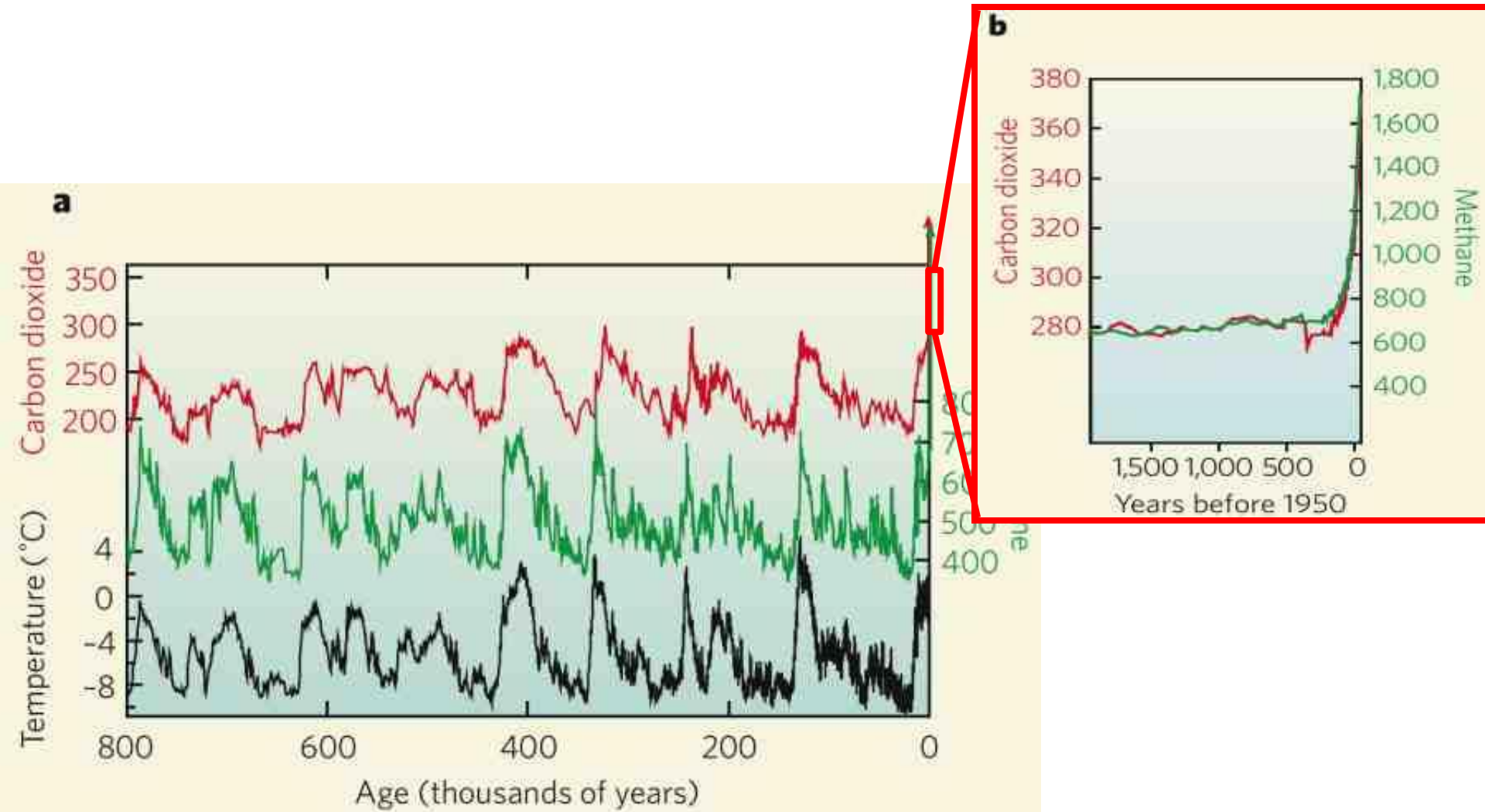
Bifurcation: changes similar to Benguela, but noisier hydro data.



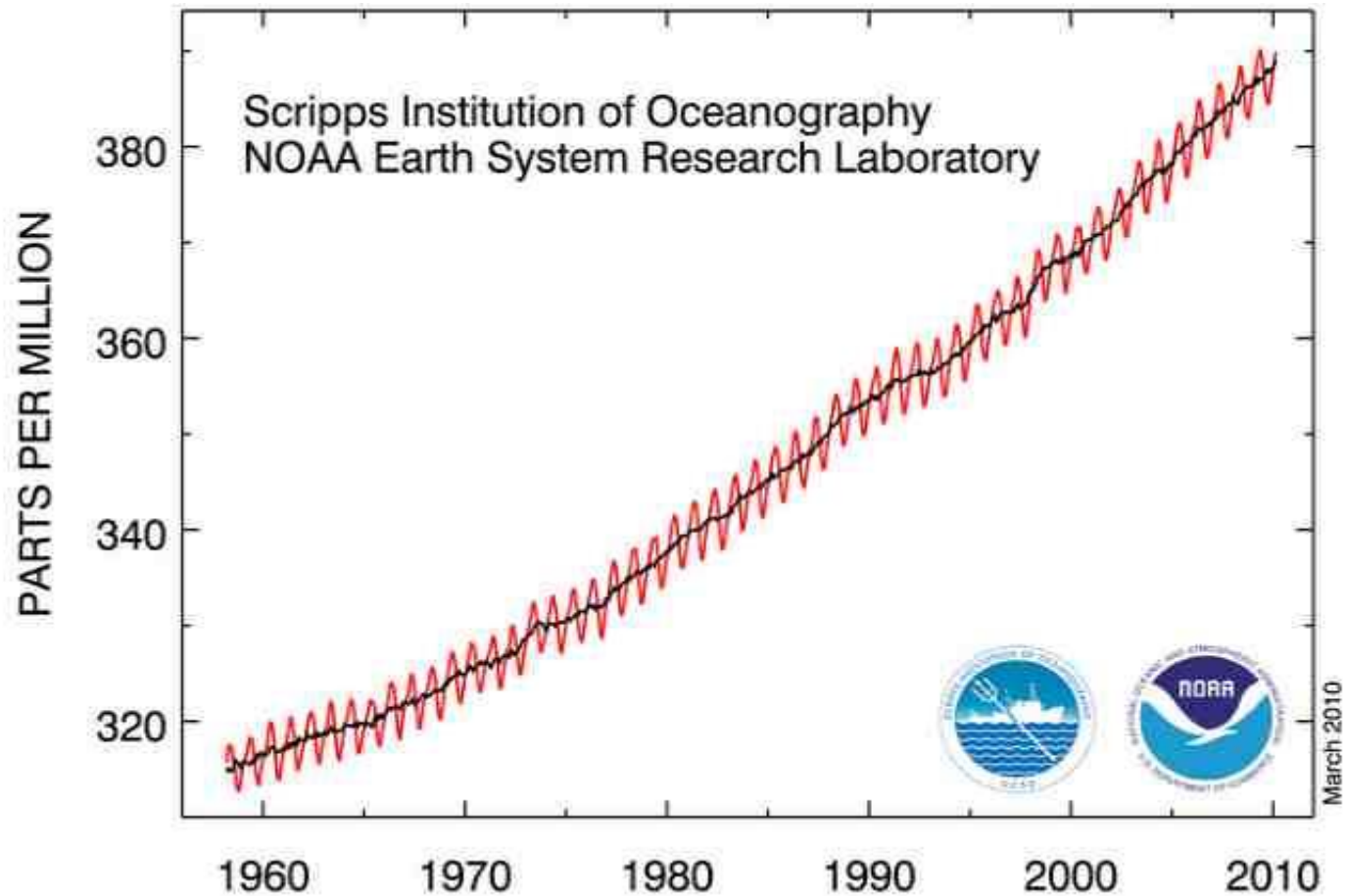
The outcrop regions



Atmospheric CO₂ and global temperature records



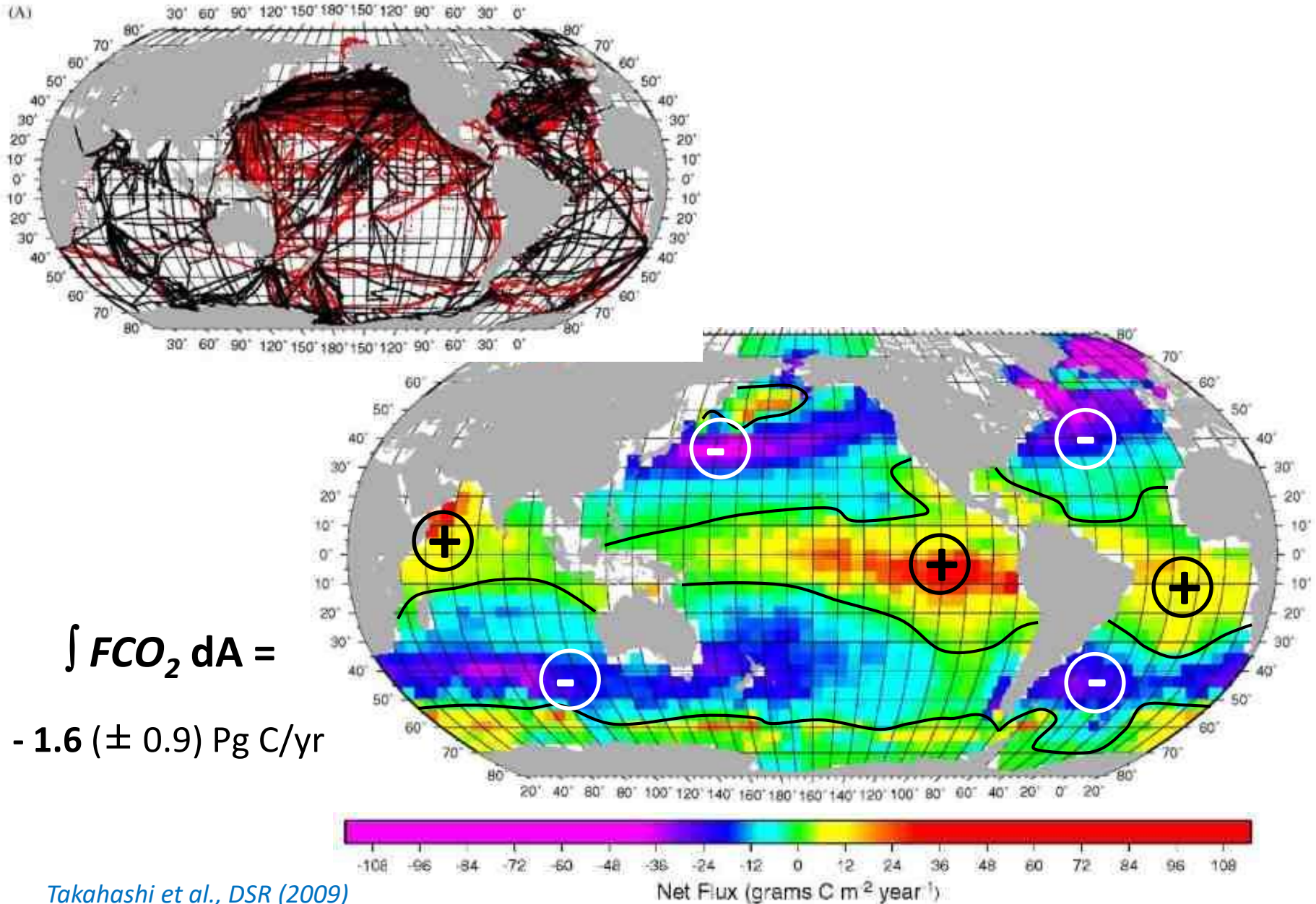
Mean global temperature (black) and atmospheric CO₂ and CH₄ (red and green) during the last 800,000 yrs. Atmospheric CO₂ concentrations increased by 30% during the industrial period.



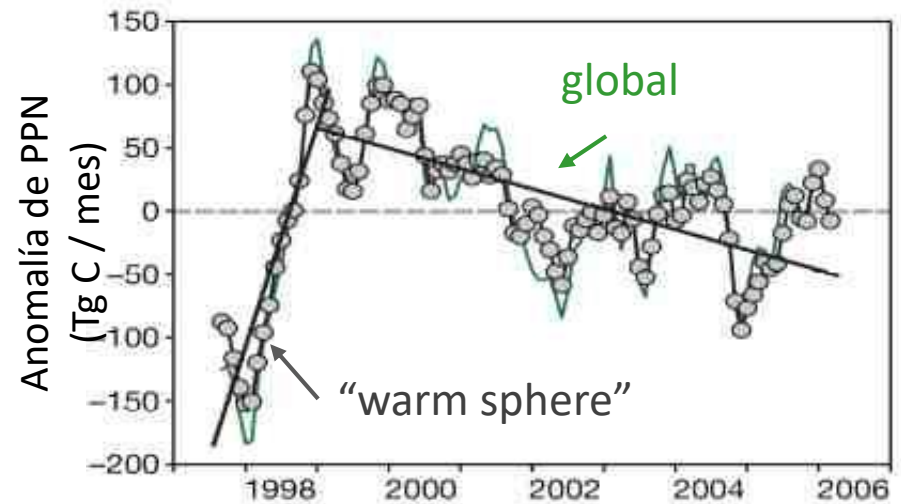
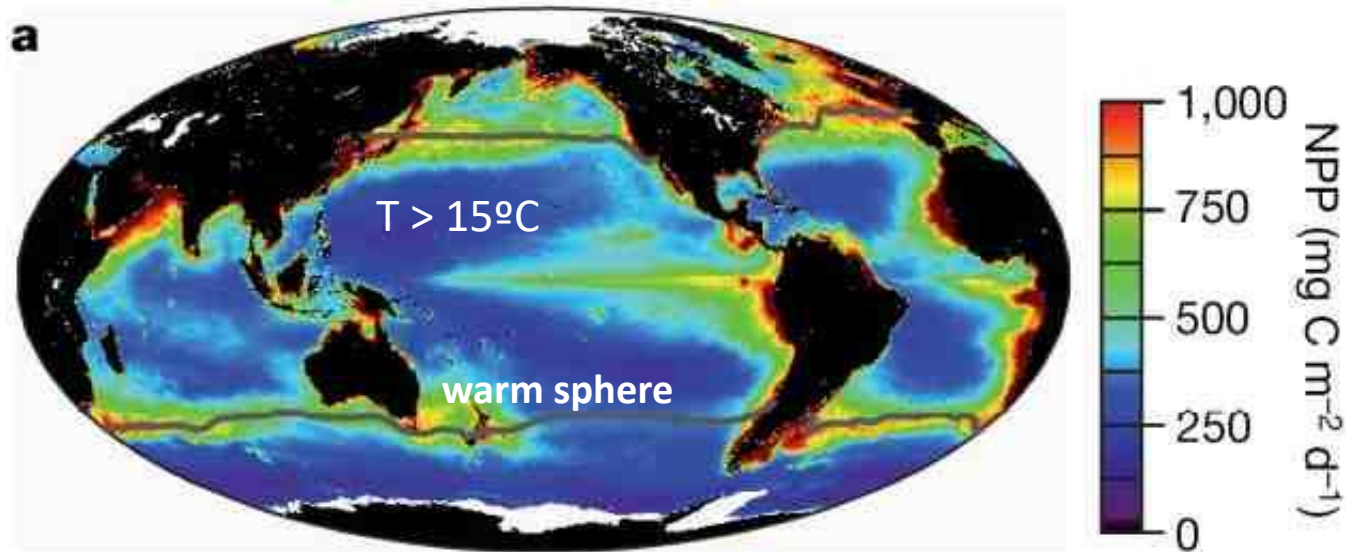
Atmospheric CO₂ data from the Mauna Loa observatory, Hawaii since 1959

<http://www.esrl.noaa.gov/gmd/ccgg/trends/>

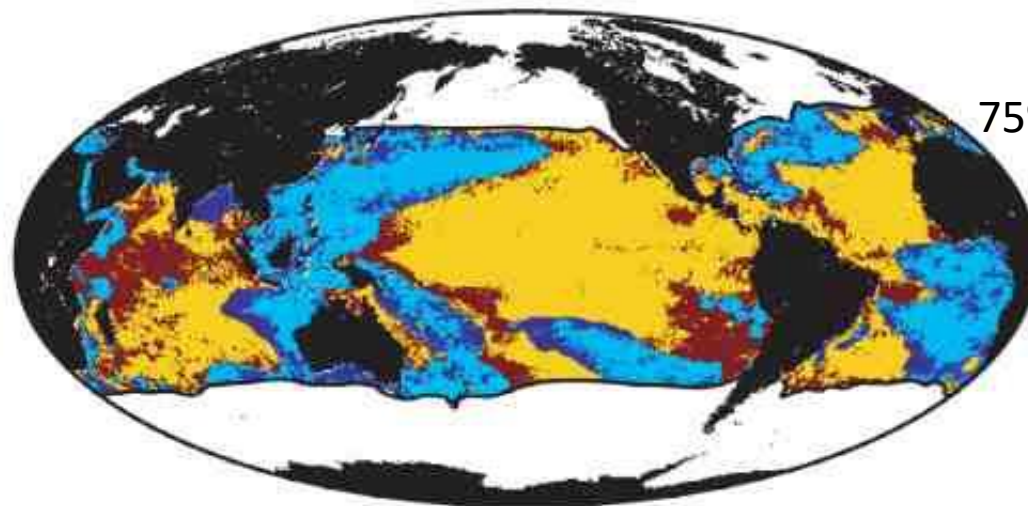
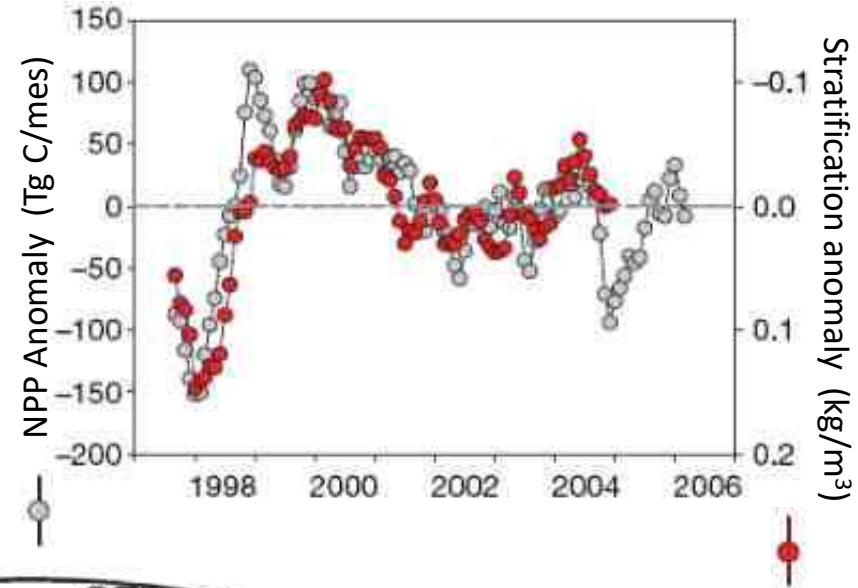
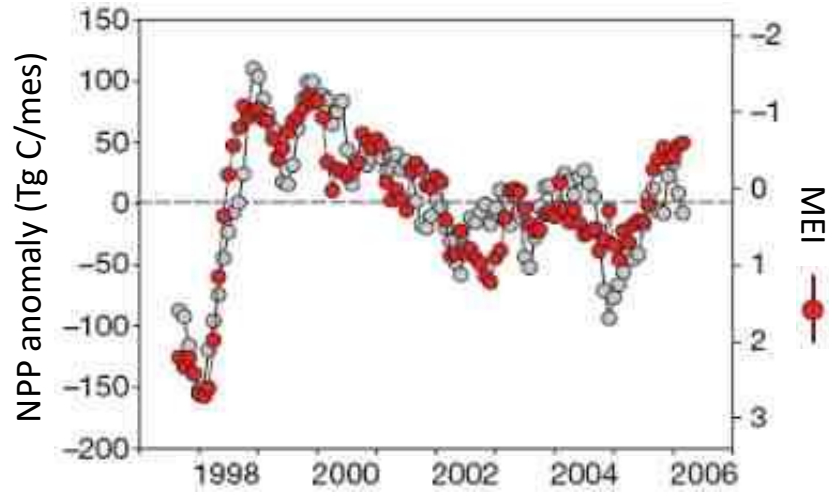
Impacts: absorption of atmospheric CO₂



Net primary production: Lessons from a strong Niño

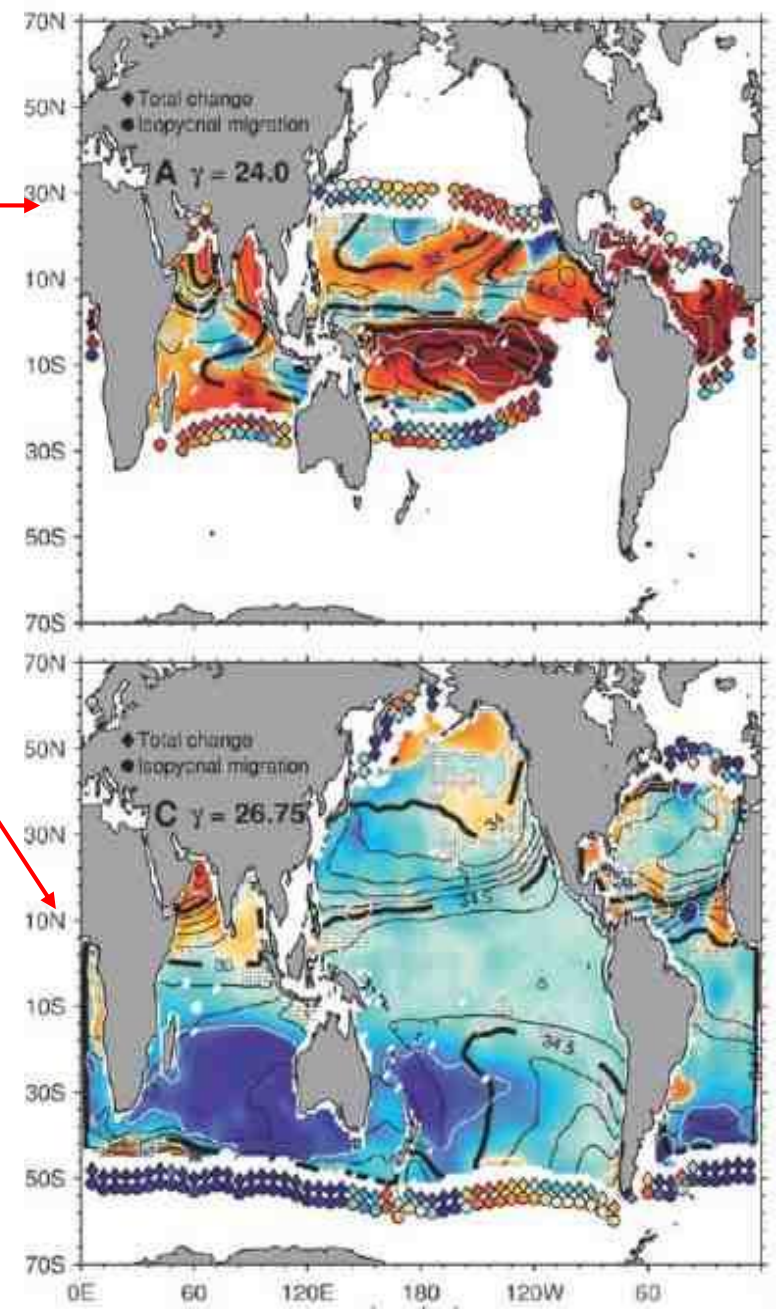
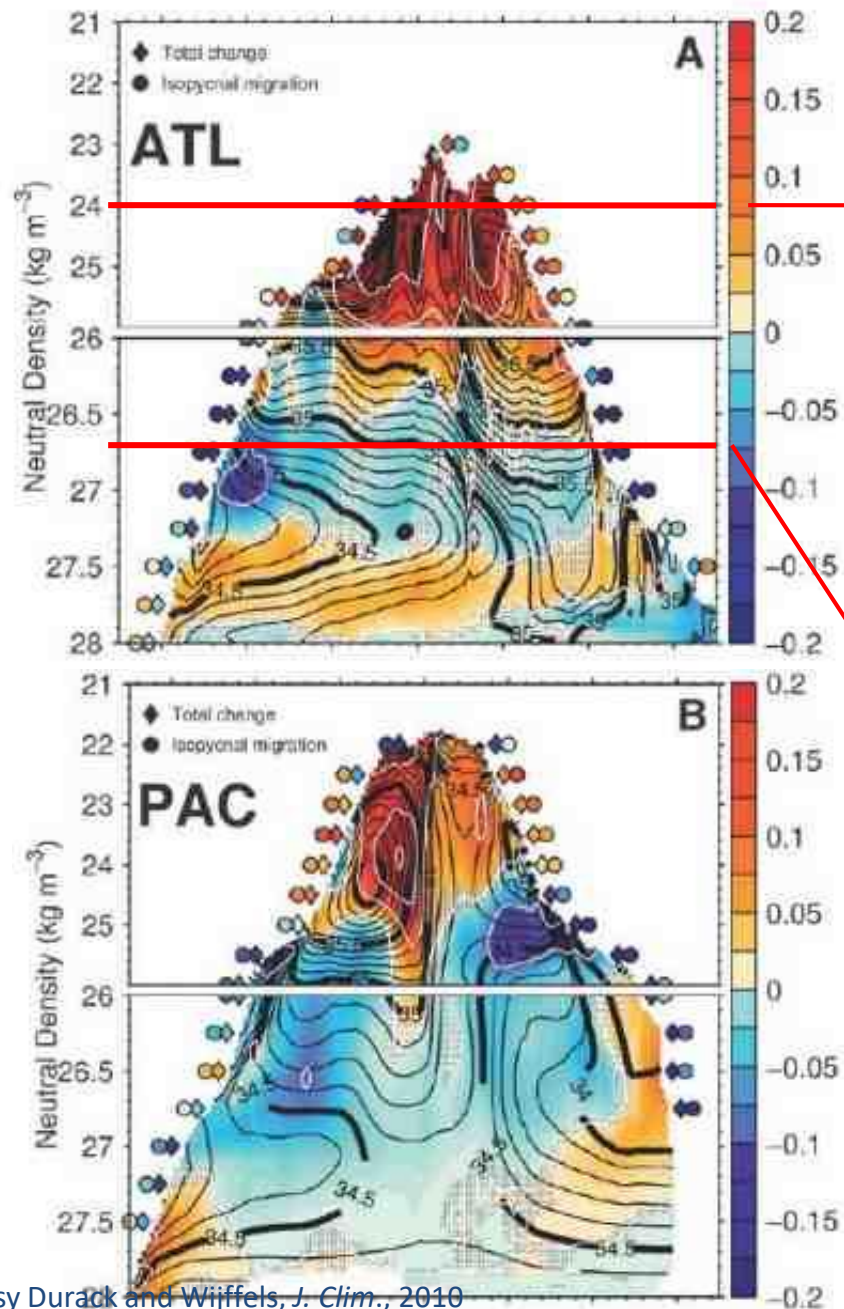


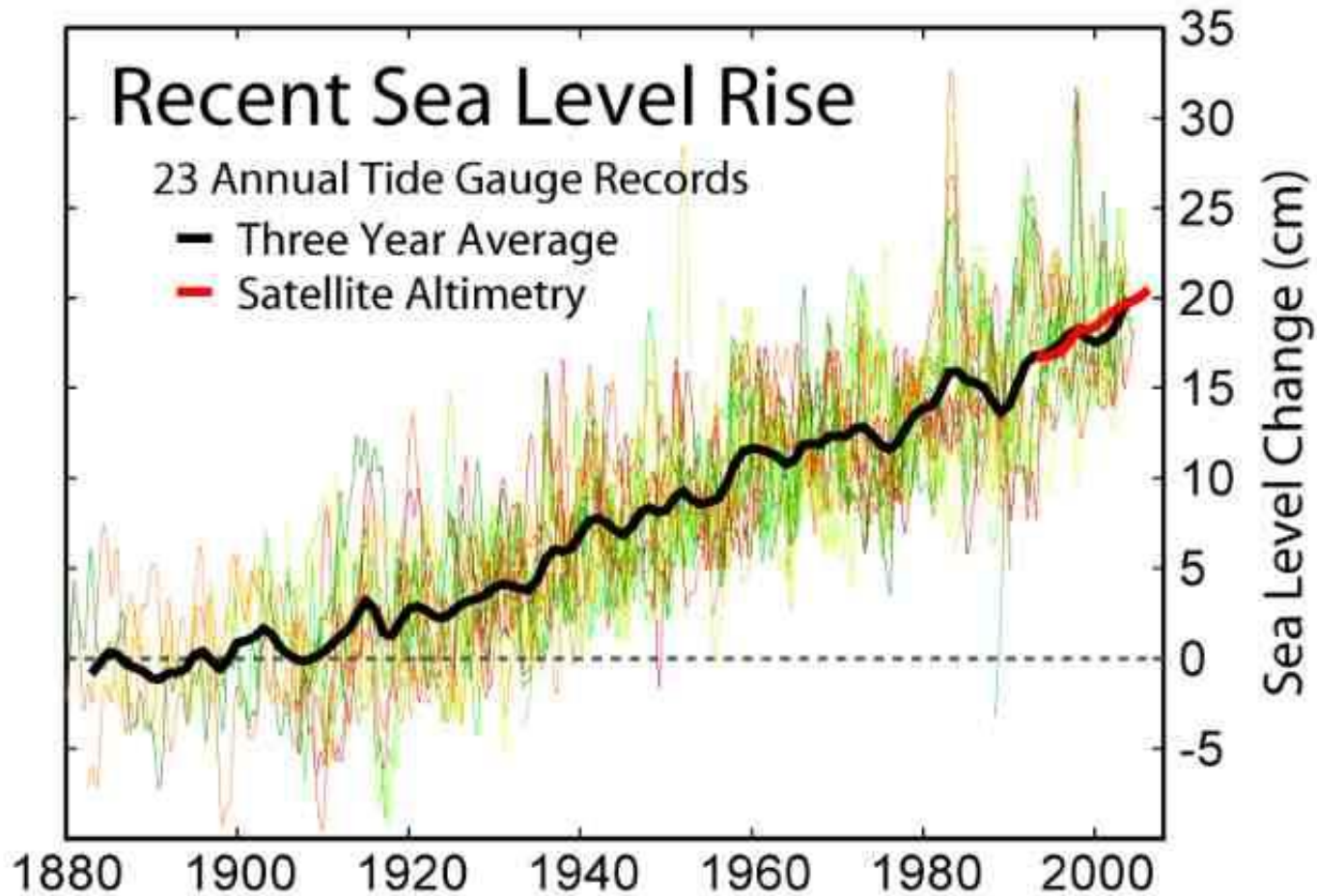
NPP variability — warm sphere, $T > 15^{\circ}\text{C}$



75% ocean area

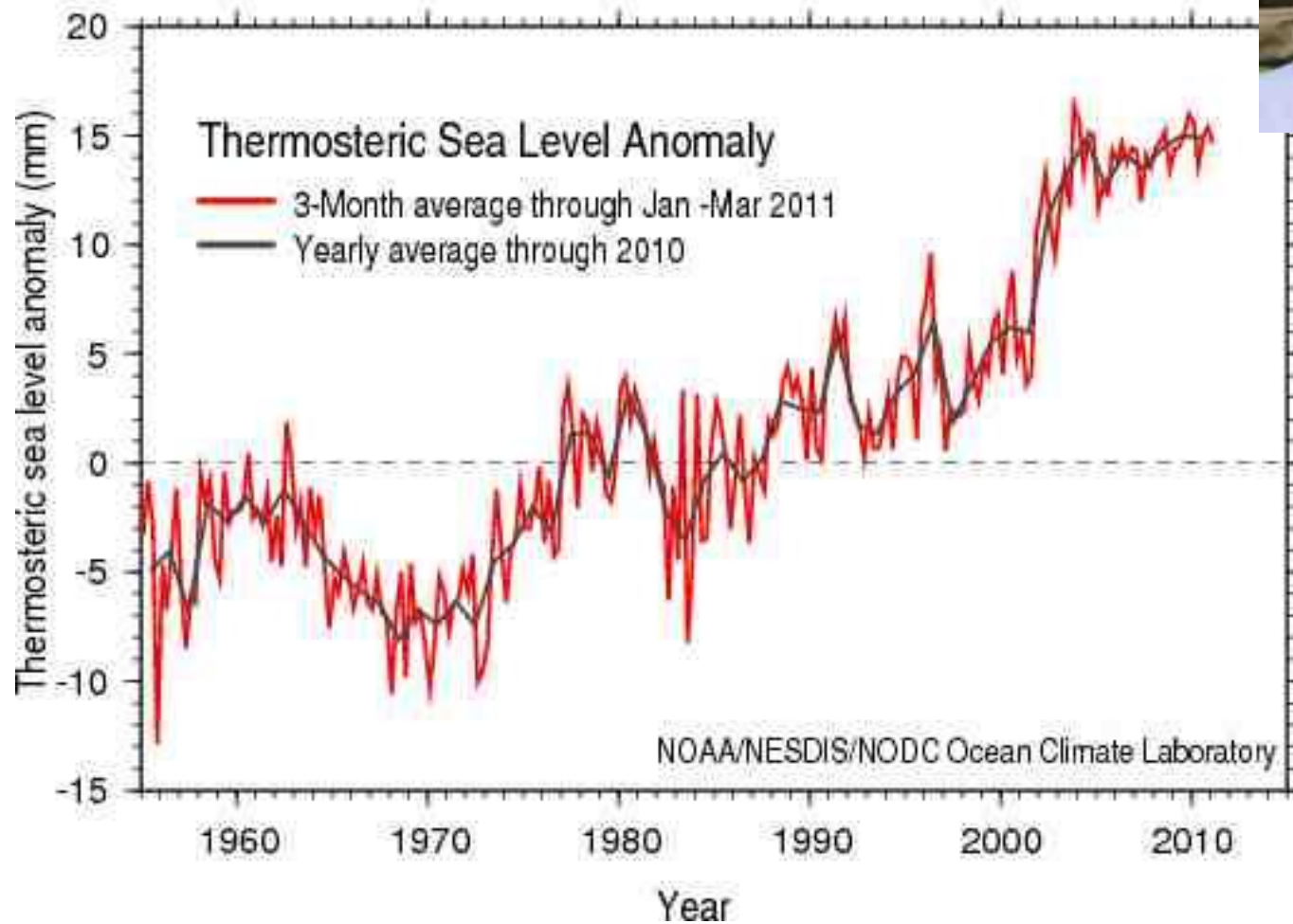
Global salinity changes - IV



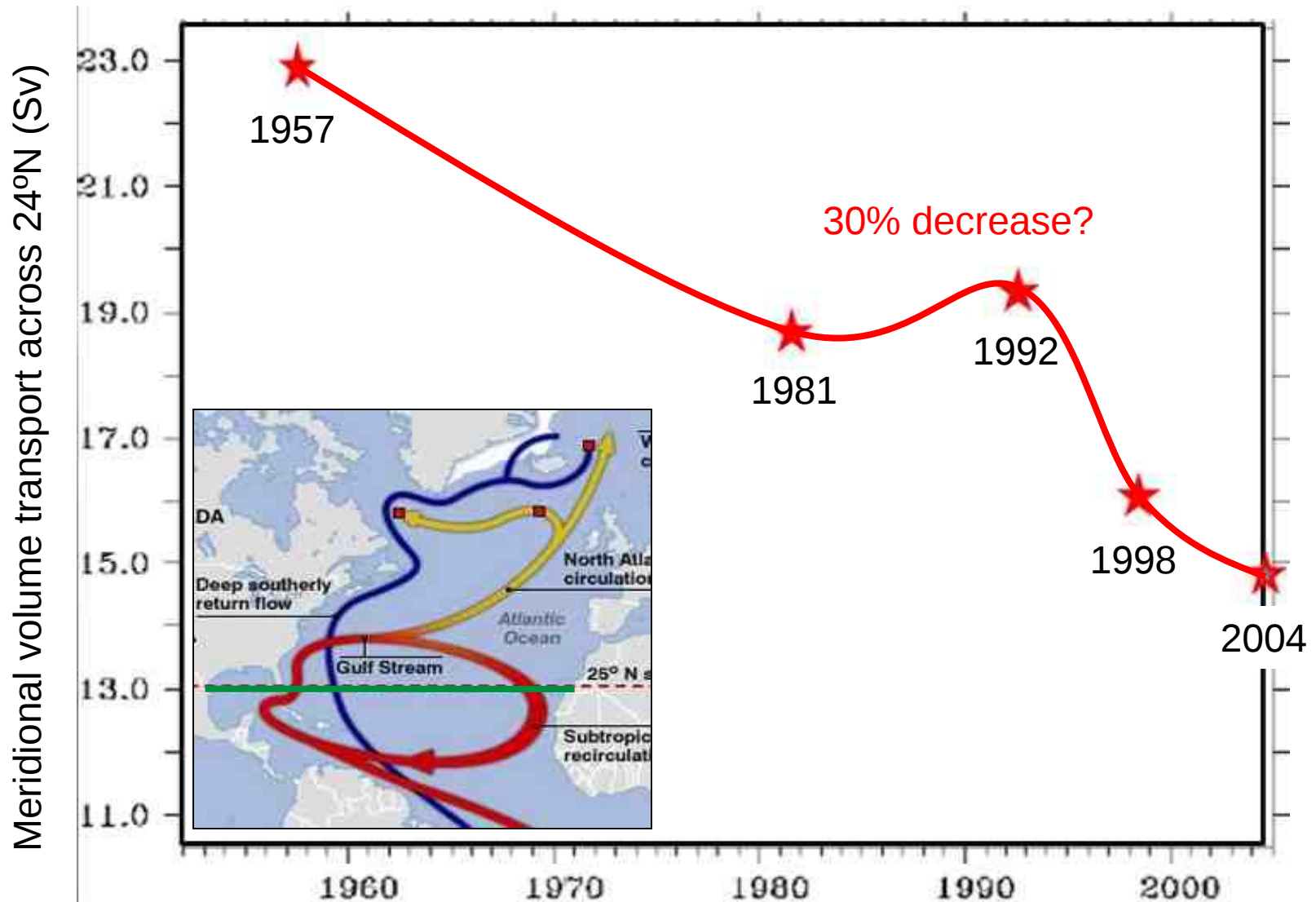


Changes in mean sea level based on 23 tide long-term gauge records (black), and satellite altimeters (red).

Thermosteric sea level rise

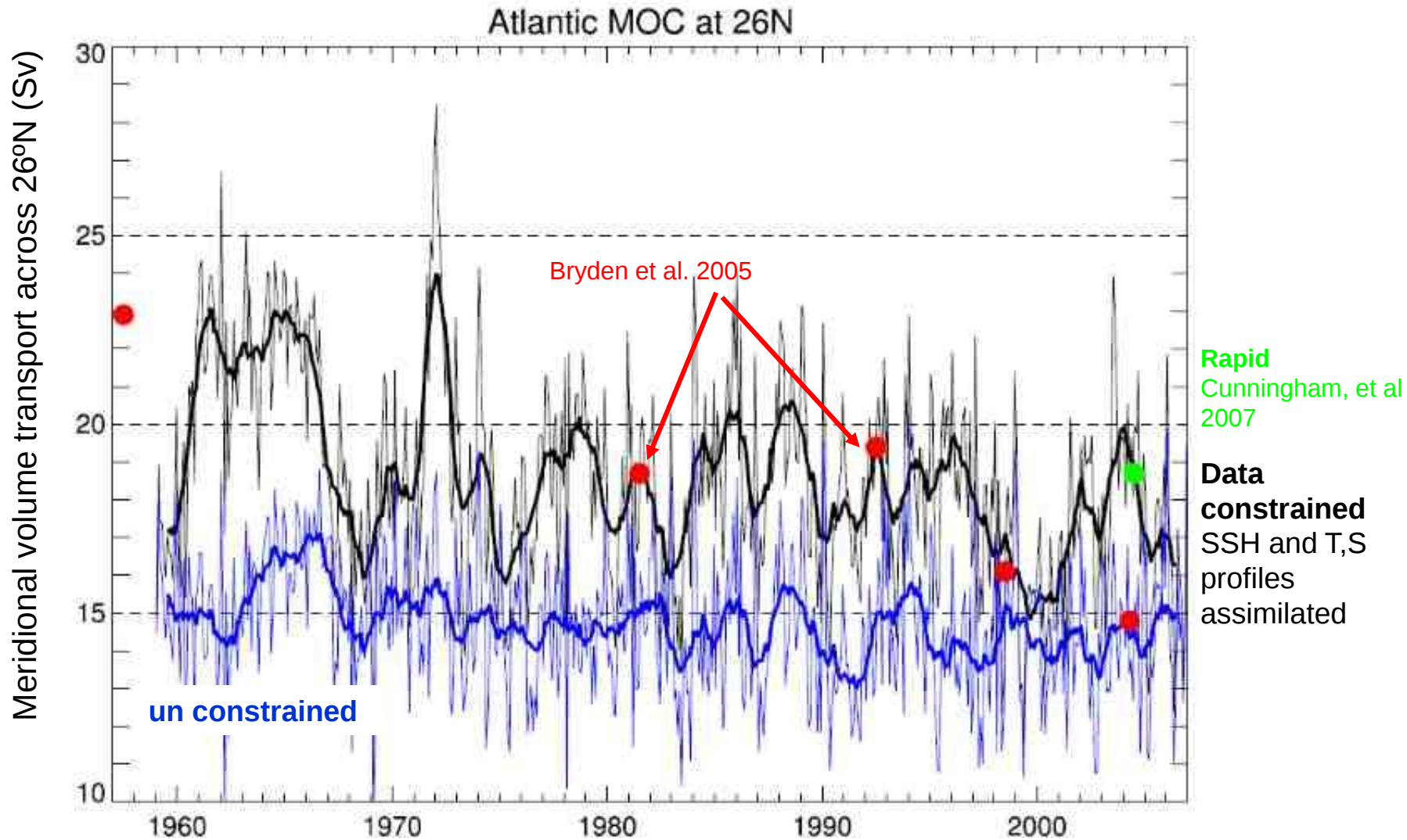


Observed Atlantic overturning – I – variability

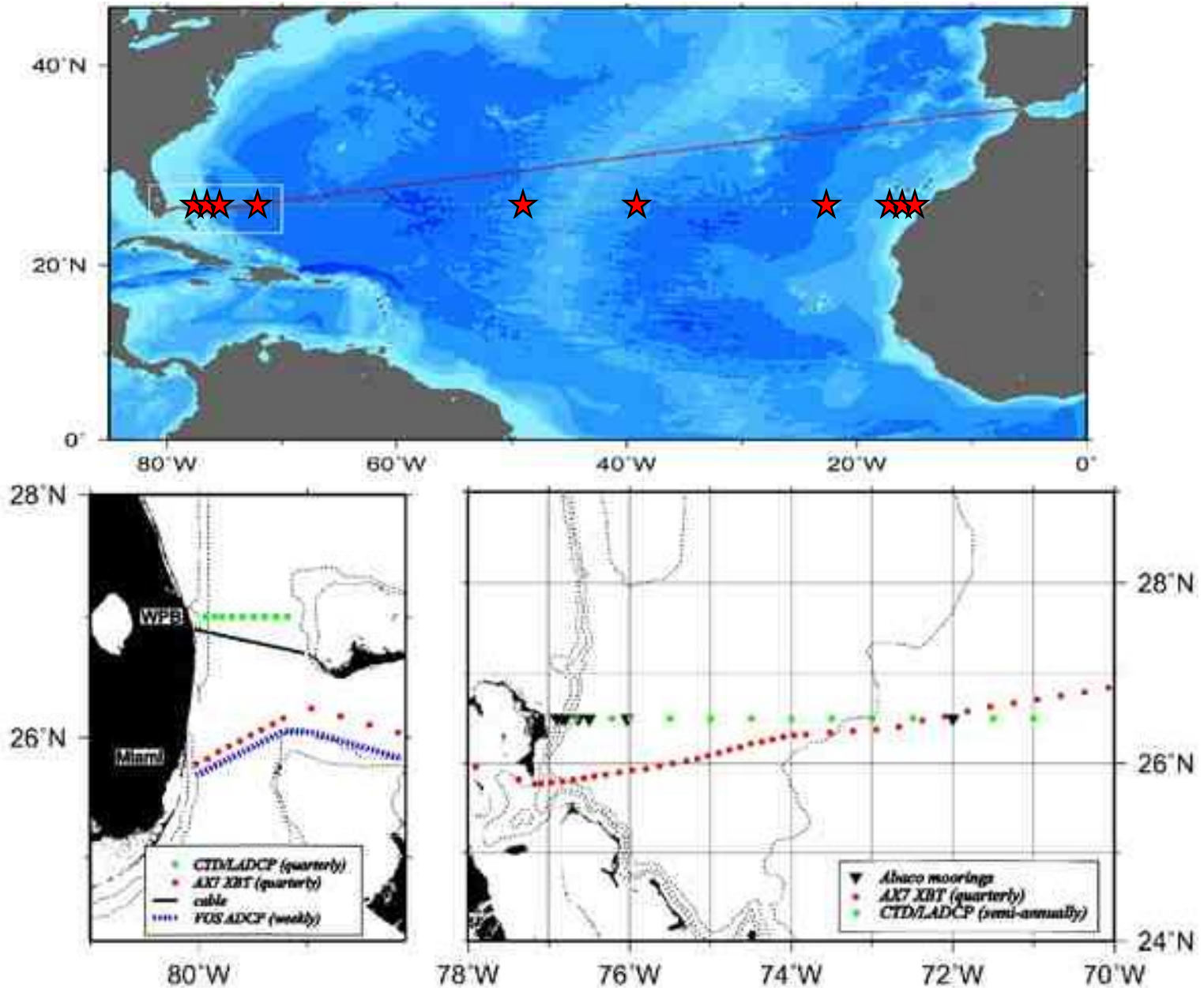


Atlantic overturning – III – variability

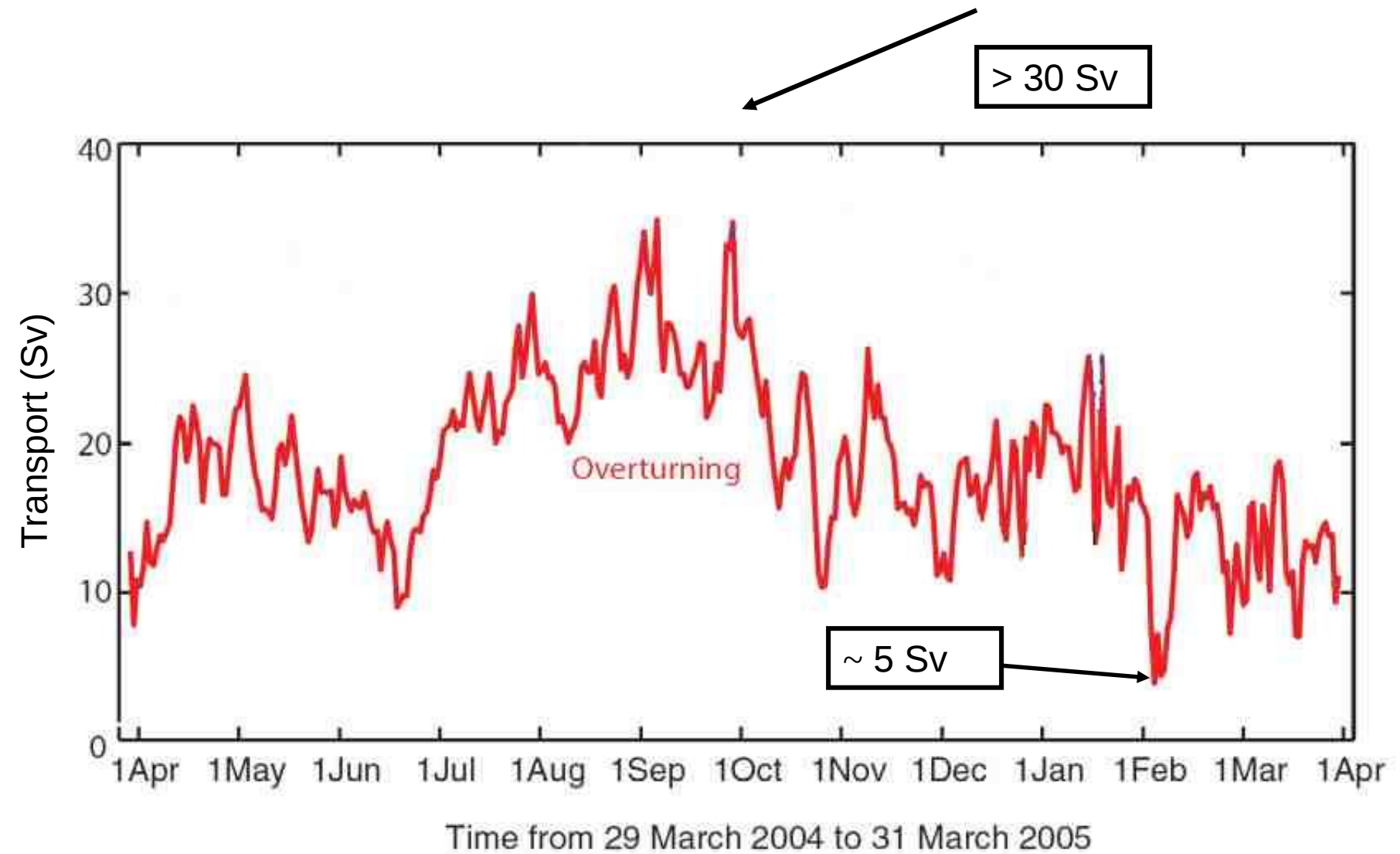
Contrasting models and observations



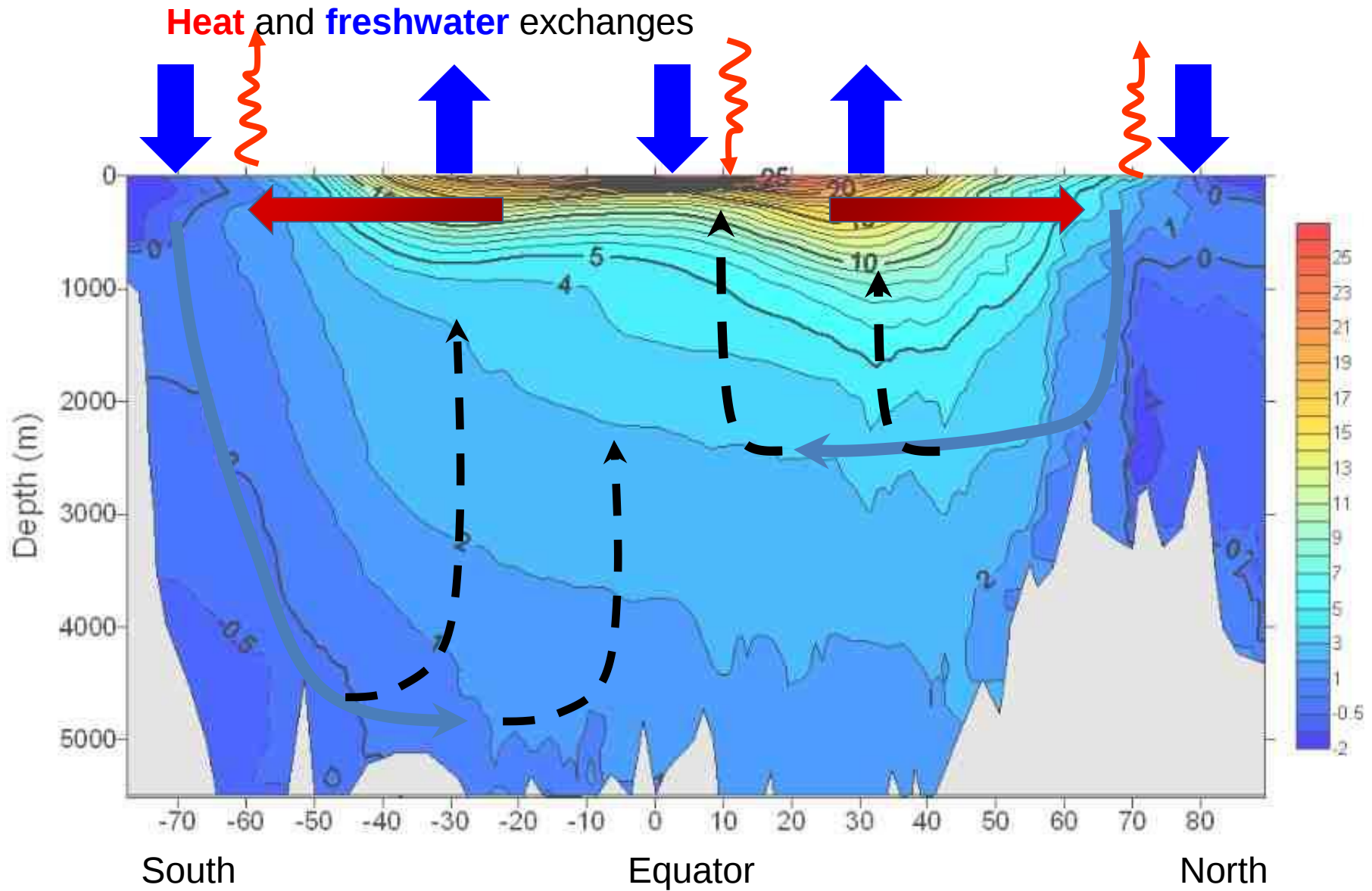
Observed Atlantic overturning – I



Observed Atlantic overturning – II

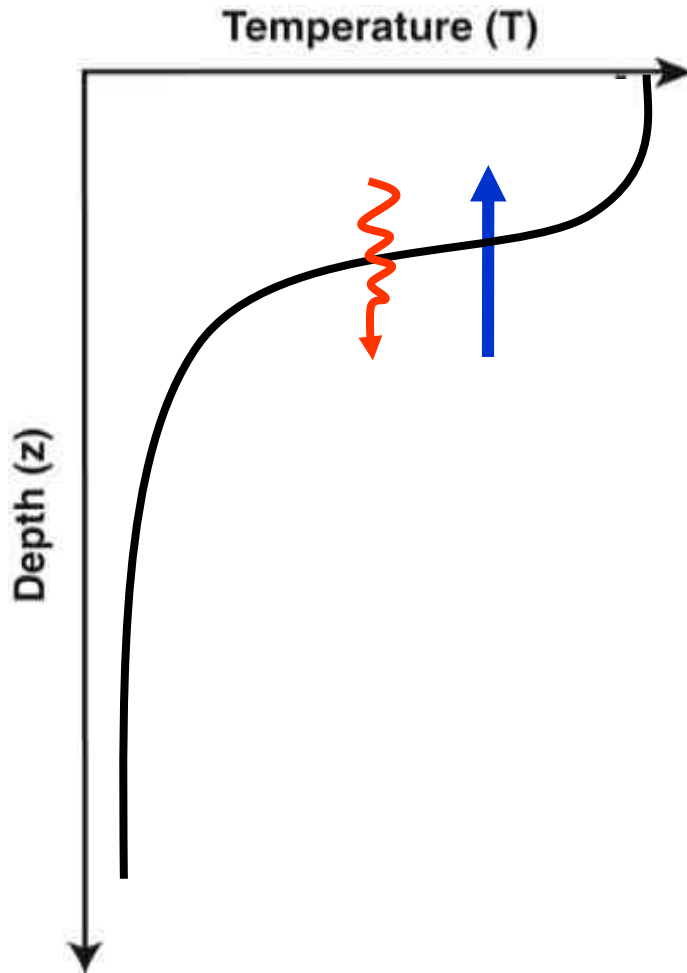


Overturning *a la* Rumford



Sustaining the vertical structure

(and the MOC)



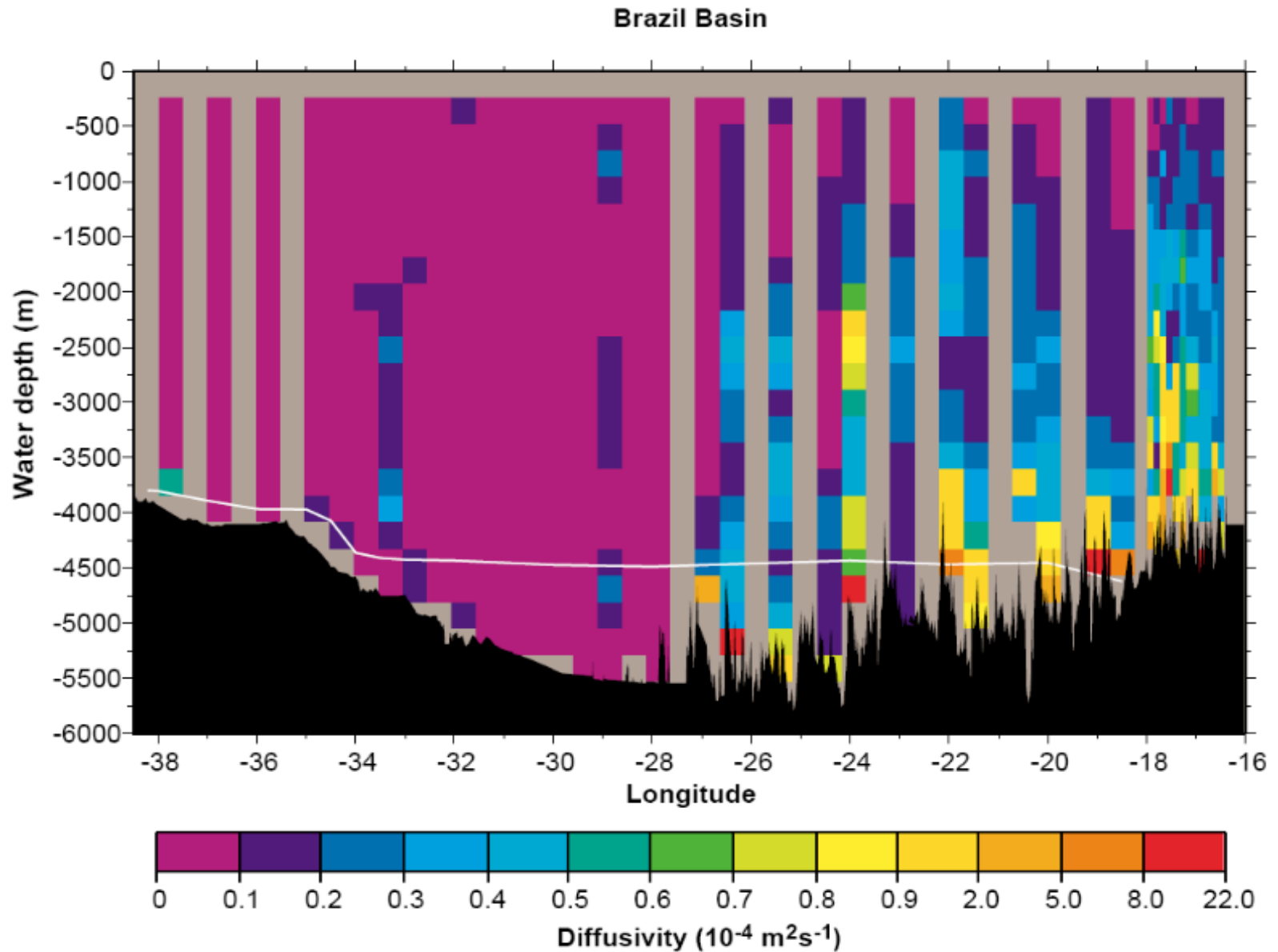
$$w \frac{\partial T}{\partial z} = \kappa \frac{\partial^2 T}{\partial z^2}$$

Downward turbulent diffusion of heat (RHS) is compensated by upwelling of cold water (LHS).

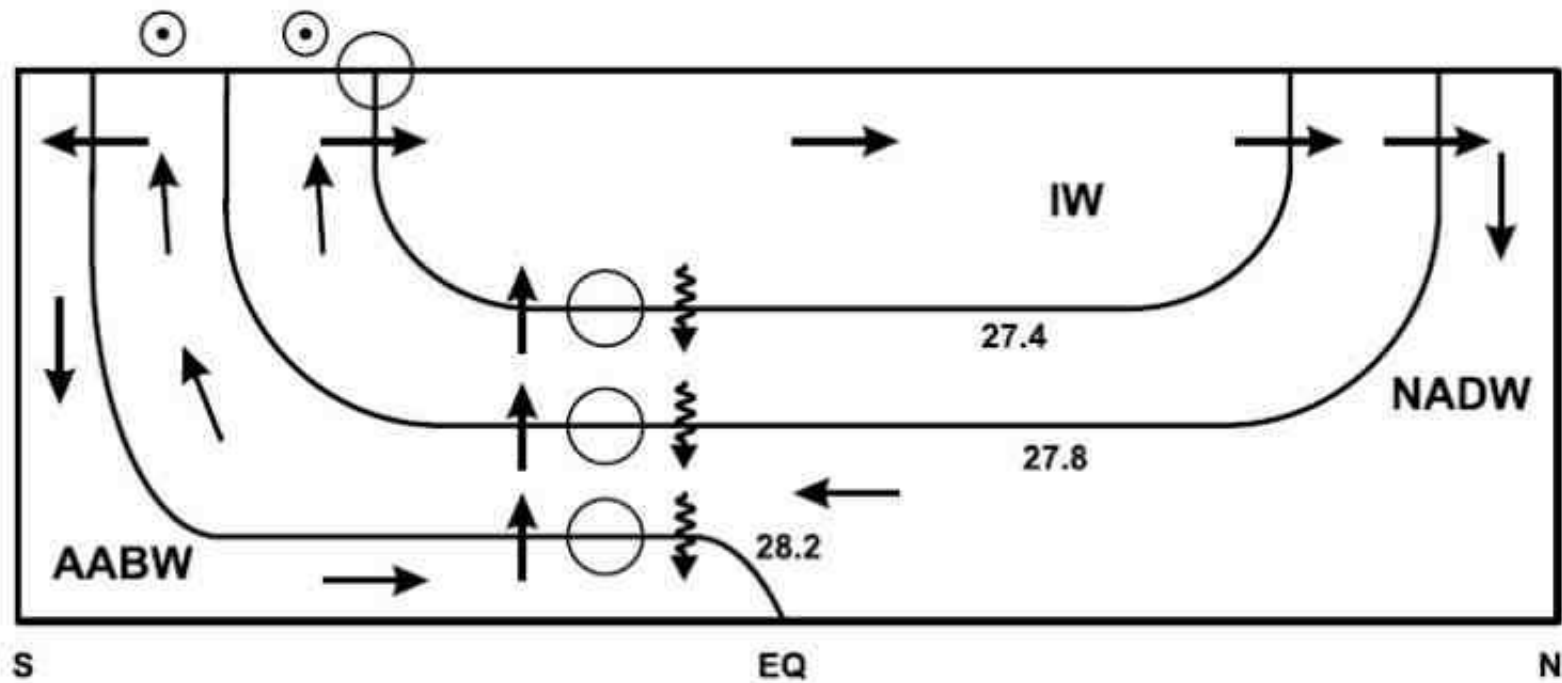
Globally averaged turbulent heat diffusivity (κ) needs to be $\sim 10^{-4} \text{ m}^2 \text{ s}^{-1}$ or about three orders of magnitude larger than molecular diffusivity.

This balance provides the upwelling necessary to balance water formation at high latitude.

Non homogeneous diffusivities



Schematic overturning modes - II



Non homogeneous diffusivities

